



Techniques of Water-Resources Investigations of the United States Geological Survey

Chapter C3

COMPUTATION OF FLUVIAL-SEDIMENT DISCHARGE

By George Porterfield

Book 3 APPLICATIONS OF HYDRAULICS

UNITED STATES DEPARTMENT OF THE INTERIOR

CECIL D. ANDRUS, Secretary

GEOLOGICAL SURVEY

V. E. McKelvey, Director

First printing 1972

Second printing 1977

UNITED STATES GOVERNMENT PRINTING OFFICE, WASHINGTON: 1972

.

For sale by the Branch of Distribution, U.S. Geological Survey, 1200 South Eads Street, Arlington, VA 22202

PREFACE

The series of manuals on techniques describes procedures for planning and executing specialized work in water-resources investigations. The material is grouped under major subject headings called books and further subdivided into sections and chapters; Section C of Book 3 is on sediment and erosion techniques.

The unit of publication, the chapter, is limited to a narrow field of subject matter. This format permits flexibility in revision and publication as the need arises.

Provisional drafts of chapters are distributed to field offices of the U.S. Geological Survey for their use. These drafts are subject to revision because of experience in use or because of advancement in knowledge, techniques, or equipment. After the technique described in a chapter is sufficiently developed, the chapter is published and is sold by the U.S. Geological Survey, 1200 South Eads Street, Arlington, VA 22202 (authorized agent of Superintendent of Documents, Government Printing Office).

III

CONTENTS

•

	Page
Preface	III
Abstract	1
Introduction	1
Types of records	1
Checklist for daily records	2
Particle-size analysis	2
Evaluation of size data	2
Tabulation of size data	4
Water temperature	5
Suspended-sediment concentration	9
Adequacy of data	11
Relation between single-vertical and	
cross-sectional concentrations	11
Cross-section coefficient	12
Variation with time	16
Analysis of cross-section concentration	
data	16
Development of a temporal concentration	
graph	17
Plotting symbols and scales	18
Theoretical considerations	19
Study of past records	19
Relation of water discharge to concen-	
tration	19
Estimates for periods of missing data	20
Visual comparison with adequately	
defined concentration graphs	21
Hydrographic comparison with rec-	
ords of upstream and downstream	
stations	21

Development of a temporal concentration graph—Continued
Estimates for periods of missing data— Continued
Water-sediment relation curves Examples of the sediment-concentration graph
Snowmelt discharge and sediment con- centration
Application of cross-section coefficient
Computation of daily mean concentration
Footnotes
Significant figures
Computer programs
Format of sediment tables
Computation of sediment discharge
Units of measurement
Computation of subdivided days
Mean-interval method
Midinterval method
Sediment-discharge worksheet
Station analysis
Station description heading
Periodic observations
Checklist for periodic records
Combined periodic and seasonal obser-
vations
Transmittal of completed data
Selected references

FIGURES

4		Page
1.	Graph showing minimum number of bottles of sample required to yield sufficient sediment	
9_7	for size analysis Form for—	3
4-1.		
	2. Annual tabulation of particle-size analyses of suspended sediment	4
	3. Maximum and minimum daily water-temperature tabulation	6
	4. Tabulation of once-daily water temperatures	7
	5. Sediment-concentration notes (short form)	10
	6. Tabulation of sediment data in the cross section	12
	7. Tabulation of sediment data in the cross section	13
8-12.	Graphs showing—	
	8. Relation of cross-section coefficient to discharge and season for San Joaquin River	
	near Vernalis, Calif	14
	9. Water discharge, sediment concentration, and coefficients for correcting observer's	
	single-vertical samples to cross section	15
	0 ·····	

CONTENTS

10-27.	Graph	s Showing—	Lage
	· ·	Relation of cross-section coefficient to gage height	16
		Advanced concentration during excess runoff periods	22
		Simultaneous concentration during excess runoff periods	22
		Lagging concentration during excess runoff periods	23
	14	Sediment-transport curve on a storm basis with indicated mean concentration	24
		Cumulative unit relation of total water discharge and total sediment discharge for	25
	16.	typical advanced, simultaneous, and lagging types of concentration graphs Typical effect of high-intensity short-duration rainfall on discharge and concentration for a small-drainage-basin stream having a very small amount of base flow or none	25 26
	17.	Gage height and sediment concentration, Corey Creek near Mainesburg, Pa.	27
	18.	Effect of two different flow conditions on discharge and concentration for the Rio Grande near Bernalillo, N. Mex	28
	19.	Gage height and sediment concentration, Colorado River near San Saba, Tex., May 1-6, 1952	29
	20.	Gage height and sediment concentration, Colorado River near San Saba, Tex., August 13-17, 1951	30
	21,	Gage height and sediment concentration, Colorado River near San Saba, Tex., May 22-27, 1951	31
	22.	Gage height and sediment concentration, Colorado River near San Saba, Tex., June 11-14, 1951	32
	23	Gage height and sediment concentration, Susquehanna River at Harrisburg, Pa	34
		Suspended-sediment concentration, sediment discharge, and water discharge, Willa- mette River at Portland, Oreg., December 21-30, 1964	35
	25.	Temporal relation of sediment concentration to water discharge during a snowmelt period	36
	26.	Relation of water discharge to sediment concentration, Green River at Green River, Utah, 1951	37
	27.	Graphical adjustment of concentration	39
28.		ams showing determination of mean concentration by graphical method	40
		showing format table—	
		1, suspended-sediment discharge	44
		2, suspended-sediment discharge for selected days	45
		3, suspended-sediment discharge measurements	45
		4, suspended-sediment discharge during periods of high flow	45
		5, total sediment discharge	45
		6, particle-size distribution of suspended sediment	46
		7, particle-size distribution of surface bed material	47
36		showing guide to subdivision, assuming accuracy about 5 percent	48
	-	am showing gage height and sediment concentration for a subdivided day	53
		heet for annual suspended-sediment discharge	54
		showing relation between daily suspended-sediment discharge and water discharge,	-
		nomas Creek at Paskenta, Calif., 1963-65 water years	56
40-45		t and content for—	00
-10-40.		Station analyses of water, sediment, temperature, and rainfall	58
		Station analyses of chemical quality and sediment	59
		Station analyses of sediment	
	42.	Heading and chemical-quality tabulation in annual water-quality-data report	
	43.	Updated station heading	62
	44. 12	Tabulation of periodic sediment data	63
	40.	LADUIANON OF PERIOUS SCUMENT UAVA	50

TABLES

	Page
1. Temperature conversion table to nearest 0.5 degree	8
2. Conversion factors, C, for sediment concentration: parts per million to milligrams per liter	43
3. Computation of subdivided day, mean-interval method	50
4. Computation of subdivided day, midinterval method	51

COMPUTATION OF FLUVIAL-SEDIMENT DISCHARGE

George Porterfield

Abstract

This report is one of a series concerning the concepts, measurement, laboratory procedures, and computation of fluvial-sediment discharge. Material in this report includes procedures and forms used to compile and evaluate particle-size and concentration data, to compute fluvial-sediment discharge, and to prepare sediment records for publication.

Introduction

Collection, computation, and publication of fluvial-sediment and related environmental data are part of a national program to evaluate effects of sedimentation on the life and economics of projects related to navigation, flood control, transportation, reclamation, water supply, recreation, pollution, and fisheries. Fluvial-sediment investigations may include determination of the sediment discharge of rivers, surveys of reservoirs, studies of channel morphology, research in basic processes, and interpretation of sediment data.

The purpose of this chapter is to combine into a single handbook the necessary information to evaluate sediment data, compute sediment discharge, and tabulate the data for publication. The content is based not only on the author's experience but includes information from the voluminous literature accumulated during the past two or three decades as well as the ideas of many experienced coworkers.

Although this chapter is limited to methods of compilation, computation, and editorial format, it also includes reference to sampling techniques, laboratory procedures, principles of sediment transport, and quality control, because knowledge of these is fundamental to computation of sediment records. The entire operation, from the collection of the sample in the field to the laboratory analysis and the computation and publication of the records, requires a high degree of coordination. Minor duplication of material in other chapters of the manual is necessary and intentional to allow use of the chapter as separate entities.

This manual was prepared by the California district, Water Resources Division, U.S. Geological Survey, Menlo Park, Calif., under the general supervision of R. Stanley Lord, District Chief. Technical advice and assistance were given by Geological Survey personnel in California, Texas, New Mexico, and Pennsylvania districts and by F. C. Ames, H. P. Guy, and J. K. Culbertson.

Types of Records

Two basic types of sediment records daily and periodic—are published by the Geological Survey.

Daily records are prepared for sites where sufficient determinations of sediment concentration and water discharge are obtained to justify computation of daily sediment discharge. The end product is a tabulation of daily mean concentration, suspended-sediment discharge, and periodic determinations of particle-size distribution of suspended sediment and bed material. These are combined with other quality-of-water data and released, usually by water year (October through following September) and on an annual basis, by the Geological Survey in basic-data reports covering a specific State or in the water-supply-paper series "Quality of Surface Waters of the United States."

Periodic records are prepared for sites where determinations of concentration and water discharge are not sufficient to justify computation of daily sediment discharges or where only miscellaneous samples are obtained. In addition to publication of the records, the data and computations are maintained on file in the district offices of the Water Resources Division and are available for examination or for use in interpretative reports or research.

Checklist for Daily Records

Steps in the procedure for the computation of daily records of fluvial-sediment discharge are given in the following checklist. A checklist for periodic and mixed records is given in the section "Periodic Observation" of this report. Data on stream stage and discharge needed in the daily sediment computation may be obtained from an A-35-analog-recorder chart or a plot of bihourly gage heights or discharge and from data forms 9-192, 9-210, and 9-207. The checklist items are as follows:

Particle-size analyses: Compute from laboratory analyses Tabulate Apply instantaneous water discharge Tabulate water temperature Sediment concentration: Compute from laboratory analyses List sediment measurements Copy size-concentration values on concentration notes **Compute coefficients** Chart computations: Plot concentration Draw concentration graph Review concentration graph Compute concentration Apply concentration coefficient Compute subdivided days Check subdivided days Sediment-discharge worksheet: Copy water discharge Copy concentration Compute sediment discharge Compute totals Check totals

Sediment-discharge worksheet-Continued Compute maximum and minimum Insert footnotes Plot sediment-transport curve Plot hydrograph Write or update station description Write station analysis Review entire record Prepare copies for records-processing center

Particle-Size Analysis

Samples of suspended sediment from each sampling site taken at specific or selected times of the year are analyzed for particlesize distribution. These samples indicate the average particle-size distribution of the material transported and should be obtained at various seasons of the year and at sufficient increments of discharge to cover the complete range in seasonal flow. Samples of bed material also are obtained to define the size distribution of bed material at various increments of flow and to define the physical properties of the material available for transport. The type and purpose of the sample dictates to some degree the sampling procedure, the methods used to analyze the sample, and the methods and forms used to present the data. Sampling procedures are discussed in detail in the manual on field methods for fluvial-sediment measurements (Guy and Norman, 1970) and by the U.S. Inter-Agency Committee on Water Resources (1963). Laboratory procedures and methods of analyses are discussed in detail in the manual on laboratory theory and methods for sediment analyses (Guy, 1969).

Evaluation of size data

Particle-size analyses should be evaluated during the review and tabulation process for—

- 1. Correct method of analysis,
- 2. Total number of analyses,
- 3. Range of water discharge,
- 4. Agreement of concentration and water discharge on the particle-size tabulation with those on the sediment-discharge sheet and published records, of water discharge, and

5. Validity of percentage-finer values.

The number of samples analyzed, methods of sample collection, and method of analysis depend partly on the purpose and scope of the sediment project or data program; results of analyses should be reviewed to determine if the number of samples and method used for analysis fulfills the goals of the sampling program.

Accuracy of the analysis is dependent, among other factors, on the quantity and physical characteristics of the sediment analyzed; an inspection of the data sheet will indicate if a sufficient quantity of material was collected for analysis. Analysis made on samples containing insufficient material may be in error. Particle-size distribution of these samples should be carefully evaluated for accuracy, and if suspect they should not be published. In general, the quantity of sediment needed for analysis is as follows:

Method	Quantity of	sediment, in grams
	Minimum	Optimum
Dry sieve	50	100
Wet sieve	.05	0.5-1.0
VA tube	.05	1.0-7.0
Pipet	.8	3.0-5.0
BW tube		0.7-1.3

The minimum number of bottles of sample required to provide sufficient sediment for size analysis may be determined from the curves in figure 1. The range of concentration values and percentage finer than 62 microns needed to use figure 1 are available in the station records for the preceding year.

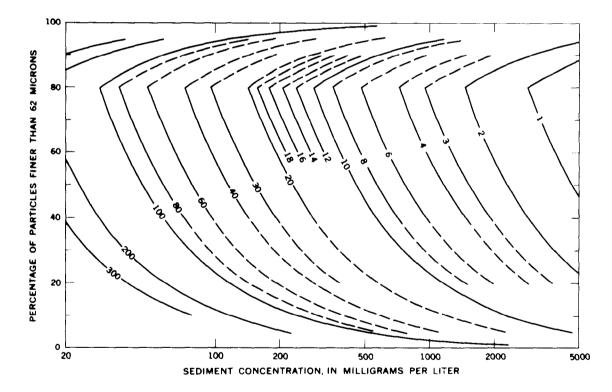


Figure 1.—Minimum number of bottles of sample required to yield sufficient sediment for size analysis. Explanation: Estimate sediment concentration and percentage finer than 62 microns by referring to analysis of samples obtained previously or by visual examination of the sediment sample. The number of bottles required is the value indicated by the line to the left of the intersection of the ordinate and abscissa. Interpolation of number of bottles is made along the abscissa. Values of the number of bottles required were computed on the assumption that each sample bottle contains 350 grams of water-sediment mixture and that a minimum of 0.2 gram of sand for a sieve or visual-accumulation-tube analysis and 0.8 gram of silt and clay for a pipet analysis in 400-milliliter suspension are needed for analysis. The number of bottles required to yield sufficient silt and clay for a bottom-withdrawal-tube analysis is five-eighths of the number indicated.

Tabulation of size data

Particle-size analyses of suspended sediment are tabulated on form 9–1539D, shown in figure 2. This form is also used to tabulate periodic or miscellaneous concentration and particle-size data. Particle-size analyses of bed material are tabulated on form 9–1539E (not illustrated). Examples of offset copy furnished by computer for publication are illustrated in the section on "Format of Sediment Tables."

Instantaneous water discharge at the time of sampling and concentration of the sample analyzed for particle-size distribution are determined and tabulated for each sample. These values must be compared for validity with the daily values published for water discharge in the surface-water records and on the sediment-discharge sheet. Water-discharge values may be computed and listed in the space provided on the particle-size forms or may be listed on the multiple-purpose form described in the section on "Analysis of Cross-Section Concentration Data."

Data from the particle-size analyses should be transcribed neatly on the form shown in figure 2. The data are arranged on the form as follows:

Date of collection.—Tabulate year, month, and day. January 1, 1970, for example, is 700101.

Time (24 hour).—Time is reported in 24hour local standard time. The hours and minutes are always written to four places and without punctuation. Do not use a.m. or p.m. For example: 0001 hours is 1 minute after midnight; 0100 hours is 1 a.m.; 1048 hours is 10:48 a.m.; 1200 hours is 12 m. (noon); 1430 hours is 2:30 p.m.; and 2400 hours is 12 p.m. (midnight).

Water temperature ($^{\circ}C$).—Water-temper-

Form
9-15390
(10-69)

UNITED STATES DEPARTMENT OF THE INTERIOR - GEOLOGICAL SURVEY - WATER RESOURCES DIVISION Coding Form for Instantaneous Suspended Sediment and Particle Size

Type Station ident. number							,							
[1] 000000×0	o Stream and <u>San</u> location <u>San</u>	mple Ri	ver	nea	11.	Som	e wh	ere	4.0	5 <u>A</u>				
Methods of analysis (74-77): E V. visual accumulation tube.	Methods of analysis (74-77): B, bottom withdrawal tube; C, chemically dispersed; D, decantation; N, in native water; P, pipet; S, sieve; W, in distilled water; , visual accumulation tube. Please do not punch the decimal point unless data appears in the field.													
	al point unless data appears													
Date Time Tempe		Sup. sediment concentration		· · · ·		<u> </u>		<u>.</u>			tize indi			Method
Yr. Mo. Day		(mg/1)	0.002	0 004	0 008	0 016	0 031	0 062	0 125	0 250	0 500	1 000	2 300	amalysis
69 10 219 08.40 14 ·			41 42 43	• -				100					<u>n_n_</u> //	5
69 10 29 08 40 14		1304					┟┈└─┶		┟╌┷┷		<u>}</u>			
1.1.10 99.40 .9.	5	4.7.8	63		7.9.	a.n.	9.8	9.8.	9.9	1				S,PW,C
1.17.17.0.49.40 17.	21	11111/18	<u>, 2171</u>	62	1/171	90	17.11.		[7.7.]	17.00			1.1	
1, 1, 2, 2, 0, 1, 3, 0, 0, 7.	0	3.05			r		T	7.4.	8.1.	9.0.	100			
1/12/210/13/010 17.0							4_4_4~	1/17L	لقىنى	mer.			<u> </u>	
1.2.2.1 1.1.0.5 .7.	2750	1,1,5,0,0	05	3,0,	3.7.	4.8.	6.0.	6.7.	8.6	94.	9.8	1.0.0		VIPING
1,2,2,1,1,05,7.	<u></u>		25	1210	121/1-	<u>462</u>	60		18 1601	1741		Loop		
12211410 81	0 18600	27200	0.0	24.	3.1.	42	5,5,	6,2,	8.4.	9.4.	9.8	1.0.0		VIPIWC
1/12/2/1/14/14	5 . 1.2699.	3,7,20,0	20	261	1211	5Z.	00	ei ^c i	U iTi.	$V_{1}T_{1}$	1101	7,0,0		VINUM
	5 1.2.8.0.0.	441.00	1.6	2.1.	2,7,	3.9	5.0	54	7.2.	07	95	97		SPINC
1,223 1.33,0/.0.	2 128,99	44100	10	21/1_	44	<u>7.7.</u>	13.0	Pier.	[141.		721	7.4	100	DUM
1228/1020 5.	0, 1990.	1.1.60.0	22	2.9	3.4.	5.2.	45.		00	9.5.	9.9.	100		VPWC
		111116010	6141	47	12171	<u> 221.</u>				1/121	1/1/2	100		
700/05/420 5	5 37.00	1,1,2,900	1.7	2.2.	3.1.	4.0	50	5.7.	7.6.	8.9	9.8	1.00		VIPIWC
POTOSITYTE PL		1 1/14/100			<u>1911 -</u>						tuar .	Trad		
01711025 6.	0, /3301	. 60.30	1.3.	2.3.	3.4.	4.5	5.6	65.	7.9.	90.	9.6.	1.00		VPWC
1 101/141/1/101-101 101									<u> </u>		11.161	1961		
0.2 2.7 1.2.20 .8.	5 8.30	1,3230	1.8	2.4	3,5,	4.7.	5,8,	64	8.2	96	100			V, P, WC
		P			10/9/			ked by			11 - 12	 		1/2/70
														1 =/ / -
	Punched by _			Date			Veri	fied by		····		_ Dat	ie	

Figure 2.—Annual tabulation of particle-size analyses of suspended sediment.

ature data are reported to the nearest 0.5 degree Celsius.

Number of sampling points.—Use when reporting bed-material analyses. This is the number of samples obtained in the stream cross section.

The laboratory may report individual analysis for each sample to show variation of bed-material size distribution in the cross section and to provide necessary data for computation of total sediment discharge by the procedure described by Colby and Hembree (1955). Generally the average size distribution in the cross section is published; however, individual analysis may be published if there is a large variation of median diameter among verticals or if there is a need for more detailed information.

Discharge (cfs).—The discharge is usually reported for the time of sampling; however, if no measurment is made or the rating does not justify reporting the instantaneous discharge, the daily mean discharge is reported. If mean discharge is used, change heading to "Mean discharge (cfs)" or use footnote "D" as explained in the section on "Footnotes."

Concentration (mg/l).—All sediment concentrations will be reported in milligrams per liter although they will be determined in the laboratory as parts per million. The supervisor must determine that all concentrations have been properly converted prior to tabulating concentrations or computing sediment discharge.

Sediment discharge (tons per day).— Values for tons per day should be determined as discussed in the section on "Computation of Sediment Discharge." Records tabulated for computer processing do not require computation of sediment discharge because values in this column are computed and listed by computer.

Particle sizes, percent finer than size (in millimeters) indicated.—The sizes for suspended sediment are 0.002, 0.004, 0.008, 0.016, 0.031, 0.062, 0.125, 0.250, 0.500, 1.00, 2.00 mm. The sizes for bed-material analyses are 0.004, 0.062, 0.125, 0.250, 0.500, 1.00, 2.00, 4.00, 8.00, 16.00, and 32.00 mm. Method of analysis.—In the method of analysis column, the symbols should be recorded in the same order that the methods were used for the analysis. For example:

SBWC
VPWC
SPWC
VPN
SV

The symbols are explained in the headnotes and are standard.

Water Temperature

Temperature is an important physical characteristic of water, and information on water temperature is a necessary part of any study of water quality. It is also an important parameter needed to compute total-sediment discharge. A temperature observation should be obtained with each chemical-quality or sediment sample.

A temperature record may consist of a tabulation of maximum and minimum observations, once-daily observations, or observations obtained during periodic visits to a station. A continuous temperature record may be obtained from one of the many devices that sense and record fluctuations of water temperature on a continuous chart. Maximum and minimum temperatures for each day are computed from the chart and listed as illustrated in figure 3. A tabulation of once-daily temperature observations obtained by field personnel or contract observers is illustrated in figure 4. The form shown in figure 4 is a modified 9-211C.

Those observations taken daily or more frequently will be included in the tables of annual reports; observations obtained at infrequent visits to the station will be published in conjunction with other data, such as the tabulation of particle-size data on form 9-265b (fig. 2) or periodic sediment data (fig. 45). The completed tabulations (fig. 3 or 4) including maximum and minimum values are sent to the records-processing center where they are prepared for publication. The tabular data must be complete, that is with an entry in each space. If no

ampleville, water year October /969 to September /970	March April May June July August September	Max Min	6.0 4.0 9.5 6.0 12.5 8.5 23.0	5.5 4.0 8.0 5.5 13.0 8.0 23.5 16.5 22.5 17.0 24.0 17.5 23.0	5.0 3.5 7.0 4.5 10.5 8.0 24.0 17.0 23.0 17.5 24.5 175 22.5	6.0 3.5 8.0 6.0 13.0 7.5 235 175 230 175 245	6.0 4.5 8.0 6.0 16.0 10.0 295 17.0 235 18.0 24.9 17.0 23.0	6.5 4.5 8.5 6.5 18.0 12.5 19.5 16.0 23.5 17.5 24.0 17.0 24.0	75 5.0 10.0 6.5 17.5	7.5 7.5 12.0 9.0 18.0 15.0 17.0 15.0 25.5 18.0 24.0 17.0 25.0	6.5 5.0 10.5	7.0 4.5 124 8.5 17.0 14.5 17.5 15.5 235 180 23.0 175 235 16.0	6.0 3.5 12.5 12.0 200 15.0 18.0 15.5 22.0 18.0 230 17.5 22.0	8.0 4.0 11.5 9.0 18.5 15.5 20.0 15.5 25.0 17.5 23.5 17.5 22.0	8.5 4.5 9.5 7.5 15.5 13.5 20.5 16.0 250 17.5 250 16.5 20.5	75 4.0 9.0 8.0 18.0 13.0 18.5 160 250 17.0 24.5 17.0	9.0 5.0 11.0 8.5 18.0 14.5 14.5 16.5 24.9 10.0 24.5 17.0 21.5
year			5 23.	53	24	23	1							202	20.	0	й К
vater	lay		5	8	00	2.5			ž L		2				1 /33		*
-	Σ	Max	12.	Э.С	Ó.	Э.	2			Ó.	2	I		Ś		Ś	100
	ril			_	¥	ġ	ف	ý	ف			00	á	Ø.	ĸ		
	Ap	Max	o.	ø	N			<i>ø</i> .		ñ	Ó.	2			9.5		
rille,	rch	Min	•.▲	4.0			4.5				•		Ø,	•	K .	4	
mplev	Ma	Max															- 1
at Se	February	Min	2.5		W,		4.0			51		6.0	5.5				5.0
- u	ā	×	D	0	5.5	Ó	0	5	0	6	5	7.0	22	Ò	'n	5.5	6.0
ive	Fe	Max					5										- 1
ple River		Min	5.5 4	6.0	6.5	6.9	6.5	6.5	5.5	₹.0	2.0	€.0	50	5.5	6.0	5.5	50
Sample River at Sampleville,	January Fe		6.5 5.	70 6.0	6.5 6.5	7.0 6.5	75 6.5	7.5 6.5	7.0 5.5	5.5 4.0	4.0 2.0	5.0 4.0	5.5 5.0	7.0 5.5	75 60	6.0 5.5	5.5 5.0
	January	Min	6.5 5.	6.0	5.0 6.5 6.5	5.0 7.0 6.5	7.5 7.5 6.5	6.0 7.5 6.5	6.5 7.0 5.5	70 55 40	9.0 4.0 2.0	8.0 5.0 4.0	50	5.5	75 60	5.5	5.5 5.0
	December January	Max Min	6.5 5.	70 6.0	5.0 6.5 6.5	7.0 6.5	7.5 7.5 6.5	6.0 7.5 6.5	6.5 7.0 5.5	9.0 7.0 5.5 4.0	10.0 9.0 4.0 2.0	11.5 8.0 5.0 4.0	8.0 6.0 5.5 5.0	6.5 5.0 7.0 5.5	70 60 75 60	75 65 6.0 5.5	7.5 6.5 5.5 5.0
	December January	Min Max Min	5.0 6.5 5.	11.0 6.0 5.0 7.0 6.0	11.0 6.0 5.0 6.5 6.5	11.0 7.0 5.0 7.0 6.5	11.5 8.5 7.5 7.5 6.5	10.0 7.5 6.0 7.5 6.5	11.0 7.0 6.5 7.0 5.5	70 55 40	130 100 9.0 4.0 2.0	13.5 11.5 8.0 5.0 4.0	12.5 8.0 6.0 5.5 50	90 65 50 70 55	80 70 60 75 60	7.0 7.5 6.5 6.0 5.5	75 75 65 5.5 50
	January	Max Min Max Min	125 10.5 6.5 5.0 6.5 5.	13.5 11.0 6.0 5.0 7.0 6.0	12.5 11.0 6.0 5.0 6.5 6.5	12.0 11.0 7.0 5.0 7.0 6.5	13.0 11.5 8.5 7.5 7.5 6.5	11.5 10.0 7.5 6.0 7.5 6.5	14.0 11.0 7.0 6.5 7.0 5.5	14.5 12.5 9.0 70 5.5 4.0	16.0 13.0 10.0 9.0 4.0 2.0	16.0 13.5 11.5 8.0 5.0 4.0	14.0 12.5 8.0 6.0 5.5 50	12.5 9.0 6.5 5.0 7.0 5.5	90 80 70 60 75 60	8.0 7.0 7.5 6.5 6.0 5.5	8.0 75 7.5 6.5 5.5 5.0
	November December January	Min Max Min Max Min Max Min	16.0 12 5 10.5 6.5 5.0 6.5 5.	14.0 13.5 11.0 6.0 5.0 7.0 6.0	14.5 12.5 11.0 6.0 5.0 6.5 6.5	15.0 12.0 11.0 7.0 5.0 7.0 6.5	15.0 13.0 11.5 8.5 7.5 7.5 6.5	135 11.5 10.0 7.5 6.0 7.5 6.5	135 14.0 11.0 7.0 6.5 7.0 5.5	12.0 14.5 12.5 9.0 7.0 5.5 4.0	11.0 16.0 13.0 10.0 9.0 4.0 2.0	14.5 16.0 13.5 11.5 8.0 5.0 4.0	15.0 14.0 12.5 8.0 6.0 5.5 5.0	135 125 9.0 6.5 5.0 7.0 5.5	120 90 80 70 60 75 60	11.0 8.0 7.0 7.5 65 6.0 5.5	12.0 8.0 75 7.5 6.5 5.5 5.0
Temperature (^o C) of water, Sample Rivel	December January	Max Min Max Min Max Min	0125105 6.5 5.0 6.5 5.	14.0 13.5 11.0 6.0 5.0 70 6.0	14.5 12.5 11.0 6.0 5.0 6.5 6.5	15.0 12.0 11.0 7.0 5.0 7.0 6.5	15.0 13.0 11.5 8.5 7.5 7.5 6.5	11.5 10.0 7.5 6.0 7.5 6.5	14.0 11.0 7.0 6.5 7.0 5.5	14.5 12.5 9.0 70 5.5 4.0	16.0 13.0 10.0 9.0 4.0 2.0	16.0 13.5 11.5 8.0 5.0 4.0	14.0 12.5 8.0 6.0 5.5 50	5/25 9.0 6.5 5.0 7.0 5.5	120 90 80 70 60 75 60	0 20 7.0 7.5 65 6.0 5.5	12.0 8.0 75 7.5 6.5 5.5 5.0

Department of the Interior - Geological Survey - Water Resources Division ŝ

9-267d

Figure 3.—Maximum and minimum daily water-temperature tabulation.

TECHNIQUES OF WATER-RESOURCES INVESTIGATIONS

Q 22

* 5 6 1 1

22.0

24.5 170 24.0 15.0 24.0 17.5 23.0 18.0 23.0 18.0 23.0 18.0 25.0 18.0 25.0 18.0 25.0 17.0 24.0 17.0 24.0 17.0 24.0 17.0 24.0 17.0

24,0 17.0 24,0 17.0 24,5 17.0 24,5 17.0 21,5 1

13.0

00505

00 6664

505

00 5

19.5

16.0 15.5 15:0 Ó

21.5 21.5 10.0 20.0

60201100

* *

Ň

170 12. 185 13. 185 13. 175 12. 175 13. 180 12

000000

00

0

4.5.4 4.5.6 4.5 7.5 7.5 7.5 7.5

0.5 50

3.0

6.0 5.0

N.4.1 N.0.12

10.20

0.990 0.990 0.990

NONN

5

¥

0

N,

ا فإني فإني في

505

\$ E & E У

23.0 00

00

0

Ŕ

21.50

23

25.0 55

25 NN

24.0

9

05% 06

000 Ø

000

135

0.0/ 2.0/ 2.0/ 2.0/ 000

> ۱. 11

00 ١ 11

Wir.

2001

8 à à à î

oh 0

17.5

Ń í. ľ.

20.02

15.0 14.5 0 ŘI ė

13.5

Ś

0

Ó.

Ľ

5 0

1

 $U(S,Department of the Interior - Geological Survey - Mater Resources <math display="inline">D_i v_{1,31011}$

Ì

water year October 1969 to September 1970 Sample Creek near Sampleville, Temperature (^oC) of water,

1

!

j

ļ

Į			Ō	Dnce-daily	observation	on between	1600 and	1800 hours	PST]			
D _{ay}	0et.	Nov.	Dec.	Jan.	Peb.	Mar.	Apr.	May	June	July	Aug.	Sept.
1	0.8/	165	001	10.5	7.5	0 8		105	16.5	195	0.0	
Ċ1	175			10.5	08	75		11.0	110	0 []) () ())
n	17.0		00/	0 //	0 8	0 2	0 .>~	105	165		2 C	01.
4	16.5	17.0	ι 1	10.5	5	65		0.11	0 %		0 (0 - 7 (5
LC;	17.0	0 1 1	0 I I	0.11	. 0) 、			, , , , , , , , , , , , , , , , , , ,	Ç 0 7		20.0
2	175				5	e		12.0	/2.0	200	200	19.0
0 1		j e	01	n .//	а 0 i	01		125	14.5	20.5	210	081
	c · / /	0.1	115	0 11	0	ر م	00/	/3.0	14.5	205	200	2 2
20	0.61	175	120	0 01	9 21	~	10.01	12 0	14.0	21.5	200	19.5
9	0 1/	081	12.5	95	0 0/	8 S	105	125	0 11	215	19.5	8 5
	17.0	17.5	02/	10.0	95	90	011	13.0	13.5	20.5	200	2.6
;;	v n 	110	200	10.5	0.6	8	011	/3.0	14.5	19 5	21.0	061
N (0.0/	ø Ö	0	00/	12.5	100%	200	20.0	200
21 :	/e.0	201	200	0.0	01	, ė	9.0	12.5	165	200	510	200
			5 5	4 1	56	0	0.01	115	0.11	20.5	21.0	205
1	10.1	15.0	00/	95	06	6 0	105	13.0	081	20.5	20.5	061
16	100	14:0	9	06	0 %	0 e	1 5	/35	185	20.5	20 5	0.81
17		< 1	0.0/	0 0	مرہ ک	65	9.5	0 +1	185	21 0	0 77	0
<u>8</u>		04/	50/	0.	50	وبح	0.0/	115	175	21.0	0 5	75
19		0+1	0 0/	06	90	7.0	95	130	165	215	200	0.8
	011	2	6.6	95	95	2,0	0.//	14.0	190	215	19.5	0.21
21	. r	0.41	<i><i>o</i></i>	9.6	25	0	077	150	18.5	22.0	210	195
		2 4 C	0.4	01	1 00	1.5	0.0/	15.5	185	22.0	21.5	0.6
3 2	0.41	0		γ N	2	2.5	95	14.0	165	21.0	20 0	19.5
н м 9 С	17.5		0.00	0 0	5	ò ò	9.5	12.0	165	19.0	195	19.0
3	001	0,11	10.01	0.1	0,1	4.0	10.0	120	160	220	20.0	19.5
0.1	00	0.1	202	0.6	یں معر	9.5	105	// 5	15.5	22.0	195	190
77	2 4	200	<i>v i</i>	0 1	<u>م</u>	95	0.11	/3.0	15.5	21.0	200	18.0
0			0.0/	1 2	0	000	105	14.0	021	205	205	17.5
87	0.27	201	5.6	0.1	11	4 1 V 1	001	14.5	175	20.5	200	0.81
2	16.0		000	<i>o</i> (~ ~ ~	10.0	16.0	18.0	21.0	19.5	0.1
				2		2		16.5		210	195	

ι,

COMPUTATION OF FLUVIAL-SEDIMENT DISCHARGE

Figure 4.—Tabulation of once-daily water temperatures.

value is available, a leader (..) should be placed in the blank space.

Although temperature data have been published by the Geological Survey in degrees Celsius since October 1967, many temperature measurements are made with a Fahrenheit thermometer and converted to Celsius when recorded on the permanent laboratory and field sheets. Thermograph records may be converted directly to Celsius by template, or the values may be tabulated in Fahrenheit (figs. 3, 4) and the conversion to Celsius made by digital computer. Values recorded on digital tape by temperature monitors also may be converted by computer. All other temperature values, including the "extremes" values for the period of record, should be converted to Celsius (table 1) in

Table 1.—Temperature conversion table to nearest 0.5 degree

[The numbers in the center columns refer to temperatures, either in Celsius or Fahrenheit, which are to be converted to the other scale. If converting Fahrenheit to Celsius, the equivalent temperature will be found in the left columns. If converting Celsius to Fahrenheit, the equivalent temperature will be found in the right columns]

	0 to 24.	5	:	25.0 to 4	9.5	t	0.0 to 7	4.5		75.0 to 10	0.0
	0	32.0	-4.0	25.0	77.0	10.0	50.0	122.0	24.0	75.0	167.0
-10.0	.5	33.0		25.5	78.0	10.5	50.5	123.0	24.0	75.5	168.0
17.0	1.0	34.0		26.0	79.0	10.5	51.0	124.0	24.5	76.0	169.0
-17.0 -17.0	1.5	34.5	3.0	26.5	80.0	10.5	51.5	124.0 124.5	25.0	76.5	170.0
-17.0	2.0	35.5	<u> </u>	27.0	80.5	11.0	52.0	125.5	25.0	77.0	170.5
-16.5	2.5	36.5	-2.5	27.5	81.5	11.5	52.5	126.5	25.0	77.5	171.5
-16.0	3.0	37.5	2.0	28.0	82.5	11.5	53.0	127.5	25.5	78.0	172.5
-16.0	3.0 3.5	38.5	2.0	40.U	83.5	11.5	53.5	121.5	26.0	78.5	172.5 173.0
-15.0	3.5 4.0	39.0	<u> </u>	28.5 29.0	84.0	12.0	53.5 54.0	129.0	26.0	79.0	174.0
	4.0	40.0	1.5	29.5	85.0	12.0	54.5	130.0	26.5	79.5	175.0
-15.5 -15.0	4.5 5.0	40.0	1.0	29.5 30.0	86.0	12.5	55.0	130.0	26.5	80.0	176.0
-15.0 -14.5	5.0 5.5	41.0	1.0 1.0	30.0	87.0	13.0 13.0	55.5 55.5	132.0	20.0 27.0	80.5	177.0
-14.5 -14.5		42.0	5	31.0	88.0	13.5	56.0	132.0	27.0	81.0	178.0
-14.0	6.0 6.5	43.0 43.5	— .5 — .5	01.0	88.5	13.5	56.5	133.0	27.5	01.0	179.0
	6.5 7.0	43.5 44.5	5	31.5 32.0	89.5	13.5	50.5 57.0	134.5	28.0	81.5 82.0	179.5
-14.0 -13.5		44.5	.5	32.0 32.5	90.5	14.0	57.5	134.5	28.0	82.5	180.5
-13.9	7.5									04.0	101.5
-13.5	8.0	46.5	.5	33.0	91.5	14.5	58.0	136.5	28.5	83.0 83.5	181.5 182.0
$-13.0 \\ -13.0 \\ -12.5 \\ -12.0 \\ -12.$	8.5	47.5	1.0	33.5	92.5	14.5	58.5	137.0	28.5 29.0	80.0	182.0
	9.0	48.0	1.0	34.0	93.0	15.0	59.0	138.0	29.0	84.0	100.0
12.0	9.5	49.0	1.5	34.5	94.0	15.5	59.5	139.0	29.0	84.5 85.0	184.0 185.0
	10.0	50.0	1.5	35.0	95.0	15.5	60.0	140.0	29.5	80.0	180.0
-12.0 -11.5	10.5	51.0	2.0	35.5	96.0	16.0	60.5	141.0	29.5	85.5	186.0
	11.0	52.0	2.0	36.0	97.0	16.0	61.0	142.0	30.0	86.0	187.0
-11.5	11.5	52.5	2.5	36.5	98.0	16.5	61.5	143.0	30.0	86.5	188.0
-11.0	12.0	53.5	3.0	37.0	98.5	16.5	62.0	143.5	30.5	87.0	188.5
11.0	12.5	54.5	3.0	37.5	99.5	17.0	62.5	144.5	31.0	87.5	189.5
-10.5	13.0	55.5	3.5	38.0	100.5	17.0	63.0	145.5	31.0	88.0	190.5 191.0
-10.5	13.5	56.0	3.5	38.5	101.5	17.5	63.5	146.5	31.5	88.5	191.0
	14.0	57.0	4.0	39.0	102.0	18.0	64.0	147.0	31.5	89.0	192.0
9.5	14.5	58.0	4.0	39.5	103.0	18.0	64.5	148.0	32.0	89.5	193.0
9.5	15.0	59.0	4.5	40.0	104.0	18.5	65.0	149.0	32.0	90.0	194.0
— <u>9.0</u>	15.5	60.0	4.5	40.5	105.0	18.5	65.5	150.0	32.5	90.5	195.0
- 9.0	16.0	61.0	5.0	41.0	106.0	19.0	6 6.0	151.0	33.0	91.0	196.0
— 8.5	16.5	62.0	5.5	41.5	107.0	19.0	66.5	152.0	33.0	91.5	197.0
- 8.5	17.0	62.5	5.5	42.0	107.5	19.5	67.0	152.5	33.5	92.0	197.5
- 8.0	17.5	63.5	6.0	42.5	108.5	19.5	67.5	153.5	33.5	92.5	198.5
- 8.0	18.0	64.5	6.0	43.0	109.5	20.0	68.0	154.0	34.0	93.0	199.5
- 7.5 - 7.0	18.5	65.5	6.5	43.5	110.5	20.5	68.5	155.0	34.0	93.5	200.5
- 7.0	19.0	66.0	6.5	44.0	111.0	20.5	69.0	156.0	34.5	94.0	201.0
7.0 6.5	19.5	67.0	7.0	44.5	112.0	21.0	69.5	157.0	34.5	94.5	202.0
- 6.5	20.0	68.0	7.0	45.0	113.0	21.0	70.0	158.0	35.0	95.0	203.0
— 6.5	20.5	69.0	7.5	45.5	114.0	21.5	70.5	159.0	35.0	95.5	204.0
- 6.0	21.0	70.0	8.0	46.0	115.0	21.5	71.0	160.0	35.5	96.0	205.0
- 6.0	21.5	71.0	8.0	46.5	115.5	22.0	71.5	161.0	36.0	96.5	206.0
— 5.5	22.0	71.5	8.5	47.0	116.5	22.0	72.0	162.0	36.0	97.0	206.5
5.5	22.5	72.5	8.5	47.5	117.5	22.5	72.5	162.5	36.5	97.5	207.5
5.0	23.0	73.5	9.0	48.0	118.5	23.0	73.0	163.5	36.5	98.0	208.5
- 4.5	23.5	74.5	9.0	48.5	119.5	23.0	73.5	164.0	37.0	98.5	209.5
4.5	24 .0	75.0	9.5	49.0	120.0	23.5	74.0	165.0	37.0	99.0	210.0
- 4.0	24.5	76.0	9.5	49.5	121.0	23.5	74.5	166.0	37.5	99.5	211.0
									38.0	100.0	212.0

the field. Temperature data are to be observed, reported, and published to the nearest 0.5 degree Celsius.

Suspended-Sediment Concentration

Sediment concentration may be determined as the ratio of the weight of the sediment to the (1) weight of the water-sediment sample, (2) weight of the water in the water-sediment sample, or (3) weight of the pure water equal in volume to the volume of the sample. Discharge-weighted concentration is usually determined by the first method and is the concentration determined by the laboratory and referred to in this manual. Because of convenience in the laboratory, it is determined in parts per million and is defined as the dry weight of sediment divided by the weight of the watersediment mixture multiplied by 1 million. As the concentration is published in miligrams per liter, however, the values determined in the laboratory must be converted to milligrams per liter prior to computation of sediment discharge or publication.

The discharge-weighted mean concentration in the vertical generally is obtained from depth-integrated samples obtained with standard velocity-weighting samplers. The mean concentration in the vertical also may be obtained from point samples, which represent equal units of depth by (1) weighting each sample by the velocity at each sampling depth or (2) recording the sampling time for each sample and using the weight of the sample collected per second in lieu of point velocity to weight each sample. A discharge-weighted mean in the vertical also may be obtained from a composite of point samples if all samples in the vertical are taken for an equal period of sampling time (U.S. Inter-Agency Committee on Water Resources, 1963, p. 46-50).

The discharge-weighted mean concentration in the cross section may be computed from the mean concentrations of the several sampled verticals. If the sampled verticals represent centroids of equal discharge (EDI method) (Guy, 1970), the mean concentration is the average of the several verticals or is the mean of the composited samples, provided all samples are of the same volume. Thus, samples obtained by the EDI method that are to be composited for particle-size analysis must be the same volume. Samples collected at centroids to define lateral distribution of sediment in the cross section should be analyzed individually and, therefore, do not require an equal volume of water in each sample. If the sampled verticals are uniformly spaced and the same transit rate is used for all verticals (ETR method) (U.S. Inter-Agency Committee on Resources, 1963, p. 41), the mean concentration is the ratio of the total weight of sediment to the total weight of the water-sediment mixture in all samples. Hence, samples collected by the ETR method must be composited either in the laboratory or arithmetically, because the concentration of any individual sample is relatively meaningless.

Concentration data obtained to compute sediment discharge should define the vertical and lateral distribution of concentration in the cross section and the variation of the mean concentration with time. Each sample obtained at daily and periodic stations is analyzed for concentration, and the results are listed in the concentration notes (fig. 5). Concentration notes also include the date and time and identify the sampling and laboratory procedures. Samples may be composited for analysis or analyzed individually.

Compositing, as used here, is the practice of combining the water-sediment mixture of all samples into one container to determine the concentration or particle-size distribution. The mean concentration of a composite sample is the ratio of the total weight of the sediment to the total weight of watersediment mixture. Samples usually composited are those collected only to define the average concentration in the cross section, those collected for analysis of particle-size distribution, and those collected by the ETR method. Samples analyzed individually are those collected to define the vertical or lateral distribution of concentration in the stream

UMIED STATES DEPARTMENT OF TRE MILITOR Garingesi Survey-Maar Messuros Diminin SEDIMENT CONCENTRATION NOTES, DEPTH INTEGRATED SAMPLES			(Short form)
	UNITED STATES DEPARTMENT OF THE INTERNOR	Geeingical Survey-Mater Reservces Omsion	SEDIMENT CONCENTRATION NOTES, DEPTH INTEGRATED SAMPLES (Short form)

<u>a 1 5</u> L RL.PC 2 4 6 -

1235 1235 19.60 10					-							
Me T15 on T20 on			2-12-0-1		12-12-63		12-12-63					
No Number 7.60			501	7:20	1220 at	1225	1235 251	2521	1242	1246	1250	1254
Model Dex Dex Age D me are cons 44.6°/ 49.6°/ 49.6°/ 49.6°/ 49.6°/ 49.6°/ 40.6°/ <t< th=""><th></th><th>Ĭ</th><th>7.60</th><th></th><th>7.66</th><th>7.60</th><th>7.60</th><th></th><th></th><th></th><th></th><th>7.60</th></t<>		Ĭ	7.60		7.66	7.60	7.60					7.60
A A A A A A A A A A A A A A B B A A A A A A B B A A A A A A B B A A A A A A B B A A A A A B B A A A A A B B A A A A A B B A A A A A B B B A A A A B B B B B B A B B B B B B A B B B B B B A B B B B B B A B B B B B B A B B B B B B A B B B B B <th></th> <th>ang Sta</th> <th>Dex .</th> <th></th> <th>Di Ber</th> <th>Bex</th> <th>Di 100 ' 201</th> <th>200 REW</th> <th>300 REW</th> <th>400 EEU</th> <th>500' REV</th> <th>600' EE 1</th>		ang Sta	Dex .		Di Ber	Bex	Di 100 ' 201	200 REW	300 REW	400 EEU	500' REV	600' EE 1
Model J:	Į		460/		/• 57	1.8+	1.84	/	1.30			1.87
Game Obs B.E. A.L. A.L. A.L. Tans Tans 1 3 4 3 4 1 </th <th>Part of</th> <th>Ŧ</th> <th>d= 8.1</th> <th></th> <th>111 0</th> <th>4</th> <th></th> <th>1.8.4 %</th> <th>1:54 3/6</th> <th>1.6.3 40</th> <th>d. 3.9 5/4</th> <th></th>	Part of	Ŧ	d= 8.1		111 0	4		1.8.4 %	1:54 3/6	1.6.3 40	d. 3.9 5/4	
Two Two 315 364 346 219 1 Met 1 34 346 219 2	1		بد 20		7.4	A. L.	A . L .	AL	۲	۹۲	۸L	٦۲
Met L 315 364 346 219 1 Interner m. 340 341 341 341 3 Interner m. 340 340 341 341 3 Interner m. 340 341 341 341 3 Interner m. 340 23.0341 31.7353 3 3 3 Interner m. 23.0341 21.7353 31.7353 3 <td></td> <th></th> <td></td>												
Million re. 340 341 341 3 Gous 22:03n1 21:7351 3 <	8 8	Ĩ	1318	364 1	396	219 1	116	301	385	312	402	225
Cross 22.03-1 21.735/ Tons 23.034/ 21.735/ Tons 23.034/ 21.7393 Nut 0.0106 0.0056 Nut 0.0056 0.0056 Nut 0.0106 0.0056 Nut 0.0106 0.0056 Nut 0.0106 0.0056 Nut 0.0106 0.0056 Nut 0.0056 0.0056	Contas	mer ne.	340		146		342	3 1 3	344	_	346	347
Two 21.024/ 21.7293 Nu 0.0106 0.0058 Nu 0.0106 0.0058 Nu 16 0	9	aross.	7 PEO .22		21.7351		21. 1177	21.1473	21. 6501	22.1333	21.6566	22.4100
Nut 0.0106 0.0058 Nut 0.0058 Nut (um) (6)	ت الا ل	Tere			21.7293		21. 2154	21.1450	21.6471	22 1302	21.6519	22.406
		Ţ			0.0058		0.0023	0.0023	0:0030	0.0031	0.0047	0.003
1 8 8 10	038 Mi	0.5 cerr										
	-	¥										
	ÿ	Î	i i		0	(7	~	00	01	12	4

$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	8		2-21-21				12-14-63		12/5-63		27-7/-21	
7.60 7.60 7.60 7.60 7.60 7.65 7.65 $\psi F'$ ψ	Ē	1	0061	1	7 /5 400	7:20	+ +	7:20	1:00 257	8.05	7:1540	7:20
D: Ber h $\frac{1}{4}$	J	The head	7.60	7,60	7.60		7.60		7.55		7. 55	
4F' $4g'$ $4f'$ $3f'$ <t< td=""><td>ğ</td><td>nehng Sta.</td><td>_</td><td></td><td></td><td></td><td></td><td></td><td></td><td></td><td>•</td><td>-</td></t<>	ğ	nehng Sta.	_								•	-
$d \cdot F_1$ 4_1 4_2 3_1 4_1 b_1 b_2 b_2 b_1 b_2		10 111 11	*		.87		/. Jt	/	1.14	/	/ 87	
AL AL bf bf bf bf bf bf bf $i = 322$ $33/i$ $33/i$ $31/i$ $35/i$ $35/i$ $357i$ $330i$ $330i$ $i = 322$ $33/i$ $347i$ $356i$ $35/i$ $357i$ $353i$ $212 2037$ $31i f 53i$ $21i f 52i$ $21i 166i$ $22i 7429i$ $22i (777)$ $21i f 53i$ $21i 9439i$ $21i 716i$ $22i 752i$ $22i (777)$ $21i f 23i$ $21i 9439i$ $21i 716i$ $22i 752i$ $22i (777)$ $21i f 23i f 21i 9439i$ $21i 716i$ $22i 752i$ $22i (777)$ $21i 716i$ $21i 716i$ $22i 752i$ $22i (777)$ $2004i$ $0.004i$ $0.00ff$ $0.00f7$ $9 - 10$ 10 13 33 10 12	E.	nerts	1.1.1		\$	æ/æ						
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$	3	Gerrobs.	77	YT	۵E	βĒ	ΒĒ		δ£		ЛE	
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	30 10	Tare										
348 349 350 351 352 353 22 20 21 1571 21 1571 21 22 1777 21 1531 21 152 22 22 1777 21 1531 21 152 22 22 1777 21 1531 21 152 22 22 1777 21 1531 21 156 21 22 21 21 21 21 155 2 0 0 0 0 0 4 10 1 1 1 1	45 18	Ĩ	1322	331 ,	336	315	335	318	418	359 ,	330	323
22 20:77 21:131 21:157 21:9524 21:150 22:752 22:1750 21:176 21:156 22:752 22:1750 0:005 0:004 0:0015 0:007 0:0050 0:005 0:0015 0:007 0:007 1 1 1 1 1	ğ	Namer no.	348		349	350	351		352		353	
22 1977 21 21 1916 21 1552 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 3 3 3 3 3 1		Gross	7202 22		21.8321	21.1579	21.9524		21.7150		22.7629	
p< 0060 0.0035 0.0045 0.0047 0.0077 q 10 13 13 13 1 12	1	Tare	22 1977		21.9286	21. 2538	21.9439		21.7061		22.7552	
	300KG 200 200 200 200 200 200 200 200 200 20	L.,			0.0035	0.004	0.0015		0.0067		0.0077	
	35 M											
	•	Ĩ										
	3	nc (p @m)	0		01	61	13		11		/2/	
	×	Se (REV. 10-62)		(6	E	2	5	2	Š	2	5	1711

Figure 5.---Short form for sediment-concentration notes.

10

cross section. Mean concentration in the cross section, or vertical, is computed by weighting the concentration of each sample by the increment of discharge it represents.

Examples of concentration notes are shown in figure 5. Samples individually analyzed are the six-bottle sample of December 12 and the two-bottle sample of December 13. All other samples collected December 12-16 were composited.

The samples collected from 1220 to 1305 hours December 12 were obtained to determine the relation between the mean concentration in the cross section and the mean concentration at the fixed sampling stations. The average concentrations for the fixed sampling station at 1220 to 1225 hours and 1300 to 1355 hours and for six verticals in the cross section at 1235 to 1250 hours are circled.

Adequacy of data

A continuous evaluation of concentration data must be maintained to insure that sufficient samples are obtained and that the samples are of acceptable quality. The stepby-step preparation of records offers a continuing base for cross consultation among personnel responsible for records, laboratory, and fieldwork to evaluate the overall efficiency of the sampling program and to determine if the quantity and quality of the basic data meet desired standards.

Errors in concentration values usually occur because of simple mistakes in sampling procedure or because too few samples were obtained to cover the natural random variation in concentration and size gradation of transported sediment. A description of the sampling procedure is given by Colby (1963, p. 40), by the U.S. Inter-Agency Committee on Water Resources (1963), and by Guy and Norman (1970). Factors that should be evaluated regularly are (1) the number of samples collected in each vertical, (2) the number of verticals sampled in each cross section, (3) the number of samples with respect to time, and (4) the relation of the concentration in the single sampling vertical to average concentration in the cross section.

The number of samples required in each

vertical and the number of verticals which must be sampled to determine the mean concentration within acceptable limits may vary with location and time. A study of the variation of concentration in sand-bed streams is given by Hubbell (1960), and a statistical method for determining the number of samples required is described by Guy (1968). Additional information is available in Guy and Norman (1970).

Relation between single-vertical and cross-sectional concentrations

If sediment samples are obtained routinely at a single vertical in a cross section, the relation of the concentration of the singlevertical sample to the mean concentration in the cross section must be determined prior to computation of sediment discharge. This relation, in the form of a coefficient, is determined by an analysis of cross-section concentration data.

Ideally, sufficient samples should be obtained routinely in the cross section to define the mean concentration both in time and space, and if cost were of no concern this procedure might be selected for all operations. In practice, however, we obtain a computation of routine daily samples at from one to three verticals plus less frequent but more comprehensive samples at sufficient verticals in the cross section to define mean concentration in the cross section. This mean concentration is used to determine the departure of the concentration observed at the single vertical, or fixed-sampling vertical, from the mean concentration in the cross section.

This information should be used (1) to relocate the fixed-sampling station at a vertical that is more representative of the average stream concentration or (2) to determine a coefficient to convert the concentration of the fixed-vertical sample to the mean value for the stream cross section. The adequacy of the sample at the fixed vertical may be determined by an inspection and analysis of the data for stations with uniform concentrations in the cross section and by statistical analysis at stations where variation in sediment concentrations exceed the desired accuracy (Hubbell, 1960).

Cross-section coefficient

The ratio of the average sediment con-

centration in the cross section to the concentration determined by daily samples at a fixed station (box) is computed on the forms shown in figures 6 and 7. This ratio is referred to as the cross-section coefficient.

		/	17		b / -		Nater y	year	a th	
for S	A 6/ A	mente	Kirer a	the		<u> </u>	ending	Sept. 30,	1964	
(I) Date	(2) Time PST	(3) of Sample	Concer (4)	(5) b b	Cross- section coef. a/b	(7) g. h.	(8) Shift	(g) Water disch. (cfs)	(10) Sediment disch. (t/d)	(11) Temp. (*C)
^{/0} 3 63	0820	X sect Conc	8	8		3.83		10,100	818	13
"12 63		X Sect S/30	10	11	7.40	2.67		£ 240	122	14
12 12 63	1240	x sect Cons	19	9	Say ho 1.11	3.74		11,300	315	9
1 15 64	A740	X Sect	4	3		e.51		8,000	86	9
1 21 64	0/00	063 Size		957	_	AL 67		60,800	157,000	7
1 21 64	//••	0 \$ 5 4178		607	_	7.93		24,200	#3 .9 @	7
2-21-64		x sect sens	7	10	~y %	303		7,230	224	11
3 29 64	///9	K Sect Geoc	5	5	1.0	1.73		5,820	79	12
5164	1240	x-sect Cons-	10	11	say 10 - 9/	3.43		10,400	28/	
6 <u>10 64</u>	1115	x-sect	11	7	Say 1.0 1.22	3.33		10,100	300	
7 15 64	0925	r-soct	7	8	Jay /.e	3.15		11,600	182	13
e 19 st	0120	x-sect	9	7	ay /0 /-29	3.14		1600	2/2	14
* 26 64	0745	x-sect	5	6	ay 1.0	2.76		8.480	114	14

SEDIMENT DISCHARGE MEASUREMENTS

Figure 6.---Tabulation of sediment data in the cross section.

SEDIMENT DISCHARGE MEASUREMENTS

U)	Mean	Type (3)	Concentr	ation	Coef.	(7) Chart	(1)	Water(9)	Sediment	Temp
Date	time	of	8 (4)	<u> </u>	a/b	G.H.	Shift	disch.	disch. (0)	-
1942	LTue	Sample	x-sect.	box	a (70	G.R.		(cfs)	(t/d)	(°C)
		Engr			, <u> </u>	1				
<u>2ct.10</u>	1340	x-sect	61.	56	1.09	11.20	-	1,050	173	21
		Engr.			use /					-
Jov. 18	1210	X-sect	31	30	1.03	12.22	- 	1,600	134	13
		Engr	44?	11.0	HJ6 / 0					_
ec. 18	1445	x-sect	44	46	.96	12.80		1,950	232	13
1963						ĺ				
//00		Engr.			<u> </u>			••••••••••		
en.25	1415	x-sect	28} zq	26	1.12	12.11	-	0/170		•
	1413	Engr			1.12	12.11		01,470	// 5	
		X-sect.	30)							
		Engr		?	<u>├</u> ──-			<u>}</u>		
eb. 2	0940	Dex		1630)	1	14.56		3,180	14,000	14
		Rogr. Size			HSE 1.0					/ /
2	1015	X-sect.	1530	(.97	14.62		3240	13,500	14
_		Ergr conc		1						
2	1020	x-sed	1560							
2	1040	Engr		1540					j	
	, , , , , ,	Box		1340						
4	1200	Engr size	167 172	218	.79	23.13		12,200	5,670	12
		Ener.	101 111	<u> </u>		43.75		12,200	3,670	
	1205	Engr. Come	177)							
		Engr Size								
12	0935	X-Seat	/22	154	.79			D 9,700	3200	
4		Eng- 5:30		.						_
Mar. 12	1010	s-sect to	65	14	. 88			D 1,750	307	13
1		Sugar size	(07		. 89	02.25				
pr. 12	1130	prover 2 set	107	120	. 87	22.35		11,200	3,240	/3
May 7	0810	Engr. X-Seet	76	70	1.09	18.56		6.820	1,400	11
7		Engr			1.01	11.36		6,620		<u> </u>
- 14	1430	x-sect.	89	78	1.14	22.62		11,600	2,790	16
,		Engr.								
June 13	0935		78	74	1.05	19.25		7420	1.560	20
		Rog-								
luly 11	0940	x-Sect.		81	1.04	13.69		2,480	562	22
1	-01-	Engr	11.1	_	2		ĺ	1	المح رور	a ./
lug. 15	0730	X-sect	164	73	·/·76	11.30		1,050	465	24
ept 17	1030	Engr X-sect	124	131	.15	12.68		1810	606	22
	1-00	A-3-21	<i>(T</i>			/ =		1.010		
oct. 10					1.00				1	
		!			- -				+	
T										

Figure 7.—Tabulation of sediment data in the cross section.

The manner in which the coefficient is applied depends on the cause of the lateral variation in the distribution of sediment con-

.

centration. This variation may be caused, among other reasons, by proximity and quantity of tributary inflow, bed form, channel alinement, source and type of sediment, season, and discharge. Based on these conditions, each record should be analyzed in detail to determine the most efficient and accurate manner for application of the coefficient.

In addition to using the data to adjust the current concentration values, coefficient analysis also may be used to reevaluate the sampling methods and the location of the sampling vertical at the station. This may make it possible to adjust the sampling locations so that a coefficient that is nearly equal to 1.0 will exist for all conditions of flow.

Common methods for determining daily coefficients involve the correlation of the cross-section coefficient with season and gage height or discharge. As an example, coefficient values for the San Joaquin River near Vernalis, Calif., for the 1963 water year (fig. 7) are plotted against discharge and season (fig. 8). The correlation of coefficient with discharge is poor; however, the correlation with season indicates a possible trend. This trend was investigated by plotting the values of the coefficient on the annual hydrograph of discharge and concentration (fig. 9). The seasonal effect indicated on the hydrograph indicates a coefficient of about 1.0 during the late summer and autum, less than 1.0 during the first few months of the storm season (February, March, and April), and more than 1.0 during the sustained high discharge during the irrigation season; this effect is verified by the repetition of this trend during successive years.

Sometimes coefficients show a reasonable correlation with stage, as indicated in figure 10. The values of the coefficient were determined for 2-percent increments or less for the corresponding range in stage (gage height) and tabulated in the figure. These

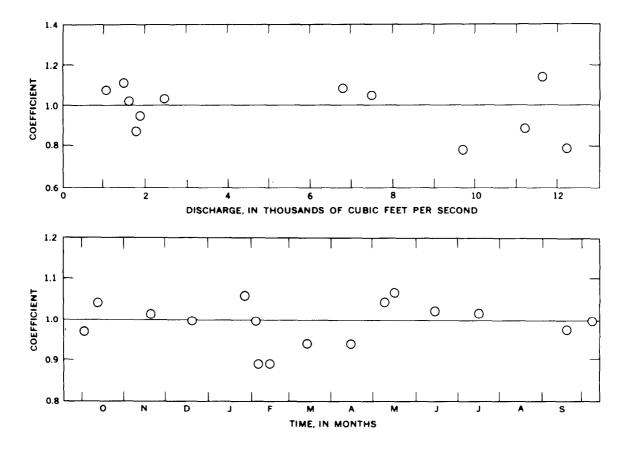


Figure 8.—Relation of cross-section coefficient to discharge and season for San Joaquin River near Vernalis, Calif.

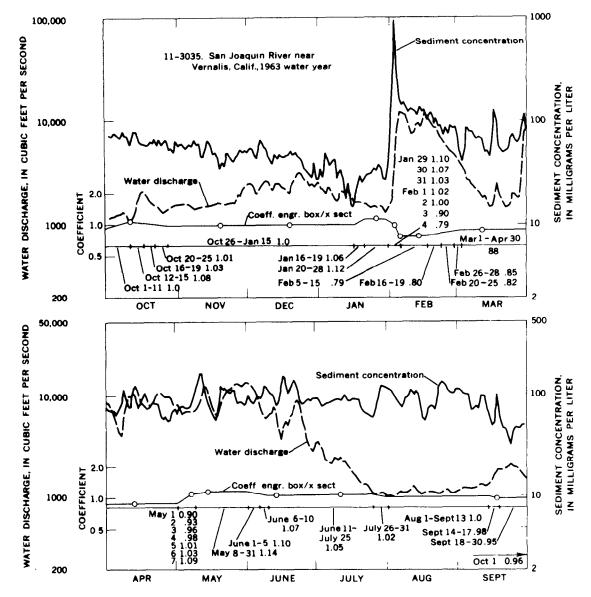


Figure 9.—Water discharge, sediment concentration, and coefficients for correcting observer's single-vertical samples to cross section.

values may be used to correct daily concentration values or concentration values for intervals of a subdivided day.

The average coefficient for the Sacramento River at Red Bluff (fig. 6) was assumed to be 1.0, and no correction was made to the daily concentration values, even though the ratios of individual measurements ranged from 0.83 to 1.33. This example illustrates that the application of a coefficient, as in applying a rating shift to a gage-height value, is a matter of judgment based on the data available. The ratio of 10/9=1.11 indicates a 10-percent (plus) error, and a correction ordinarily would be made. However, the difference of 1 mg/l (milligram per liter) between the cross section and the box sample may be the result of error in laboratory procedures and the result of rounding numbers; therefore, for all practical purposes, such coefficients are ignored. A variation of a few milligrams per liter above 50 mg/l also is considered negligible, and a coefficient is not applied if the indicated corrections

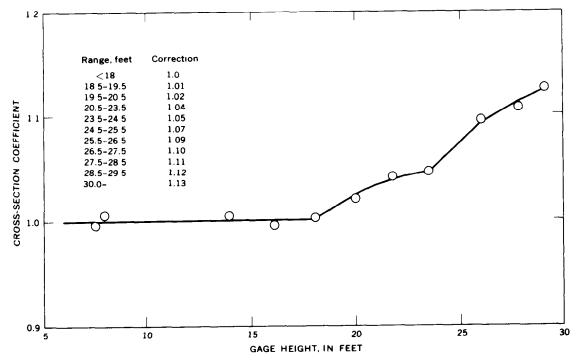


Figure 10.—Relation of cross-section coefficient to gage height.

are random. Coefficients should be applied, however, if all corrections are in the same direction and if the trend persists seasonally and is evident in the record for preceding years.

Variation with time

The number of samples required to define the variation of concentration with time may be difficult to determine from visual inspection of the concentration notes. The effectiveness of the sampling schedule should be evaluated after each storm event. An efficient way to evaluate the adequacy of sampling is to plot the concentration values on the gage-height record as soon as possible after the data are available. Plotting and evaluating the concentration data with respect to time are described in the section "Development of a Temporal Concentration Graph."

Analysis of cross-section concentration data

Concentration values obtained from

cross-section samples are listed on the multiple-purpose form shown in figures 6 and 7, which may also be used to (1) list particlesize analyses and compute the instantaneous discharge required to complete tabulation of size analyses (fig. 2) and (2) list samples obtained at periodic stations or miscellaneous sites.

The tabulation of cross-section samples is used to compute the coefficient needed to adjust the concentration of samples obtained at a single, or fixed, vertical to the average concentration determined by cross-section sampling. The average concentration of the cross section determined from multiple-vertical sampling is recorded in column 4 (fig. 6), and the concentration for the corresponding date and time of the observer's fixed-sampling vertical or three-vertical set is recorded in column 5. The coefficient used to adjust the observer's samples is the ratio of a/b and is recorded in column 6.

The gage height at the time of sampling is obtained from the corrected gage-height record and recorded in column 7. Any gage height recorded on the bottle, particularly by the observer, should be considered as uncorrected data and generally used only to fix the sample in time, to aid in making necessary corrections to the pen record, or to estimate missing gage-height records. The gage height forms the base for computation of the instantaneous water discharge and the sediment discharge.

An evaluation of the quality of the coefficient for a given sampling design may be made by the method developed by Guy (1968).

Samples collected by the EDI method must have nearly equal volumes for each sampling vertical if they are composited; otherwise the bottles must be analyzed individually, and the concentrations for each cross-sectional set of samples averaged. If this procedure is not followed, the quality of the data obviously is affected. Experience has shown that suspended samples from sandbed streams occasionally will be contaminated with varying quantities of bed material. Considerable judgment must be exercised in the field and laboratory to insure that these samples are eliminated from the composite.

The determination of sediment discharge requires the use of water discharge; therefore, the accuracy of the computed value for sediment discharge is dependent on the accuracy of measurements of both water discharge and sediment concentration. In many locations, the water discharge can be determined to a high degree of accuracy from the relation of discharge to stage. If, however, the relation of discharge to stage is not stable, as for most sand-bed streams, or an accurate relation is not available, as for a new station, a measurement of water discharge is necessary at the time sediment concentration is sampled in the cross section. Discharge measurements and discharge ratings at gaging stations are discussed by Buchanan and Somers (1965, 1968) and Carter and Davidian (1965, 1968). An earlier detailed description of stream-gaging procedures is found in U.S. Geological Survey Water-Supply Paper 888 (Corbett and others, 1943).

The evaluation and application of daily

values of the cross-section coefficient are discussed in the section "Cross-Section Coefficient."

Development of a Temporal Concentration Graph

The next step in the computation procedure for sediment discharge is to translate individual values of concentration into a continuous temporal concentration curve. This step may be reasonably simple if values for water discharge or sediment concentration do not vary greatly and (or) if sufficient samples are obtained to define adequately the changes in concentration with time. Accurate results are obtained from a concentration curve defined adequately by samples because a large number of samples successfully integrate the many complex interrelations among variables affecting the availability and movement of sediment in streams (Guv. 1970).

Development of a temporal concentration graph may be difficult if too few samples were obtained. Preparation of the concentration graph will require application of theoretical and practical principals of sedimentation. Inadequate sampling results in a less accurate graph, and much more time is required to prepare the graph. Because of the extra time, in addition to loss in accuracy, it is usually less expensive to collect additional samples than to estimate the concentration graph.

A sampling program for each station should be designed to obtain optimum results when the desired accuracy of record is balanced against the many physical and economic conditions. A few samples properly spaced with time may adequately define the concentration of a flood event at certain stations, providing that the personnel computing sediment discharge have detailed knowledge of seasonal sediment trends for the complete range of flow conditions experienced. Lack of knowledge of these trends, such as at a new station or a station with a large number of variable conditions affecting sediment erosion and transport, requires an intensive sampling program. Successful station operation requires continuous modification of the sampling program to obtain the best accuracy possible with a reasonable expenditure of time and effort.

Concentration data should be interpreted and the graph drawn by personnel with a knowledge of the sampling program, the physical and cultural environments affecting the stream regimen and sediment sources, and the fundamentals of sediment transport. After the graph is drawn, it should be reviewed and modified as required prior to computation of daily mean concentration values and sediment discharges. Changes in the graph are made easily at this point and may eliminate possible future recomputation.

Difficulties may be encountered while drawing the continuous graph because of paucity of samples, unusual storm events, or periods of missing records. Valuable guidance may be available from past records of sediment discharge at the site and at nearby sites. A study of these records before plotting the data and drawing the graph should be a required part of the computation procedure. Some of the factors that should be considered prior to drawing the concentration graph and examples of concentration graphs are included in the following section.

Plotting symbols and scales

Concentration values are plotted on a gageheight chart or a copy of the chart. If an analog record of stream stage is not available because of the use of digital recorders, a plot of gage height or discharges from the digital record must be made for the important periods of changing stage and concentration, such as during rapid snowmelt or storm runoff.

The symbols and scales used for plotting should be chosen carefully and, if possible, be consistent with those used in preceding years. Suggestions relating to the plotting of concentration values and the choosing of scales are summarized as follows:

1. Adjust concentration values from parts

per million to milligrams per liter prior to plotting.

- 2. If necessary, adjust the plotting times for chart-time corrections and travel time between sampling site and gage.
- 3. Plot the average value for each set of samples. Individual values of each bottle should be plotted if poor agreement exists among bottles.
- 4. Use plotting symbols such as the following:
 - Observer samples—mean value.
 - Observer samples—individual samples.
 - \triangle Technician sample at observer's fixed station (box).
 - Technician cross-section sample---mean value.
 - Particle-size sample. Use above symbols and circle if sample analyzed for particle-size gradation.
- 5. Use of a proper plotting scale facilitates computation and checking, increases accuracy of daily mean concentration values picked from the graph, and provides a visual method for comparison and study of various flood events; therefore,
 - (a) Use simple scales such as 1 to 1, 1 to 2, 1 to 5, or multiples of 10 thereof, with zero at the base line.
 - (b) Use as few scales as possible, but do not hesitate to change scale as needed.
 - (c) Plainly mark each change in scale. Use previous year's record as guide to scales. Use the same scale for all events of similar magnitude; such a scale provides a visual means for comparing and evaluating graphs and assists in development of characteristic curves that are extremely helpful in shaping the graph when incomplete sampling data are available.
- 6. Use a maximum height of the graph 5-8 inches above the base line (0 mg/l) on the gage-height chart. As the concen-

tration decreases after a storm, change the scale when the concentration graph approaches to within 1 or 2 inches of the base line. Experience indicates that personnel drawing a graph near the base line tend to be influenced by the limiting 0-mg/l base line, and therefore values determined from a graph approaching the base line usually are high.

7. Choose a scale, if possible, so that concentration values can be plotted to three significant figures. For example, if a stream has concentrations that range from 300,000 to 400,000 mg/l, a scale of 1 inch=50,000 mg/l allows a maximum height for the graph of 6-8 inches, and concentrations above 100,000 mg/l (2 inches) may easily be plotted to three significant figures. Below 100,000 mg/l the scale can be read only to two significant figures and should be changed.

Theoretical considerations

Considerable information is available on theory of sediment transport and the factors affecting the availability of sediment for transport. Colby (1964a, p. A3) states that

Relationships of sediment discharge to characteristics of sediment, drainage basin, and streamflow are complex because of the large number of variables involved, the problems of expressing some variables simply, and the complicated relationships among the variables. At a cross section of a stream, the sediment discharge may be considered to depend: on depth, width, velocity, energy gradient, temperature, and turbulence of the flowing water; on size, density, shape, and cohesiveness of particles in the banks and bed at the cross section and in upstream channels; and on the geology, meterology, topography, soils, subsoils and vegetal cover of the drainage area. Obviously, simple and satisfactory mathematical expressions for such factors as turbulence, size and shape of the sediment particles in the streambed, topography of the drainage basin, and rate, amount, and distribution of precipitation are very difficult, if not impossible, to obtain.

References that will aid in understanding the interrelation of some of the above-listed variables and sediment discharge are cited in pertinent text sections and are listed at the end of this manual. This list is by no means complete, but will serve as a starting point for those interested in furthering their understanding of sediment transport.

Study of past records

A study of the variation and range of suspended-sediment concentration with time at a given point, or sampling station, reveals many similarities among different flood events. A plot of concentration values with time and with flood stage will define graphs that can be used to estimate concentration graphs for missing periods or for inadequately sampled periods. The absolute values and duration of these values may vary considerably from event to event; however, the shape of the temporal graph may be similar among the several events. Thus, the first step in drawing the concentration graph is to study the plotted points for trends, sketch in the parts of the graph well defined by samples, and study those parts defined previously-for the entire historical record if necessary.

A file of historical concentration graphs that are characteristic of the variation and range of suspended-sediment concentration should be assembled to facilitate the use of these graphs during development of the temporal concentration graph and to reduce the number of past records stored in current files. Characteristic graphs may be different for different basins, and many characteristic graphs may exist for each station.

Relation of water discharge to concentration

The relation of water discharge to concentration is an important aspect to consider when developing the temporal concentration graph. The variation of water discharge, as depicted by the continuous graph of stage on an analog chart or a plot of bihourly discharges from a digital record, provides a valuable clue to the time and magnitude of changes in the sediment concentration of the stream. The relation between water discharge and sediment concentration is not fixed. It is affected by many variables, and the variation and range of concentration during one storm period or during one low or medium streamflow period may differ from the concentrations during other periods, even though the streamflow may be identical or similar. Therefore, interpretation of concentration data and the drawing of the temporal concentration graph always requires consideration of the variables that affect the relation between water discharge and concentration.

Availability of sediment is a major variable affecting sediment concentration. Factors affecting availability are discussed in detail by Guy (1970) and in many texts. The availability for a short period may be considered relatively constant, and curves characteristic of the relation of water discharge to concentration for diverse storm periods, tributary inflow, and seasonal effect may be assembled for ready reference. (See previous section and the section on "Examples of the Sediment-Concentration Graph.") Changes in natural availability of sediment may be caused by such events as forest fires or channel changes, landslides, and mass wasting associated with or accelerated by catastrophic floods. These changes should be noted and considered during development of the concentration graph.

Availability of sediment also is influenced by the activities of man. Activities which may cause rapid and large changes in sediment availability include road construction, dam construction, diversions, land-use changes, logging, urbanization, and gravel mining.

Basin size may affect the correlation of concentration and water discharge and the shape of the concentration graph. In general, the smaller basin has a more predictable relation between water discharge and concentration than the large basin, which is often affected by a larger number of variables. The Colorado River at Grand Canyon, for example, has tributaries affected by many variables. These tributaries include rivers with large flows and very low concentrations as well as streams with small flows and large concentrations. All these water and sediment conditions, plus regulation by upstream dams, imposed on one downstream station cause a large range in concentration for a given water discharge.

Other factors affecting the relation between sediment discharge and streamflow are listed in the section on "Theoretical Considerations."

Estimates for periods of missing data

The shape and magnitude of the temporal concentration graph for individual rises have characteristics based on the principles previously discussed. A knowledge of the typical patterns from past records is helpful when interpreting the concentration data and constructing the concentration graph for periods of inadequate concentration data.

Concentration data are considered inadequate when a significant part of a record cannot be defined within probable limits of 5 or 10 percent. The efficient and reasonably accurate development of a continuous concentration graph or determination of sediment discharge during the period of missing data requires careful study, in which experience and ability to make sound estimates based on concentration data collected during other periods are most helpful. The length of the inadequately defined period may range from 20 minutes to several days. The short period usually occurs on streams having rapid changes of water discharge and concentration and very frequently occurs at the beginning of a rise resulting from intense rainfall. This situation is particularly critical on streams in arid regions and on streams with small drainage areas. Long periods of missing data may occur because the sampling site is inaccessable during floods or because of loss of equipment or samples.

An estimated concentration graph is preferable to direct estimates of sediment discharge. During short periods of missing data, a continuous concentration graph may be estimated accurately and used to compute daily mean concentration and sediment discharge. During long periods of missing data. an accurate estimate of concentration may not be possible, and daily values of sediment discharge must be estimated directly from the historical relation between water and sediment discharge by interstation correlation or by comparison with records obtained at an upstream or downstream station. A complete record of daily values facilitates interpretation or statistical evaluation of the data by computer techniques; therefore, if possible, estimates of both sediment concentration and discharge should be made. During periods that sediment discharge was estimated directly, daily concentration values must be estimated independently of sediment discharge if the period includes rapid or large changes in concentration or water discharge. An independent estimate of daily mean concentration is necessary because published values of concentration are time weighted, and daily time-weighted values of concentration cannot be computed from daily values of water and sediment discharge that represent periods of changing streamflow and concentration. If an acceptable estimate of concentration is impossible, no daily concentration will be published, and a leader (..) will be placed in the concentration column.

The methods or combination of methods used to estimate missing data may vary from station to station and seasonally for the same station. Each period of missing data, therefore, must be studied, and the best estimate made on the basis of existing data and circumstances; regardless of the method chosen the estimate should be verified by a second method. A partial list of methods commonly used to estimate sediment data follows.

Visual comparison with adequately defined concentration graphs

The visual procedure, when supplemented by the two succeeding methods, probably is the most common and accurate method used to construct concentration graphs for periods when data are insufficient. The principles involved are discussed in more detail elsewhere in the manual and especially in the sections on "Study of past records" and "Examples of the sediment-concentration graph." Each station should be sampled in detail during sufficient runoff events to provide a catalog of the shape and magnitude of the sediment curves pertinent to the station.

The shape of the concentration graph with respect to the gage-height graph should be carefully considered as to the time the rapid increase starts, the time of peak concentration, and the slope of the recession curve. Typical concentration graphs of the various types, such as advanced, simultaneous, and lagging, are illustrated in figures 11, 12, and 13.

Hydrographic comparison with records of upstream and downstream stations

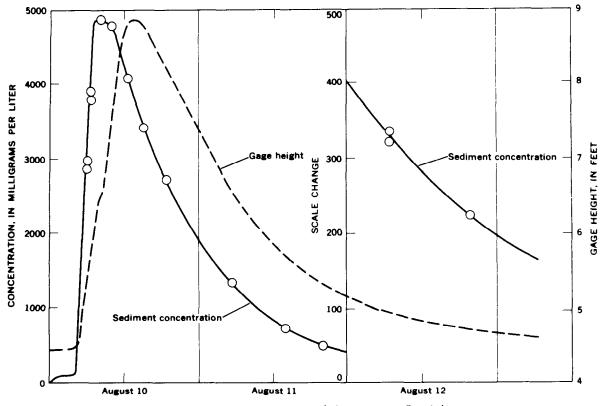
Hydrographic comparison is an excellent tool to check the accuracy of the concentration record and sampling program, as well as to estimate periods of missing records. Each record should routinely be compared with adjacent station records wherever possible, and consideration should be given to significant natural and manmade differences that would account for discrepancies in the computations.

Short periods of missing concentration data can be estimated on the basis of the concentration curve for an adjacent station. Longer periods of missing sediment data can be estimated by comparing values of daily sediment discharge plotted on hydrograph form 9-284 (fig. 25).

Water-sediment relation curves

The relation between water discharge and sediment discharge may be expressed by an average curve. This curve is called a sediment-transport curve and is used frequently to estimate periods of missing data or to extend records.

The types of sediment-transport curves are numerous, and the selection of the correct type for each use is important. According to





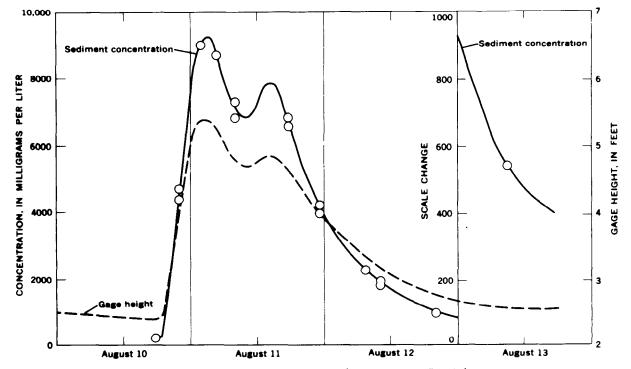


Figure 12.—Simultaneous concentration during excess runoff periods.

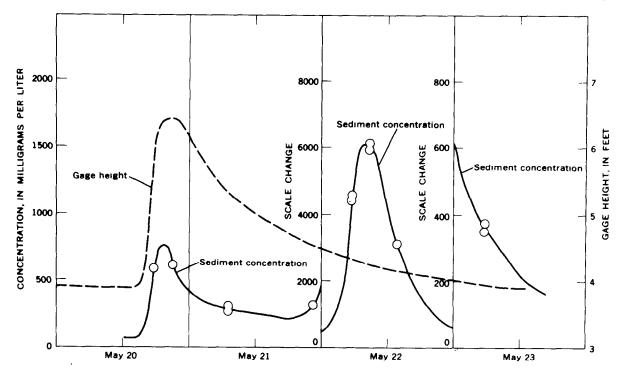


Figure 13.—Lagging concentration during excess runoff periods.

Colby (1956a), sediment-transport curves may be classified according to either the period of the basic data that define a curve or the kind of sediment discharge that a curve represents. Thus sediment-transport curves may be classified as instantaneous. daily, monthly, annual, or flood-period curves. The instantaneous sediment-transport curves are defined by concurrent measurements of sediment discharge and water discharge for periods too short to be materially affected by changes in flow or concentration during the measurements. Daily, monthly, annual, and flood-period sedimenttransport curves usually are defined by and expressed as average sediment and water discharges for periods of days, months, years, or flood periods, respectively. They can be defined by and expressed as total quantities of sediment and water discharges during the respective lengths of time. On the basis of the kind of sediment that they represent, sediment-rating curves may be classified as suspended-sediment rating curves, unmeasured sediment-rating curves, and total-sediment rating curves. These sedimentrating curves may be further divided according to size of particles for which the defining sediment discharges were computed.

The simplest relation between sediment discharge and water discharge is represented by an instantaneous sediment-transport curve. Such a curve is not affected by the extent or pattern of changes in concentration or flow. It is likely to be the most suitable curve from which to determine the effect of different factors on the basic relation between sediment discharge and water discharge and on departures from this relation. On the other hand, an instantaneous sediment-transport curve is not theoretically applicable to the direct computation of daily sediment discharges from daily water discharges except for days on which the rate of water discharge was about constant throughout the day. In practice, however, an instantaneous curve may agree with a daily curve within limits of accuracy of their definition.

Daily or instantaneous sediment-transport curves adjusted for factors that account for some of the scatter from an average curve may be used to compute approximate daily, monthly, and annual sediment discharge. Colby (1956a) describes in detail the selection of the proper type transport curve and the use, preparation, and adjustment of transport curves.

Methods to improve water-sediment relation curves are discussed also by Guy (1964). Two methods using water-sediment relation curves to estimate concentration or discharge are: (1) a plot of the total water discharge versus total sediment discharge or concentration for each storm event or flood period (fig. 14) and (2) a cumulative unit graph relating total water discharge and total sediment discharge for individual storms (fig. 15).

These plots generally are most applicable to small streams with a uniform source of sediment and a low base flow. For many streams the correlations may be greatly improved if the base flow is subtracted from the water discharge. Data must be available for a number of adequately defined hydrographs representing a range of flow and seasons to insure reasonable success with these methods.

The procedures for using these methods are apparent from the illustrations. The limitations of their use will depend on the

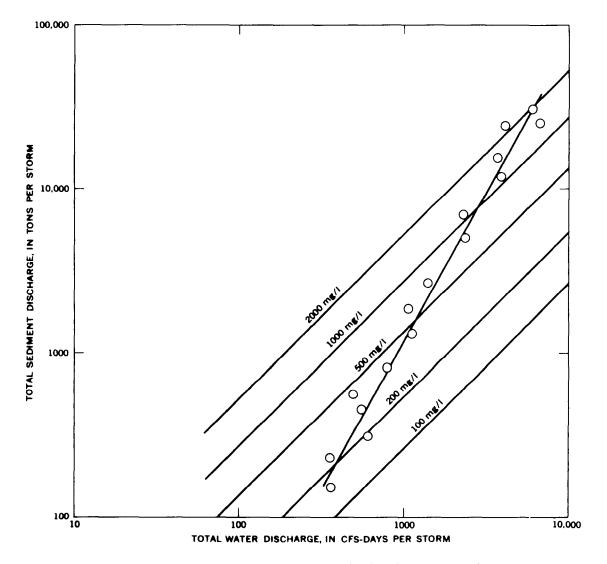


Figure 14.—Sediment-transport curve on a storm basis with indicated mean concentration.

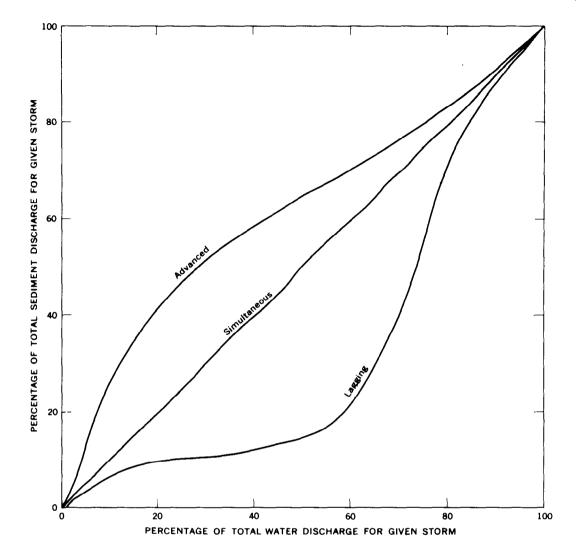


Figure 15.—Cumulative unit relation of total water discharge and total sediment discharge for typical advanced, simultaneous, and lagging types of concentration graphs.

circumstances encountered with the individual record; consequently, as in other interpretive studies, judgment is required. For example, the sediment-transport curve plotted on a storm basis (fig. 11) may be biased (1) if several years of record were used in its preparation and the sediment yield from the basin was changing significantly with time as a result of changing land use or (2) if an unusual number of off-season storms occurred. The change in the sediment yield of a basin was illustrated by a study of 9 years of data of Brandywine Creek, Pa. (Guy, 1957). The cumulative unit graph (fig. 12) is used in conjunction with the total sediment-discharge method to provide estimates for subdividing the storm hydrograph into smaller increments. These methods may or may not be useful in the development of a continuous concentration graph for extended periods of insufficient data, but they are useful for estimating sediment discharge.

Instantaneous values of concentration from advanced and lagging graphs (figs. 11, 13) plot as "loops" on the sediment-transport curves, and this loop effect should be considered if values from the transport curve are used to estimate the shape and magnitude of the continuous-concentration graph for periods of missing records. Daily values also may plot as a loop on the transport curve because of the variation with time of the factors affecting sediment transport and of the subdivision effect.

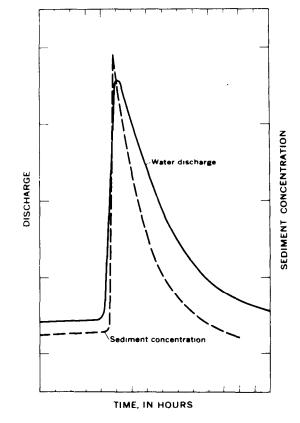
Suspended-sediment discharges computed from any sediment-transport curves, except curves for some streams that transport mostly sands, will be less accurate than sediment discharges computed from frequent samples. The difference in accuracy may or may not be worth the difference in cost of operation. This decision will depend on the particular sampling station and the use to be made of the sediment records.

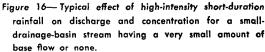
Flow-duration curves have been widely used with instantaneous or daily sedimentrating curves to compute the average sediment discharge for long periods of time when no samples were collected. In principle, and within the limits of averaging and multiplying averages, the method is equivalent to computing average sediment discharge from a daily sediment-transport curve and daily water discharges. The flow-duration curve is used only as a convenient method for abbreviating the distribution of daily water discharges and thereby shortening the computations.

Examples of the sediment-concentration graph

The preceding sections discuss many reasons for the variation of sediment concentration with time and discharge. This section presents examples of (1) the relation between concentration and discharge (or gage height) for basins of various size, climatic conditions, geology, and land use and (2) variations of this relation that may occur in a large basin.

Figure 16 is an example of the typical, sharp discharge peak and concentration graph produced when high-intensity rainfall of short duration occurs over a small basin and the stream channel is dry or has only low flow prior to the storm. The typical con-





centration graph will rise rapidly and peak at or slightly before the discharge peak, after which it decreases rapidly, generally at a faster rate than the recession in water discharge. The shape of the recession curve usually is parabolic. At the discharge peak, the concentration may fluctuate rapidly for a short period before starting to recede. The duration of the concentration peak is seldom greater than that of the water-discharge peak. Note that the concentration did not start to increase prior to the increase in water discharge.

An example of a concentration graph of a stream in a small basin, Corey Creek near Mainesburg, Pa., (12.2 sq mi) when the runoff increased at a slower rate is shown in figure 17. This basin generally has better vegetal cover, less intense precipitation, a more humid climate, and a higher base flow than the basin illustrated in figure 16.

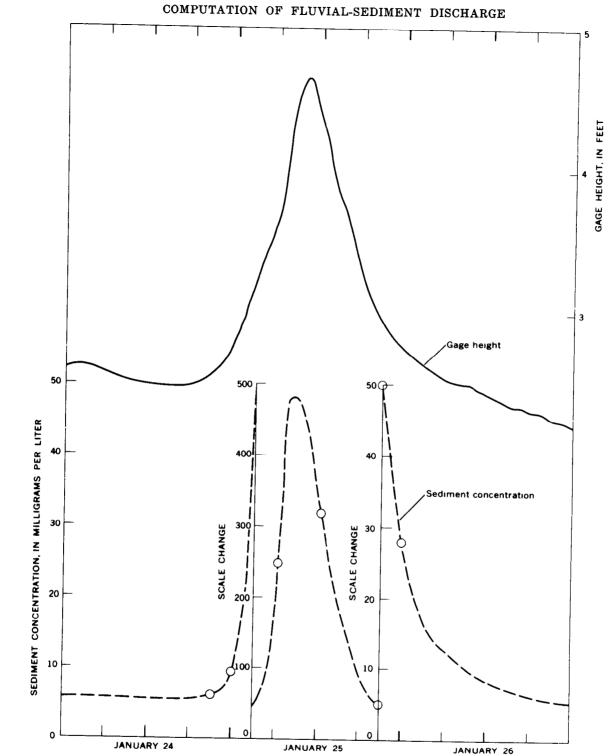


Figure 17.---Gage height and sediment concentration, Corey Creek near Mainesburg, Pa.

Figure 18 shows the effect on sediment concentration in the Rio Grande near Ber-

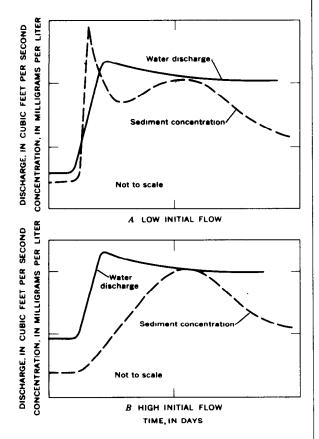


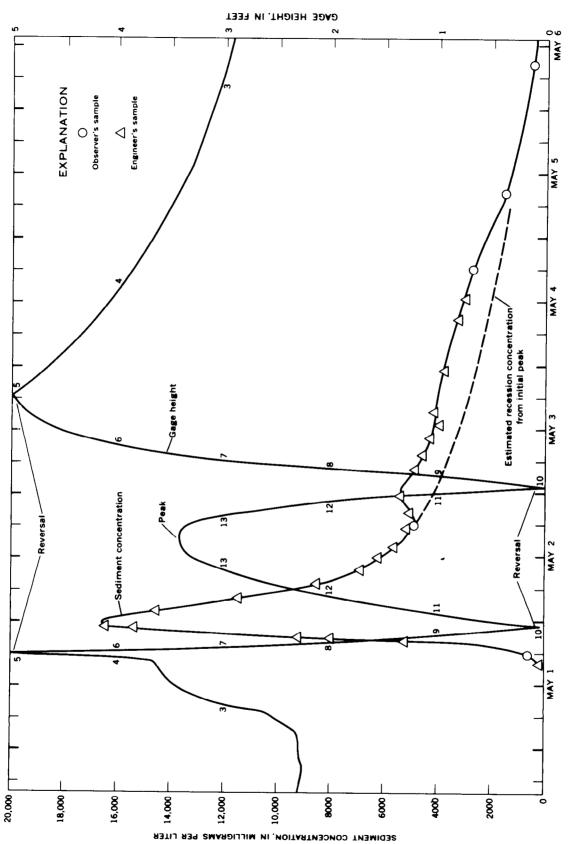
Figure 18.—Effect of two different flow conditions on discharge and concentration for the Rio Grande near Bernalillo, N. Mex.

nalillo, N. Mex., of two separate releases of water from a tributary reservoir over 100 miles upstream. In both instances, the release is at the same rate of discharge; the major difference is in the quantity of water in the stream at the time of release (the initial flow). The shape of the hydrograph is similar in both cases, but there is a marked difference in the sediment-concentration graph owing to the initial flow conditions. Figure 18A illustrates low initial flow conditions. The released water erodes sediment from the bed and the banks of the stream and causes an initial sediment peak. followed by the usual recession, similar to that illustrated in figure 16. After the initial recession another rise in concentration occurs which represents the suspended material contained in, or picked up by, the released water. Figure 18B illustrates the effect of initial channel storage on concentration. Because of high initial flow, the change in stage and velocity is less, and there is little or no additional erosion of sediment from the bed and banks of the stream by the initial increase in flow. The concentration pattern for the released water, however, is the same as that for figure 18A. The interface between the water initially in the river and the released water is defined not only by the changes in suspended sediment but also by a change in temperature and conductivity. In other words, the water represented by the hydrograph peak preceding the sediment-concentration graph is water that was in the channel prior to the release and moved downstream ahead of the release.

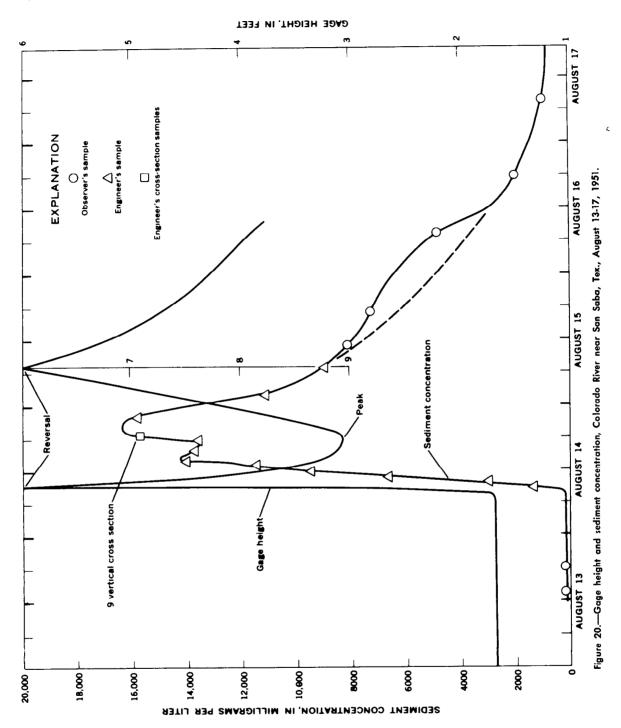
The examples shown in figures 19-22 illustrate for the Colorado River near San Saba, Tex., the range of concentration peaks and the variation of concentration with time which can occur in a river that drains a large basin of diverse geologic, topographic, climatic, and land-use characteristics.

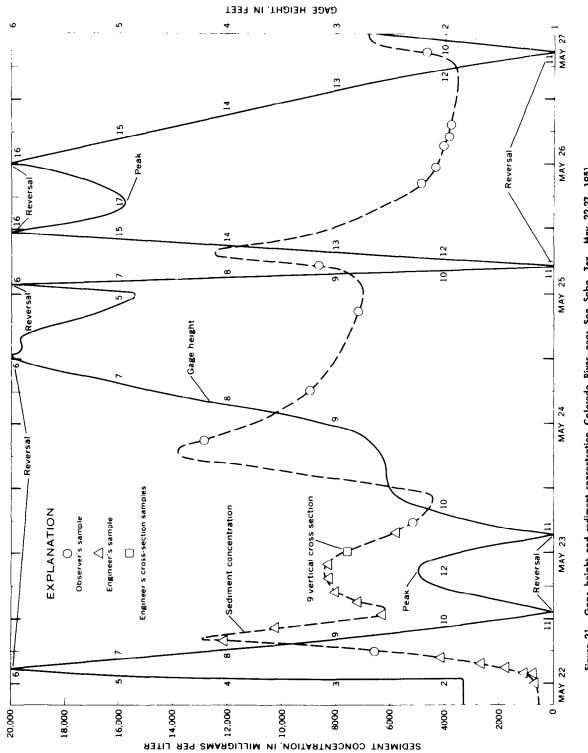
The graphs for the period May 1-6, 1952 (fig. 19), illustrate a typical water-discharge peak and sediment-concentration graph for a large stream when the flow was caused by thunderstorm activity in a small area of the basin. The graphs differ from those shown for a small basin (fig. 16) in that (1) the increase in discharge from 0400 and 1700 hours May 1 is water previously in the channel and (2) the rate of increase of discharge was attenuated by the distance from the source to the station. These two differences cause the significant rise in concentration to be delayed.

Several general conclusions regarding the sediment characteristics of this station can be inferred from figure 19 and illustrate the type of analysis that should be applied to each station record. First, the concentration from 0400 to 1700 hours on May 1 is only slightly larger than the concentration on the preceding day and illustrates a general rule that the concentration graph seldom will show a large increase before the actual storm











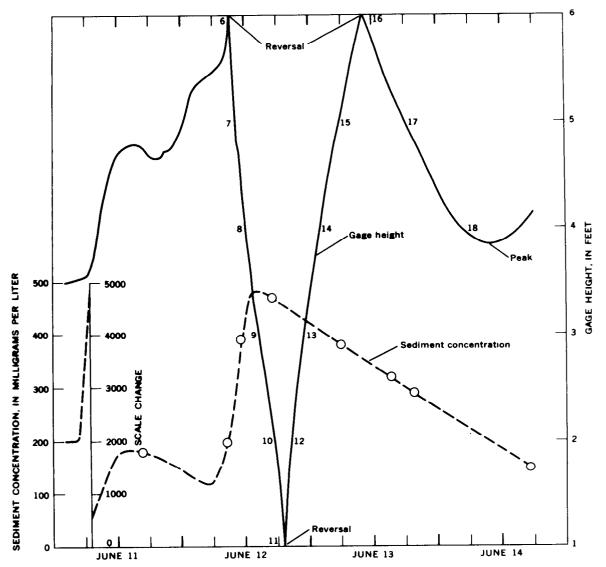


Figure 22.—Gage height and sediment concentration, Colorado River near San Saba, Tex., June 11-14, 1951.

water reaches the station—in this instance, at about 1630 hours. Second, the water peak occurred about 24 hours after the first storm water reached the station, although the concentration peak occurred about 7 hours after the first storm water reached the station.

These graphs illustrate that, for this station, the concentration peak usually precedes the water peak and indicate that, by a comparison of the initial peaks in figures 19-22, the longer the time period between the first arrival of storm water and the storm peak, the longer the time interval between the concentration peak and water peak. Or, conversely, the concentration peak occurred about 7 hours after the initial storm water reached the station, even though the time interval between the initial storm water reaching the station and the water peak increases. Although this time interval (7 hours) should not be considered a firm rule at this station, it could be used in conjunction with the general shape of the concentration curve shown in figure 19 to describe adequately the curves in figures 19-22 even though only two samples had been collected each day.

The May 1-6 rise (fig. 19) has a near

classic hydrograph recession; however, the concentration graph fails to follow the classic pattern. The sediment recession seems normal until 1800 hours May 2, after which the concentration increases and is somewhat above the normal recession curve until about 1200 hours May 5. For purposes of illustration, a normal concentration recession line was estimated for May 2-5 and is represented by a dashed line. The sediment represented by the difference in the estimated graph and the graph based on samples probably was introduced into the main stem by inflow from a small storm on one or more tributaries in the lower part of the basin. The tributary flow contained a higher concentration of suspended sediment than the river, but the water discharge was insufficient to be noticed on the stage record. The effect of various sediment sources superimposed on one hydrograph is more pronounced in the examples to follow.

The period August 13-17, 1951 (fig. 20), has a hydrograph similar to that previously discussed (fig. 19), and runoff apparently came from one source. Correspondingly, the sediment-concentration graph would be expected to have a single rise and characteristic recession. The sediment samples indicate, however, that possibly three major sources of water and suspended material combined to form the single water peak. The initial concentration peak occurred about 4 hours prior to the water peak. Then a tributary flow of higher concentration combined with the initial flow and caused a secondary, and higher, concentration peak. Evidence of a third source of material is indicated by the change in recession rate of concentration about 0300-0800 hours August 16. Finally, on August 16 the sediment concentration dropped abruptly to a level that may have occurred August 15 had the flood peak contained water and sediment from only one source.

The graphs for May 22-27, 1951 (fig. 21), indicate the effect of several peaks produced from several rainstorms or from drainage of several subbasins, or from both. The first increase in discharge was rapid, and the initial concentration peak was conventional, although the peak concentration was not as high as that previously experienced (fig. 20). The difference between this graph and those in the previous examples may be the result of different antecedent conditions in the basin or sediment from a different subbasin. The second concentration peak superimposed on the original sediment recession could not be predicted from the gage-height trace. The third concentration peak may be anticipated because of the abrupt decrease in rate of recession about 2200 hours May 23. The fourth concentration peak, that of May 25, apparently follows the characteristic pattern. The fifth peak (May 27) could not be anticipated from study of the hydrograph and may have been caused by small downstream tributary flow or more likely by bank sloughing which followed the extensive period of high flow.

The period June 11-14 (fig. 22) has a higher water discharge than the preceding examples and a longer delay time between arrival of the first floodwater and the peak discharge, as usually characterized by long periods of general low-intensity rainfall. The sediment concentrations are lower than in the preceding examples. The low concentration may be attributed to antecedent conditions caused by the May storms or, more likely, to the less intense rainfall but longer duration of the June storms.

The examples discussed previously demonstrate some of the variations in concentration graphs that may be expected in a large basin when the runoff events are produced in upstream tributaries of diverse characteristics by isolated rainfall of short duration and high intensity. Figure 23 illustrates a storm event on a large stream, Susquehanna River at Harrisburg, Pa. (drainage area, 24,100 square miles), that drains a basin consisting of three major physiographic provinces with generally good vegetal cover. The March 3-14 flood was caused by intermittent rainfall that occurred March 2-10 throughout the State. The sediment concentration started to increase with the increase in water discharge, unlike the example in

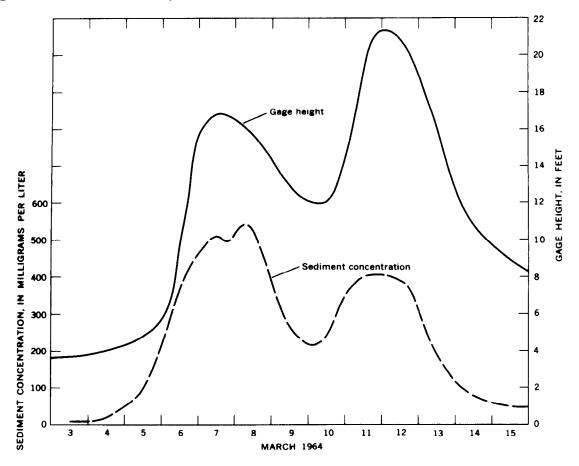


Figure 23.—Gage height and sediment concentration, Susquehanna River at Harrisburg, Pa.

figure 19, because the source area of the water and sediment was local as well as upstream and the concentration continued to increase until the discharge started to decrease. Even so, there was a small secondary concentration peak March 8. The second water-discharge peak on March 11-12, although higher than the first peak, had a lower concentration because less soil was readily available for erosion after the first few days of rain.

The hydrograph of the discharge and suspended-sediment concentrations of the Willamette River at Portland, Ore., during the recordbreaking floods of December 1964 (fig. 24) is a good example of the relation between discharge and concentration for a large flood on a large river. The discharge continued to increase for 4 days until it reached a peak. Sediment concentration, however, reached the maximum value the second day following the beginning of the rise and decreased over 50 percent by the time the water discharge reached a maximum value. Several common characteristic trends may be noted here: (1) The large increase in discharge at the outset caused a minor increase in concentration, (2) the discharge increased slowly for several days to reach a maximum value whereas the concentration increased rapidly and reached a maximum value, in less time, and (3) the water discharge receded slowly, being sustained by additional rainfall and contributions from bank and channel storage, whereas the concentration receded rapidly after reaching the maximum value.

Snowmelt discharge and sediment concentration

The relation between water discharge and

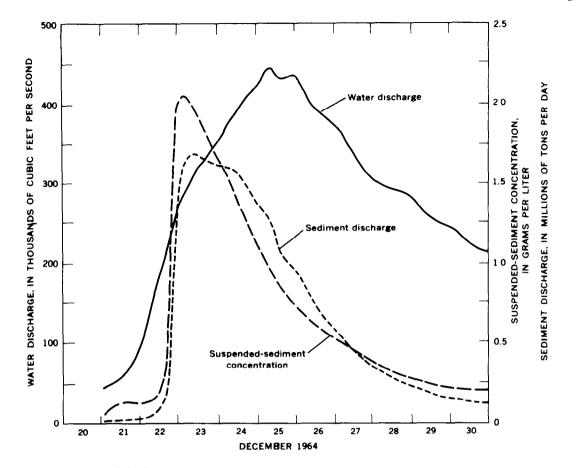


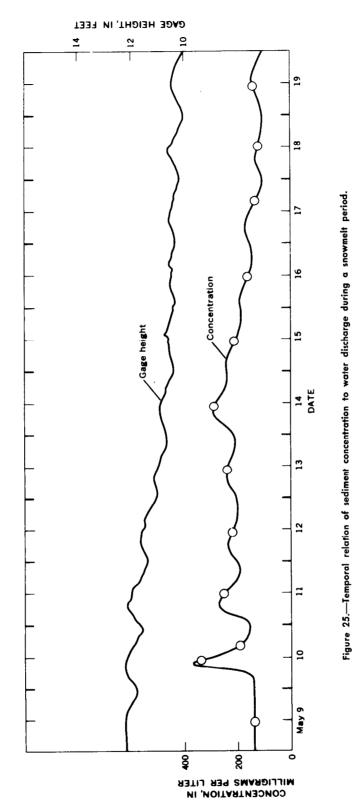
Figure 24.—Suspended-sediment concentration, sediment discharge, and water-discharge, Willamette River at Portland, Oreg., December 21-30, 1964. (After Waananen and others, 1971, p. 116.)

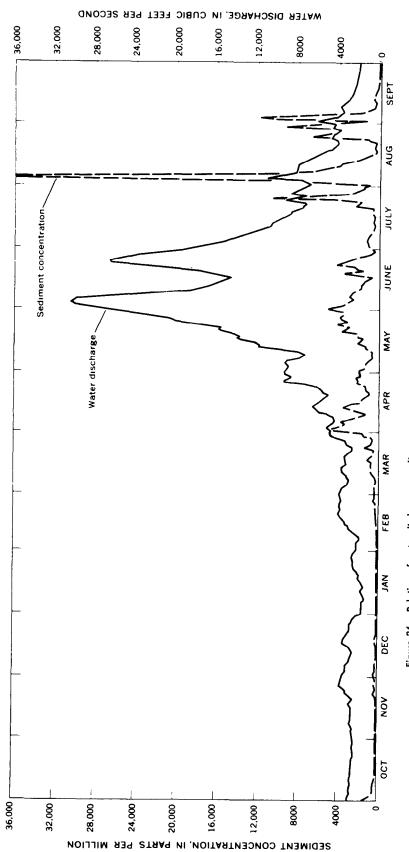
suspended-sediment concentration during snowmelt generally is different from that during periods of runoff from rainfall; also the correlation is poorer during snowmelt. This change in relation must be considered when constructing the concentration graph and interpreting the sediment-transport curve.

Increases in concentration relative to increases in water discharge are small during snowmelt runoff as compared with rainfall runoff, and sharp concentration peaks do not occur at or near the time of the water peak such as those illustrated in figures 16 and 17. Gage-height and concentration graphs of a 10-day snowmelt period are shown in figure 25.

The peak and trough of the diurnal fluctuations of sediment may or may not be in phase with those for water. The time the water peak occurs at the station is related to the distance between the snowmelt source and the gage, and the amount of lag between water and concentration peaks is a function of the distance between the sampling point and the source of the material. This relation may be determined usually from two samples per day.

The relation between daily values of water discharge and sediment concentration in a stream with substantial snowmelt runoff is shown in figure 26. The water discharge and concentration were low during the winter months. The first appreciable snowmelt in late March contained a large quantity of sediment. The snowmelt runoff continued to increase throughout April, although the concentration decreased after the initial rise and remained generally below 4,000 mg/l during the major snowmelt period in May







and June.

During the storm-runoff period July 22-September 1, sediment concentrations were considerably larger than those during the snowmelt period. The rainfall on August 1-2 caused an increase in discharge from 7,000 to 11,000 cfs (cubic feet per second), but the corresponding concentration increased from 1,000 to 39,000 mg/l. The maximum snowmelt runoff which started May 9 caused an increase in discharge from 7,000 to 30,-000 cfs and an increase in concentration only from 500 to 5,000 mg/l.

Application of cross-section coefficient

Coefficients may be applied to concentration values in several ways. Two methods discussed here are (1) correction of the concentration value arithmetically after values have been computed from the concentration graph and (2) correction or adjustment graphically of the concentration graph prior to computing concentration values from the graph. The second method is needed especially in a subdivided period where the coefficient may change with stage and discharge.

In the first method, concentration for each day, or part of day, for the water year is computed from the concentration graph and tabulated on a copy of the recorder chart. When the coefficients are computed, as illustrated in figures 9 and 10, the appropriate correction to each concentration value is written below the daily concentration. The adjusted concentration value then as computed, entered below the other values, and underscored so no error will be made when the value is transferred to the sediment-discharge worksheet. If the subdivided-day method of computation is used and the coefficient values vary during the day, then the concentration value for each interval must be corrected. This is accomplished by tabulating coefficient values taken from a table similar to that given in figure 10 in an additional column in table 3 and by correcting each concentration value in column 5 prior to computing the sediment discharge.

In the second method, the correction is made graphically by adjusting the concentration graph. This method is useful particularly if the transported material is coarse and the ratio (a/b) of mean concentration in the cross section to mean concentration in the single vertical (col. 6, fig. 6) is correlative with gage height or discharge. The concentration graph in figure 27 was adjusted graphically by using the coefficient values determined in the plot in figure 10. The range in gage height affected by each coefficient value was marked on the trace of the gage-height graph. The concentration value determined from the sediment-concentration graph for each range was multiplied by the appropriate coefficient, and an \times placed at the location of the corrected value. An adjusted sediment-concentration graph, following the trend of the original (unadjusted) concentration graph, was then drawn through the adjusted values. Computation of suspended-concentration values, either for daily values or for intervals of a day, may be made directly from the adjusted graph.

Because the reasons for the lateral variation of sediment concentration may differ from station to station and even at the same station, the application of corrections to concentration cannot be reduced to a few simple methods. Hence, as in the construction of the concentration graph, each application of coefficient requires imagination and ingenuity as well as a knowledge of the station and the fundamentals of sediment transport.

All coefficients and adjustments should be applied to the concentration graph, or values, prior to computation of sediment discharge. Any corrections not yet made, and specifically those which vary during a subdivided period, should be made while computing the subdivided period. Additional columns may be added to table 3 to show the coefficient and corrected concentration for each interval.

Adjustments are necessary also if the intervals of a subdivided period do not equal 24 because the recorder is operating too fast or too slow. If the divisor, d (footnote 1,

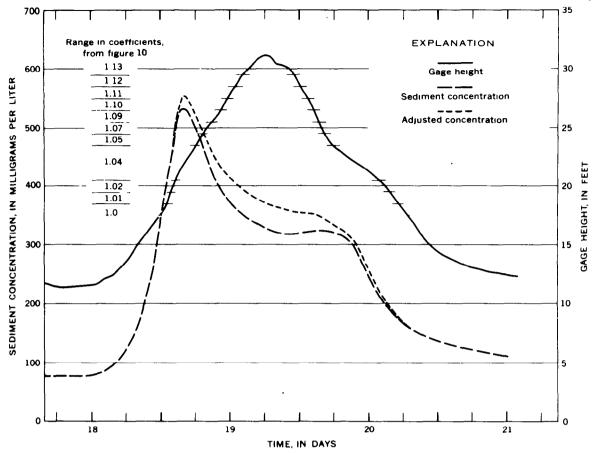


Figure 27.—Graphical adjustment of concentration.

table 3), based on 24 intervals per day, is used to compute the discharge per interval when the total interval, T, is different from 24, each interval must be corrected by 24/T.

Computation of Daily Mean Concentration

The daily mean sediment concentration is the time-weighted mean value and is computed from the concentration graph described in previous sections. Each computed daily value is listed near the bottom of the chart at the midpoint (usually noon) of the day it represents.

Mean concentration values may be determined arithmetically or graphically. The arithmetic mean is obtained either by (1) selecting values from the concentration graph that represent uniform intervals of time and taking the arithmetic average or by (2) selecting values from the graph that represent unequal intervals of time and weighting these values by the time intervals represented by the concentrations. These methods or variations of these methods are also used to determine the average concentration for each interval of a subdivided period.

The mean concentration is determined graphically by use of an integrator—a simple device consisting of a thin sheet of clear plastic about 6 inches long and 3 inches high and with a fine visible line scribed through the center of the sheet parallel with the long axis. A clear plastic 45° or $30^{\circ}-60^{\circ}$ triangle also may be used. A small hole is drilled through the plastic near the center of the line to accommodate a pencil point. The average is determined by placing the line on the plastic over the concentration

39

graph with the hole at the midpoint of the time interval. The line is adjusted vertically or rotated until half the area is above the graph and below the line and half the area is below the graph and above the line. The average concentration is the value where the line intersects the midpoint of the time interval and may be marked on the chart by placing a point of a pencil through the hole in the plastic. This graphical method is used to compute the daily mean concentration for days not subdivided as well as to determine mean concentrations for smaller intervals of time when the day is subdivided.

Because most personnel familiar with computations of water-discharge and sediment records are thoroughly familiar with the graphical method, only a brief discussion is included. A demonstration of the method by experienced personnel plus a little practice and ingenuity is sufficient to obtain acceptable values of mean concentration from the concentration graph.

Figure 28 illustrates ways to use the integrator. The line on the plastic, represented by the line A-B, is placed on the graph so that half the area is above the graph and below the line and half the area is below the graph and above the line. In figure 28A the line A-B is parallel to the base line and intersects the graph in two places. The area below the graph (designated 1) is equal to the sum of the two areas above the graph (designated 2). By placing the line on the integrator to intersect at the beginning of the 24-hour period (point x in fig. 28B), the graph is intersected only once. Thus, only two areas need be balanced, thereby sim-

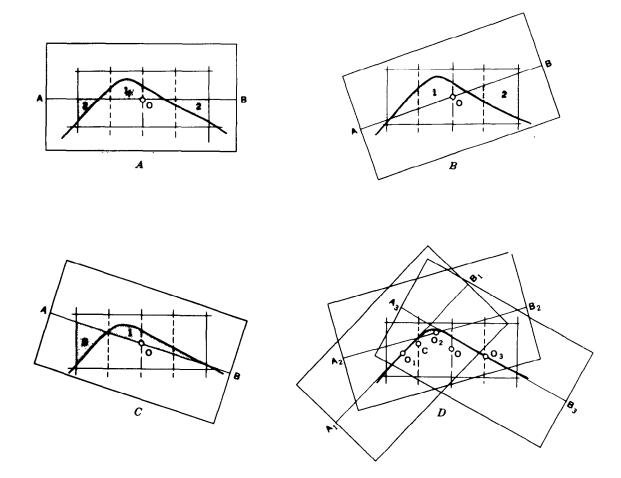


Figure 28.—Determination of mean concentration by graphical method.

plifying the visual balancing of the areas. The balancing of the areas can be further simplified by trial-and-error placing of the integrator to produce areas of uniform geometric appearance. Hence, in figure 28C the integrator line was placed on the end of the 24-hour-time period (point x) to test for uniformity of areas. As the graphs analyzed became more complex, a combination of the approaches in A, B, and C was used, and the time period (usually 24 hours) was divided into several parts, each representing nearly straight lines or small areas to be balanced. This approach is the one usually practiced by experienced personnel and is illustrated in figure 28D.

Figure 28D is divided for convenience into three time intervals-0-0600, 0600-1200, and 1200-2400. Line A_1 - B_1 establishes the mean of the first 6-hour interval at O_1 at the midpoint of the interval. Likewise, A_2 - B_2 establishes the mean of the interval 0600-1200 at O_2 , and A_3 - B_3 establishes the mean of the interval 1200-2400 at O_3 . The mean of the first two intervals, which are equal in time, is found by placing the integrator line on O_1 and O_2 and marking the intersection (point C) of the line $O_1 - O_2$ with the midpoint of the combined time interval, in this case 0600. The mean of the two 12-hour intervals is found at intersections of a line O_3 -C and the midpoint of the combined time interval, O. In each of the examples, the mean is at point O on the midpoint of time intervals.

A value of mean concentration is determined for each day for which a graph has been constructed. For periods of constant, usually low, discharge and low constant concentration, the daily concentration sample will approach the daily mean concentration, and there will be no need to construct a graph to compute daily values. For these periods, sample values may be transferred directly to the sediment-discharge sheet with no intermediate calculation.

Footnotes

Footnotes are used to describe laboratory and computation procedures and to qualify estimated values. The appropriate footnote should be determined as each value is computed and entered after the concentration value on the gage-height graph. The footnote then can be transferred to the sedimentdischarge worksheet simultaneously with copying of the concentration values on the worksheet.

Footnotes generally are not processed or stored by electronic computer; therefore, no footnotes will be listed on the tables of data processed by computer for publication. If footnotes are desired on these tables, they must be added manually by typewriter. Footnotes, however, should be limited to the minimum needed to explain and clarify the data or to describe unusual conditions that affected collection or quality of data.

Sediment footnotes and symbols do not follow the rules listed in the United States Government Printing Office "Style Manual," revised edition, January 1967; they are standardized here and have the same meaning regardless of their position in the table. The proper placement of footnotes in the table is described in the section on "Sediment-Discharge Worksheet."

Footnotes and references and the sequence in which they must appear if used on the sediment-discharge worksheet are listed as follows:

- E Estimated. (When there are 6 or more days in a year with the footnote E, delete E and insert the following headnote under the table title: "When no daily concentrations are reported, loads are estimated.")
- S Computed by subdividing day.
- A Computed from estimated-concentration graph.
- C Composite period.
- D Daily mean discharge.

Footnotes E and A are used to identify noteworthy estimated values, such as maximum and minimum values of concentration and sediment discharge, and to explain any missing values of concentration. Daily sediment-discharge values based on interpolation for short periods of low or uniform flow are not identified; hence, the sampling program must be scheduled to assure that samples are collected at some predetermined frequency, even during periods of low and uniform flow. Records that include long periods of estimated daily sediment discharges and contain the maximum sediment discharge of the year should be explained in the remarks paragraph of the station analysis. Also, conditions that influence sediment discharge such as unusual flow conditions, diversions, or construction activities should be described.

Footnote C is used to identify concentrations that are average values for a period. These average concentrations are obtained by compositing samples during periods of low concentration when fluctuations in concentration are assumed to be small. Daily values of concentration are more useful than average values for statistical treatment or interpretation of sediment trends. Daily values of concentration, therefore, should be published whenever available, and unless the analytical costs are prohibitive, the concentration of each sample should be determined in the laboratory.

Footnote D is used on tabulations of particle-size analyses to identify daily values of water discharge.

Significant Figures

Rounding of sediment-concentration and discharge values is standardized to accommodate and facilitate the computation, processing, and storage of data by electronic computer. Beginning with the 1968 water year, the number of significant figures reported in data releases is as follows:

Sediment discharge
< 0.005 report 0 tons.
0.005-0.01 report 0.01
ton.
.0109 report to near-
est 0.01 ton.
.10 -99 report two
significant
figures.
>100 report three
significant
figures.

The monthly and annual sediment discharges are the unrounded sums of the daily values. No monthly or annual values will be shown for concentration.

The use of two significant figures below 1.0 is not intended to imply accuracy, but does make the published records of sediment compatible with those of water discharge and will facilitate computer processing of the data. A statement explaining the accuracy of published values may be included in the introductory text of the basic-data report.

Computer Programs

The description of the computation procedure thus far has been related entirely to manual preparation of records. Data-processing equipment has outmoded many manual procedures, and many additional changes undoubtedly will occur in the future. Automation, however, will not change the fundamental methods described for manual processing; therefore, each individual involved in the computation of sediment discharge should be familiar with and thoroughly understand the computation procedure and be able to compute records by the manual method.

A simplified computer program is now available for general use. This program performs the routine computations described in the sections on "Computation of Sediment Discharge," "Computation of Subdivided Days," and "Sediment-Discharge Worksheet" and lists the data for offset reproduction. The program does not eliminate the need to interpret concentration and water data and define a concentration graph.

Basic data needed for computer computation of sediment discharge are identical with those needed for manual processing and include the following:

- 1. Water discharge and sediment concentration for each day as described in the section on "Sediment-Discharge Worksheet."
- 2. Water discharge, concentration, and clock time or time interval as described in

the section on "Computation of Subdivided Days."

Format of Sediment Tables

Prior to October 1967 sediment data were typed or listed by automatic data-processing equipment on forms similar to those illustrated in figures 2, 38, and 45. Data published subsequent to October 1967 are processed and the results listed by electronic computer for publication in one or more of the methods in figures 29-35.

Computation of Sediment Discharge

Sediment discharge is determined by multiplying the water discharge, in cubic feet per second, by the concentration of suspended sediment, in milligrams per liter, and a coefficient:

 $Q_s = Q_w \times C_s \times k$.

where

- Q_s is the sediment discharge, in tons per day (tons/day).
- Q_w is the water discharge, in cubic feet per second (cfs),
- C_s is the concentration of suspended sediment, in milligrams per liter (mg/l),

and

k is a coefficient that is based on the unit of measurement of water discharge and that assumes a specific weight of 2.65 for sediment.

Units of measurement

The concentration of suspended sediment, C_s , is reported in milligrams of sediment per liter of water-sediment mixture. However, as a matter of convenience, it is determined in the laboratory in parts per million, which is the dry weight of suspended material per million equal weights of water-sediment mixture, or milligrams per kilogram, and is found by the formula

parts per million= weight of sediment×1,000,000

weight of water-sediment mixture

or

$$1 \text{ ppm} = \frac{0.0010 \text{ g}}{1,000 \text{ g}} = \frac{1 \text{ g}}{1,000,000 \text{ g}}$$

Concentration in milligrams per liter is the weight in milligrams (mg) of sediment per thousand milliliters (ml) of mixture and is the ratio of dry weight of sediment to the volume of mixture, or

milligrams per liter =
$$\frac{0.0010 \text{ g}}{1.000 \text{ ml}} = \frac{1 \text{ mg}}{1.000 \text{ ml}}$$

The numerical values of parts per million and milligrams per liter are equal when the density of the mixture is equal to 1.00, and for all practical purposes, 1 liter weighs 1,000 g (grams). An increase in concentration increases the density of the mixture, and the laboratory value of concentration in parts per million must be corrected by the ratio, C, of the density of the water-sediment mixture to the density of water. Thus,

milligrams per liter = $C \times \text{parts}$ per million.

The values of the conversion factor C, are given in table 2. These values are based on

[[]The factors are based on the assumption that the density of water is 1.000 (plus or minus 0.005), the range of temperature is $0^{\circ}-29^{\circ}$ C, the specific gravity of sediment is 2.65, and the dissolved-solids concentration is less than 10,000 ppm. This table supersedes table 1 in Guy (1969)]

Concentration range (ppm)	С	Concentration range (ppm)	с
0- 15,900	. 1.00	322,000-341,000	1.26
16,000- 46,800	. 1.02	342,000-361,000	1.28
46,900- 76,500	. 1.04	362,000-380,000	1.30
76,600-105,000	1.06	381,000-399,000	1.32
106,000-133,000	1.08	400,000-416,000	1.34
134,000-159,000	. 1.10	417,000-434,000	1.36
160,000-185,000	1.12	435,000-451,000	1.38
186,000-210,000	. 1.14	452,000-467,000	1.40
211,000-233,000	. 1.16	468,000-483,000	1.42
234,000-256,000		484,000-498,000	1.44
257,000-279,000	. 1.20	499,000-514,000	1.46
280,000-300,000		515,000-528,000	1.48
301,000-321,000		529,000-542,000	1.50

Table 2.—Conversion factors, C, for sediment concentration: parts per million to milligrams per liter

	SED IMENT D I SCHARGE (TONS/DAY)	3820 670 - 17к 65 19	8.6 3.9 2.8 313	2430 125 30 13	7.5 3.6 3.6 461 461	258000 5070000 2110000 1020000 346000	259000 135000 60900 44800 34100 23000 23000
SEPTEMBER 1968 DECEMBER	MEAN CONCEN- TRATION (MG/L)	913 318 140 70 35	20 11 65 121	742 110 22 22	17 10 15 249	B100 57900 46300 36500 25500	22300 15500 11000 10700 9430 7950
10	MEAN DI SCHARGE (CFS)	1410 780 470 345 205	160 130 117 241 330	1010 422 275 218 196	163 134 117 134 500	8630 29800 15600 9300 5030	4300 3230 2050 1550 1340 1070 89257
WATER YEAR OCTOBER 1967 NOVEMBER	SED IMENT D I SCHARGE (TONS/DAY)	87 6.7 3.4 53	•15 •15 •15 •15 690 1290	388 289 26 1.2	- 45 - 20 - 16	.09 .11 .11 .11 .11	41 8.8 1330 723 66 5563.43
WATER YEAR NOVEMBER	MEAN CONCEN- TRATION (MG/L)	24 24 25 25 25	4 4 6 5 1 6 7 1	326 300 22 8	*~~~	369 369	22 25 25 25 25 25 25 25 25 25 25 25 25 2
I SCHARGE,	MEAN DISCHARGE (CFS)	ი6 58 28 19	14 18 818 318 459	254 340 148 80 55	4 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	33 42 378 378 378	232 148 486 595 315 315
SUSPENDED-SEDIMENT DISCHARGE, CTOMER	SED IMENT D I SCHARGE (TONS/DAY)	0°00°0	ccccc	0 00.00 005	.07 .07 .05 .03	02 02 02 02	.04 .05 .15 8.6 8.6 .59
SUSPENDEI OCTOHER	MFAN CONCEN- TRATION (MG/L)	+ + + + 	~~	シャシート	みみでくろ		226 226 10 10
25 196A	MEAN DISCHARGE (CFS)	• • • • • • • • • • • • • • • • • • •		. но . но . но . но . но . но . но . но	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	7•0 7•0 7•0 7•0	7.5 9.0 19 80 48 22 22
1138212	ן 104	→ ∿ω4फ	с г а с С 1 о а и с	11 12 15 15	15 17 18 19 20	22 22 24 25	26 27 28 29 30 31 31 10TAL

Figure 29.---Format table 1, suspended-sediment discharge.

44

.

TECHNIQUES OF WATER-RESOURCES INVESTIGATIONS

SUSPENDED-SEDIMENT DISCHARGE FOR SELECTED DAYS, WATER YEAR OCTOBER 1969 TO SEPTEMBER 1970

DATE	MEAN DISCHARGE (CFS)	MEAN CONCEN- TRATION (MG/L)	SEDIMENT DISCHARGE (TONS/DAY)	DATE	MEAN DISCHARGE (CFS)	MEAN CONCEN- TRATION (MG/L)	SEDIMENT DISCHARGE (TONS/DAY)
NOV 28, 1969	4.4	519	13	MAY 7	14	3670	265
JAN 27, 1970	6.0	2180	213	MAY 8	1.9	140	. 72
JAN 28	1.5	50	.20	JUN 22	4.2	1410	122
MAR 6	3.4	143	2.6	JUN 23	.42	80	.09
MAR 7	42	449	28	AUG 3	2.0	1030	61
MAR 8	1.5	70	. 28	AUG 4	5.7	3700	372
MAR 15	6.6	795	26	AUG 5	.98	490	4.0
MAR 16	1.8	50	.24	AUG 24	14	3950	968
APR 17	1.5	326	3.5	AUG 25	27	1270	640

Figure 30.—Format table 2, suspended-sediment discharge for selected days.

SUSPENDED-SEDIMENT DISCHARGE MEASUREMENTS, WATER YEAR OCTOBER 1969 TO SEPTEMBER 1970

DATE	TIME	DISCHARGE (CFS)	CONCEN- TRATION (MG/L)	SEDIMENT DISCHARGE (TONS/DAY)	DATE	TIME	DISCHARGE (CFS)	CONCEN- TRATION (MG/L)	SEDIMENT DISCHARGE (TONS/DAY)
DEC 6, 1969	0450	42	14	1.6	JUN 10	1430	45	12	1.5
JAN 12, 1970	1700	126000	100	34000	JUN 22	2300	219	69	41
JAN 14	1830	1.3	8	.03	JUL 1	1530	.80	12	.03
FEB 22	1200	13	33	1.2	AUG 12	1200	14	8	.30
FEB 28	0100	770	212	441	AUG 15	1845	108	15	4.4

Figure 31.—Format table 3, suspended-sediment discharge measurements.

SUSPENDED-SEDIMENT DISCHARGE DURING PERIODS OF HIGH FLOW, WATER YEAR OCTOBER 1969 TO SEPTEMBER 1970

	DATE	TIME INTER- VAL (HRS)	MEAN DISCHARGE (CFS)	MEAN CONCEN- TRATION (MG/L)	SEDIMENT DISCHARGE (TONS/DAY)	DATE	TIME INTER- VAL (HRS)	MEAN DISCHARGE (CFS)	MEAN CONCEN~ TRATION (MG/L)	SEDIMENT DISCHARGE (TONS/DAY)
JUL	27	2	1340	41000	12400	JUL 28	15	625	11600	12200
JUL	27	2	944	30000	6370	JUL 28	3	400	5700	770
JUL	28	2	730	21000	3450	JUL 28	4	278	3600	450

Figure 32.—Format table 4, suspended-sediment discharge during periods of high flow.

TOTAL SEDIMENT DISCHARGE, WATER YEAR OCTOBER 1969 TO SEPTEMBER 1970

		WATER TEM-		SUSPENDED SEDIMENT	SUSPENDED	TOTAL	STREAM	CHARAC	TERISTICS
DATE	TIME	PERA- TURE (°C)	DISCHARGE (CFS)	CONCEN- TRATION (MG/L)	SUSPENDED SEDIMENT DISCHARGF (TONS/DAY)	TOTAL SEDIMENT DISCHARGE (TONS/DAY)	WIDTH (FT)	MEAN DEPTH (FT)	MEAN VELOCITY (FPS)
DEC 10, 1969	1800	12.0	1400	1200	4540	4590	30.3	8.5	4.00
DEC 30	1545	14.0	440	22	26	33	25	6.6	2.9
JUN 12, 1970	1430	26.5	120	200	65	70	10.2	5	3.2
JUN 14	0900	28.0	1.6	25	.11	.17	1.8	.2	.45
JUN 15	1200	28.5	12	30	.97	.17	2.0	.6	.50
JUN 16	1300	28.0	.88	15	.04	.04	1.5	.1	. 2

Figure 33.—Format table 5, total sediment discharge.

1970

SEPTEMBER

9

OCTOBER 1969

SEDIMENT, WATER YEAR

SUSPENDED

Ч

D I STR I BUT I ON

PARTICLE-SIZE

a water density of 1.00 g/cc and a sediment density of 2.65 g/cc. They are calculated to three significant figures in 2-percent steps. Computation errors from use of these factors will be less than 1 percent and will average zero.

Values of concentration determined in the laboratory as parts per million should be converted to milligrams per liter by laboratory personnel on the form, sediment-concentration notes (fig. 5). For concentrations under 16,000 ppm, the values of parts per million and milligrams per liter are assumed to be equal. For concentrations over 16,000 ppm, the corrected value in milligrams per liter should be penciled plainly below the concentration on the form as a reminder to the user that parts per million is different from milligrams per liter. In any event, personnel who transfer the data from the laboratory forms for other uses are responsible for the correct use of the data and should check the basic data to determine that the conversion was correctly made.

After the data are converted to milligrams per liter, the only coefficient needed to compute sediment discharge is k, which is equal to 0.0027 in the English system of units. The coefficient, k, includes the conversion of water discharge and is a constant for a given system of measurement. The derivation and value of k to determine tons per day in English and metric units follows.

If the weight of a cubic foot of watersediment mixture is assumed to be 62.4 pounds, then

and the computation of sediment discharge is

 $Q_s = Q_w \times C_s \times 0.0027$ tons per day (English short tons).

A metric ton is 2,204.62 pounds avoirdupois, or 1.1023 English short tons, and a cubic meter is 35.31 cubic feet. To convert English short tons to metric tons multiply by 0.9072, or substitute factor 0.00245 in place of 0.0027 in the preceding equation.

	COH T H		SIS	VPWC	PWC	PWC	VPWC	VPWC	VPWC	VPWC	VPWC	VPWC	
		INDICATED	1.00 2.00										
		ERS)	.500	100									
ATER)		LIMET	.250	93	ļ	ł	100	1	100	100	100	100	
ED W	SIZE	N MIL	.125	68		ł	66	100	66	98	97	97	
STILL	PARTICLE SIZE	ZE (I	.062	58	100	100	94	97	93	06	83	88	
Id NI	PART	HE SI	.031										
, W,		HAN TH	016	41	66	96	63	78	67	68	52	63	
TUBE		LER T	. 800.										
ATION		IT FIN	004.	36	78	71	44	59	50	48	37	46	
CUMUL		ERCEN	002.										
SIEVE; V, VISUAL-ACCUMULATION TUBE; W, IN DISTILLED WATER)		CONCEN- SEDIMENT PERCENT FINER THAN THE SIZE (IN MILLIMETERS) INDICATED TRATION DISCHARGE	(MG/L) (TONS/DAY).002 .004 .008 .016 .031 .062 .125 .250 .500 1.00 2.00	1390	18600	21800	116000	86900	189000	89200	92300	103000	
SIEVE; V,				254	11400	10100	17900	39600	46000	36000	39900	12600	
P, PIPET; S,		DISCHARGE	(CFS)	2020	603	800	2410	784	1470	885	8260	3040	
۹ ۹	WATER Tem	PERA- Ture	()°)	7.0	11.5	16.5	14.0	19.0	19.0	14.5	21.0	17.5	
			T4 ME	1050	0550	1255	1700	1615	1130	1000	1355	1235	
,			DATE	27, 1970	24	26	10	10		19		•	
				۹R	¥	¥	N	ц	ц С	5	0G	5	

AAAU UUAAA

PARTICLE-SIZE DISTRIBUTION OF SURFACE BED MATERIAL, WATER YEAR OCTOBER 1969 TO SEPTEMBER 1970 (METHODS OF ANALYSIS: O, OPTICAL ANALYZER; P, PIPET; S, SIEVE; V, VISUAL-ACCUMULATION TUBE)

			. –	NUMBER						PAF	RTICL	E SIZE	2				METHOD
			TEM- PERA- TURE	OF SAM- PLING	DISCHARGE	PERC	ENT F	INER	THAN	THE S	51ZE	(IN MI	ILLIM	ETERS)	INDIC	ATED	OF ANALY-
	DATE	TIME		POINTS		.004	.062	.125	.250	.500	1.00	2.00	4.00	8.00	16.00	32.00	515
NOV	29, 1969	1515	9.0		440	0	3	10	53	97	100						sv
	30, 1970				206	0	1	10	48	86	97	98	99	100			sv
	22				143	0	1	10	51	86	98	99	99	100			sv

Figure 35.—Format table 7, particle-size distribution of surface bed material.

Water-discharge data in the metric system usually are reported in cubic meters per second, and sediment-discharge data in metric tons per day. If water data are in cubic meters per second, the weight of 1 cubic meter of water is 1 metric ton, and the time interval is 24 hours; thus

and the computation of sediment discharge becomes

 $Q_s = Q_w \times C_s \times 0.0864$ (metric tons).

Computation of subdivided days

The term "subdivide" refers to the division of data for a calendar day into shorter periods of time to obtain correct daily mean values of water or sediment discharge when one or both change beyond certain limits during the day. The values for these periods are then summed to obtain values of water or sediment discharge.

The basic reasons for subdivision of data for a day to compute water discharge and the reasons for subdivision to compute sediment discharge are different. Computation of daily mean water discharge requires subdivision of the day when the gage height is changing because the relation between stage and discharge is curvilinear, and the daily discharge corresponding to a mean gage height will be less than the true mean discharge. Computation of a daily mean sediment discharge requires subdivision of the day when both water discharge and concentration are changing because the average of the products of two variable quantities is not the same as the product of the averages of the quantities. Subdivision is not required if either discharge or concentration is constant during the day.

Water-discharge values for subdivided days frequently are calculated in advance of sediment calculations and are listed on the recorder chart of the gage-height record. However, because the reasons for subdivision to determine water discharge and the reasons for a subdivision to compute sediment discharge are not the same, a subdivision for water discharge may not be adequate for sediment discharge. Colby (1956b) sums it thus:

A common source of errors is incorrect division of the day into parts. An inept type of subdivision may be little better than no subdivision at all. Far too often the same separation of the day into parts is used for both the computation of water discharge and the computation of sediment discharge.

The concentration graph should be prepared prior to computation of water discharge, and the initial subdivision made on the basis of change in both water discharge and concentration. If subdivisions and computations of water discharge were made prior to preparation of the concentration graph, a new subdivision of water discharge may be necessary to compute sediment discharge. The new subdivision may divide the day into different parts and into a larger number of parts. Any new water-discharge values added to the subdivision should be carefully computed and checked. Usually when a period is subdivided into a larger number of intervals, the accuracy of the value of daily mean water discharge is improved, and the finite value of the daily water discharge is increased. If the increase in daily mean water discharge over the previously computed value is significant, the original value should be revised. This significant increase will not occur, however, if the original subdivision was adequate.

Figure 36 is a guide to indicate if subdivision is necessary to maintain computations within specified limits of error. The maximum and minimum discharge and the maximum and minimum concentration are determined for the day, or for a trial period, and the ratios of these values plotted on the guide. In practice, a precise computation and plotting of ratios is not necessary, and a quick mental calculation of the ratios and a visual location of these on the guide will suffice.

Subdivided days are computed primarily by two methods which are referred to herein as the mean-interval method and the mid-

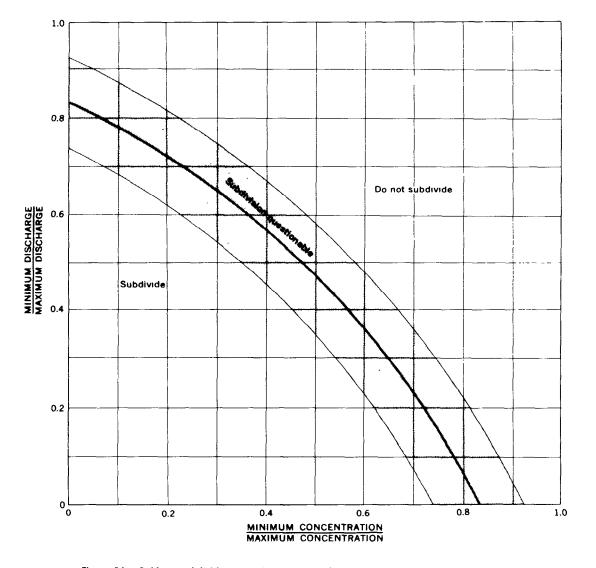


Figure 36.—Guide to subdivision, assuming accuracy about 5 percent (after Colby, 1956b).

interval method. These methods are analogous to the mean-section and midsection methods of computing streamflow from water-discharge measurements described by Buchanan and Somers (1965).

Mean-interval method

The mean-interval method of computation of sediment discharge requires the determination of the mean water discharge and mean sediment concentration for several short periods of time during a day. The products of the two variables for each interval are then time weighted and summed to obtain daily mean discharge. This method is convenient if an analog record of gage height is available and mean values of gage height (or water discharge) and concentration can be determined graphically. Graphical determination of these values for an interval of a day may be made in the manner described in the section on "Computation of daily mean concentration."

An example of the mean-interval method is given in table 3. The daily sediment-discharge data were computed by two procedures:

- 1. Sediment discharge computed for each time interval and summed to obtain daily discharge (col. 7).
- 2. Sediment discharge computed by summing the product of water discharge, time interval, and concentration in the calculator and multiplying the summation by k/T, where k is the coefficient, 0.0027, and T is the summation of hours in the time-interval column, or number of intervals used in subdividing the day. In this latter procedure, column 7 is not computed. This procedure is the one most commonly used.

An advantage resulting from use of the first procedure is the listing of actual sediment-discharge values or tonnage for each time interval (col. 7, table 3). This method is most convenient to use if the computations are made by slide rule or by a calculator that does not cumulatively multiply. These data are useful for direct application to special studies of water and sediment-discharge relations.

Midinterval method

The midinterval method of computation of sediment discharge assumes the values of water discharge and sediment concentration for a specific time represent the average values for the time interval that extends ahead and behind halfway to the preceding and following clock times. The term "midinterval" is not precisely descriptive because if the successive time intervals are unequal the values are not at the midpoint. However, if the time intervals are equal then the values are at the midpoint of the time interval. This method is usually used when the gageheight record is a printout of a digital record.

An example of the midinterval method is given in table 4. Values of gage height and concentration are tabulated for the clock time (col. 1), and the time interval and water discharge are computed. Sediment discharge is computed by summing the product of the water discharge, sediment concentration, and time interval in the calculator and multiplying the summation by k/T, where k is the coefficient, 0.0027, and T is the summation of hours in the time-interval column, or the number of intervals used in subdividing the day.

If subdivision was not used for the data in table 4, then the 24-hour sediment discharge would be

 $Q_s = Q_w \times C_s \times 0.0027 = 2,260 \times 1,830 \times 0.0027$ = 11,200 tons per day.

The error caused by not subdividing is 25,400-11,200=14,200 tons per day,

and the ratio of Q_s computed from daily mean values to Q_s computed by the sub-

Table 3.-Computation of subdivided day, mean-interval method

- Time interval (col. 1): Time during which there is insignificant change in water discharge and sediment concentration or during which the maximum change is within limits specified by the allowable range in stage and by the guide to subdivision (fig. 36). The number of intervals used will depend on the accuracy required.
- Gage height (col. 2): The average gage height for the time interval corrected for any mechanical errors. It is usually determined graphically.
- Shift correction (col. 3): The adjustment applied to gage height to account for variations in the rating. This column may be omitted if no shift is needed.
- Water discharge (col. 4): The mean discharge for the interval; obtained from rating tables.
- Sediment concentration (col. 5): The mean concentration for the interval; usually determined graphically from the concentration graph.
- Interval \times concentration (col. 6): The product of the time interval (col. 1) and concentration (col. 5). The daily mean concentration, published in the concentration column of the sediment-discharge sheet, is computed by dividing the sum of col. 6 by the sum of col. 1. Col. 6 can be omitted for simple computations; col. 1 \times col. 5 may be accomplished mentally and entered directly in the calculator and multiplied by water discharge (col. 4).
- by water discharge (col. 4). Computer sediment discharge, procedure 1 (col. 7): The discharge for the interval is computed by multiplying the product of the qtc (col. $4 \times col. 6$) by 0.0027 and dividing by the time, or total length of the interval. The calculation for periods divided into 24 intervals may be simplified by using the divisor, d. If this is done, col. 6 may be omitted. Values of d, given in footnote 1, are T/tk, where t is the interval, k, is 0.0027, and T is 24. Tons per day is the numerical sum of interval values. An example of the computation of the first and last intervals, using d from fcotnote 1, is

tons per interval =
$$\frac{\text{col. } 4 \times \text{col. } 5}{d} = \frac{q \times c}{d} = \frac{174 \times 8}{1,481} = 0.94$$

= $\frac{q \times c}{d} = \frac{4,920 \times 3,500}{8,889} = 1,940.$

Computed sediment discharge, procedure 2:

Total interval = $T = \sum (\text{col. 1}) = 24$ Daily mean gage height = $\frac{\sum (\text{col. 1} \times \text{col. 2})}{T} = \frac{144.29}{24} = 6.01$ Daily mean water discharge, $Q_{**} = \frac{\sum (\text{col. 1} \times \text{col. 4})}{T} = \frac{54,286}{24} = 2,260$ Daily mean concentration, $C_* = \frac{\sum (\text{col. 1} \times \text{col. 5})}{T} = \frac{43,648}{24} = 1,820$ Tons per day, $Q_* = \frac{\sum q(tc)}{T} = \frac{\sum (\text{col. 4} \times \text{col. 6})}{T} \times 0.0027$ $= \frac{225,767,432}{24} \times 0.0027 = 25,400$

Time interval, t (1)	Gage height (ft) (2)	Shift correction (ft) (8)	Water discharge, q (cfs) (4)	Sediment concentration, c (mg/l) (5)	Interval \times concentration, tc (col. 1 \times col. 5) (6)	Computed sediment discharge, procedure 1 (tons/interval (7)
6	4.17	0	174	8	48	1
	4.32	ŏ	234	80	60	2
2 2	4.57	Ō	855	80	160	6
2	4.66	0	403	100	200	9
ī	4.83	Ó	498	180	180	10
1	5.75	0	1,210	650	650	88
1	6.80	Ó	2,520	1,350	1,350	383
1	7.55	0	3,710	2,350	2,350	981
1	8.10	0	4,760	3,350	3,350	1,790
1	8.39	0	5,380	4,600	4,600	2,780
1	8.53	0	5,690	6,200	6,200	3,970
1	8.62	0	5,890	6,950	6,950	4,610
1	8.58	0	5,800	5,500	5,500	3,590
1	8.47	0	5,550	4,600	4,600	2,870
1	8.37	0	5,330	3,950	3,950	2,370
1	8.18	0	4,920	3,500	3,500	1,940
Total 24	144.29		54,286		43,648	25,400
Weighted						
mean	6.01		2,260		1,820	_

¹ Divisors used to compute sediment discharge in procedure 1 for a given time interval:

Fime interval, t (hr)	Divisor, d	Time interval, t (hr)	Divisor, d
1	8,889		683.8
2	4,444	14	634.9
3	2 963	15	592.6
4	2,322	16	555.6
5	1 778	17	522.9
6	1 481	18	193 8
7	1,270	19	467.8
8	1.111	20	444 4
9	987 6	21	423.3
10	888 0	22	404.0
11	808 1	23	386 5
12	740 7	24	370 37

Table 4.—Computation of subdivided day, midinterval method

Clock time (col. 1): The actual time, given in 24-hour time, for which values are tabulated. Sufficient values must be chosen to assure that the maximum change in successive values of water discharge and sediment concentration is within the limits specified by the allowable range in stage and by the guide to subdivision (fig. 36).

Time interval (col. 2): The sum of one-half the time back to the preceding clock time and one-half the time to the following clock time. The first interval, of 2.5 hours, is one-half the time from midnight (0000 hrs) and 5 a.m. (0500 hrs). The second interval, of 3.5 hours, is one-half the time from midnight to 5 a.m. (2.5 hrs) plus one-half the time from 5 a.m. to 7 a.m. (1 hr). This may also be computed by taking one-half the difference of alternate hours (except the first and last) as follows: First interval: $\frac{0500-0000}{2} = 2.5$ Second interval: $\frac{0700-0000}{2} = 3.5$ Third interval: $\frac{0900-00500}{2} = 2.0$

Cols. 3-6: See explanation for columns 2-5, table 3. Interval × concentration (col. 7): See explanation for column 6, table 3.

	Clock time (hrs) (1)	Time interval, t (hrs) (2)	Gage height (ft) (3)	Shift correction (ft) (4)	Water discharge, (cfs) (5)	Sediment concentration, c (mg/l) (6)	Interval \times concentration, tc (col. 2 \times col. 6 (7)
	0000	2.5	4.17	0	174	8	20
	0500	3.5	4.19	0	182	8	28
	0700	2.0	4.32	0	234	40	80
	0900	2.0	4.60	0	370	90	180
	1100	1.5	4.67	0	408	100	150
	1200	1.0	4.73	Ó	442	120	120
	1300	1.0	5.22	Ó	744	320	320
	1400	1.0	6.22	Ō	1,740	1,000	1,000
	1500	1.0	7.20	0	3,120	1,900	1,900
	1600	1.0	7.83	Ō	4,220	2,850	2,850
	1700	1.0	8.26	Ó	5,090	4,000	4,000
	1800	1.0	8.50	Ō	5,620	5,300	5,300
	1900	1.0	8.56	Ō	5,750	7,000	7,000
	2000	1.0	8.60	Ō	5,840	6,400	6,400
	2100	1.0	8.54	Ó	5,710	4,950	4,950
	2200	1.0	8.41	Ō	5,420	4,200	4,200
	2300	1.0	8.31	Ō	5,200	3,700	3,700
	2400	.5	8.10	Õ	4,760	3,300	1,650
Total		24	144.365		54,168		43,848
Weighted					-		-
mean			6.02		2,260		1,830

Computed sediment-discharge procedure:

Total interval = $T = \Sigma$ (col. 2) = 24 Daily mean gage height = $\frac{\Sigma$ (col. 2 × col. 3)}{T} = \frac{144,365}{24} = 6.02 Daily mean water discharge, $Q_{\bullet} = \frac{\Sigma$ (col. 2 × col. 5)}{T} = \frac{54,168}{24} = 2,260 Daily mean concentration, $C_{\bullet} = \frac{\Sigma$ (col. 2 × col. 6)}{T} = \frac{43,848}{24} = 1,830 Tons per day, $Q_{\bullet} = \frac{\Sigma q}{T}$ $k = \frac{\Sigma$ (col. 5 × col. 7)}{T} × 0.0027

$$= \frac{226,035,716}{24} \times 0.0027 = 25,400$$

divided-day method is

$$\frac{11,200}{25,400} = 0.44$$

In practice, this ratio is computed for each subdivided day during the computation process because it gives an index to the value, or need, to subdivide. For large streams, it also provides an excellent basis for experienced personnel to estimate the need to subdivide the next day. As the flood peak recedes and concentration values decline, the ratio for each successive day will approach 1.00; subdivision is not necessary when the ratio for the preceding day is 0.95 or greater.

The excellent agreement between average values for corresponding variables in tables 3 and 4 is notable. Such close agreement indicates sufficient subdivision. A similar check also is possible when additional intervals are added to previously subdivided water-discharge values. If weighted gage-height values are the same, but water discharge increases significantly, then the original subdivision of water discharge was insufficient. This reasoning can be extended through computation of sediment discharge. If additional intervals fail to significantly change daily mean values of water discharge, concentration, or sediment discharge, then the period is subdivided into sufficient intervals.

Data for the examples in tables 3 and 4 are taken from the gage-height and the concentration graph in figure 37. Data may be recorded and computations made on a copy of the gage-height chart or on a separate sheet.

Sediment-Discharge Worksheet

At this point a continuous sediment-concentration graph has been drawn and reviewed. Concentration values for daily periods and subdivided periods have been computed, and all coefficients applied. Sediment discharge and time-weighted concentrations have been computed for each subdivided day. These values and computations now are ready for the next step—entering the data on the sediment-discharge worksheet (fig. 38).

The sediment-discharge worksheet and the tabulation of particle-size analyses (fig. 2) are the major products of the basic-datacollection program and are published annually. These are the forms from which copies are made to furnish preliminary information to the various data users and which are furnished to the records-processing center. These forms, therefore, should be filled out completely, neatly, and accurately-both in content and editorially. Editorial procedure and rules are based on the United States Government Printing Office "Style Manual," revised edition, January 1967. Specific rules are further explained in "Suggestions to Authors of the Reports of the United States Geological Survey," fifth edition, 1958.

If the electronic computer program is used, values of water discharge and sediment concentration are compiled and punched on data cards; sediment discharge is not computed by the method that follows. Part of the output of the computer program is a listing of sediment data in a form suitable for publication. Records computed manually on the worksheet also are sent to the recordsprocessing center to be listed by computer for publication. An example of the offset copy—prepared from data on the worksheet (fig. 38)—produced by the automatic dataprocessing equipment is illustrated in figure 29.

The technical and editorial accuracy of the product depends on each individual who has a part in its preparation. Each value should be checked for accuracy, and each tabulation should be evaluated for editorial correctness. Data should not be listed with the intent of allowing the checker, or reviewer, to discover and correct errors. It is much easier to keep errors out of the tabulation than to correct them at a later date, if ever.

A list of the technical and editorial rules used in the preparation of the sedimentdischarge worksheet (fig. 38) follows. *Heading of table:*

- 1. Enter the number and name of station in capital letters.
- 2. Fill in the correct time period. If the period is not a complete water year, change the months to agree with period of record.

3. Fill in headnote, if applicable. See rules for footnote E in section on "Footnotes." Body of table:

- 1. Mean discharge:
 - (a) Copy the water discharge from form 9-192 or printout. Use exact figures listed in surface-water data because perfect agreement must exist between data listed in surface-water and water-quality publications.
 - (b) Omit footnotes in this column.
 - (c) Omit explanation or qualifications concerning water discharge.
- 2. Mean concentration:

COMPUTATION OF FLUVIAL-SEDIMENT DISCHARGE

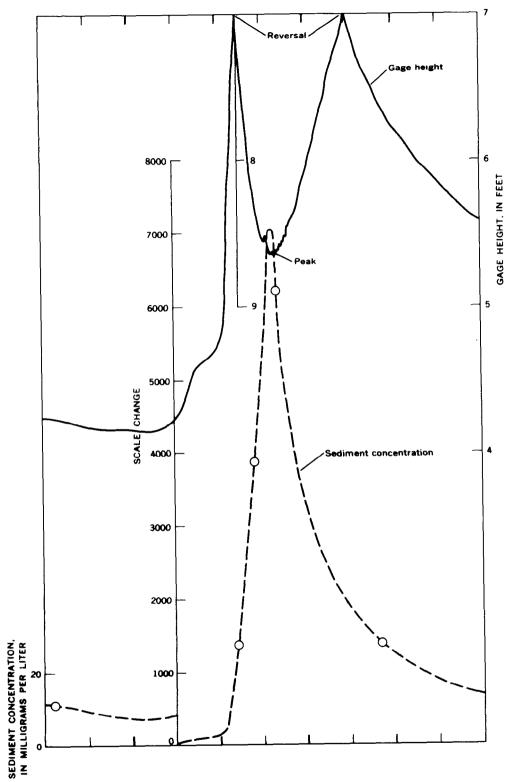


Figure 37.—Gage height and sediment concentration for a subdivided day.

UNITED STATES DEPARTMENT OF THE INTERIOR-GEOLOGICAL SURVEY WATER RESOURCES DIVISION

11-582 × SAMPLE CREEK NEAR SOMEWHERE, CALIF Sespended sediment, weter year Ostaber [967] to September 1962.

L		October			Novamb			Decembe	н Т			January			Pobruary			March	
ŀ	No.		nd partiment				-	Suspende					d andiment	-	Suspende	d undiment	then r	· · ·	l nediment
	(cfs)	testion (ng/1)	Tons per day	diacharga (cfa)	thean communication (mg/1)	Taka por day	diacharge (cfs)	(mg/1)	Teas per day	Dey	discharge (cfa)		Tons per day	discharge (cfe)		Tons par day (discharge (cfs)	(mg/l)	Tuas per duy
	0.69	#	e.ei		247		4410	9/3	5 3820	1	160	Zeez		_ (BAR)	<u>.</u>	- 16.100	190	. 25 -	
	. 50		•	42	40	- T	710	3/8	<u>670</u>	1,	\$10	-	16,500		5 . 10			630	21
ł		I	•	51	22		470	140	171		930	2000	17,600		4100. 5650	10200 \$670		540 515	51
Ĺ	0		0	21	4	ريع. ادر ا	205	-70 -35	45.	5		5,200	//		4900	13100	371	420	43
F		2	0	14	4	15		20	\$6		1,270		5 29 800		3400	1130		600	69
Ì	30	3	0	11	3	16		11	3.9	7	761	6.900	12,709		2700	5100		540	5
	. 60		0,	10	/00	5 46		9	2·1	•	770	4.600	1,120		\$400	+***		415	40
ł	.10	<u>í</u>	•	311	651	5 690	241	65	42		780		10,400		2.050	3440			- 34
┝	10			<u>459</u> 254	542	5 1,240	330	12/	5 2430		45/	6550	_/ 6 100		1660	2520		410	
ł	10 10		0	340	374	5 381 5 289	411	110	125	12	1,100	15.0	4 27300) 4 25100,	490	1780	25/0		432 504	31 5
İ	1.1	2	.01	148	65		275	40	30	11	995	-	17,700	-	1220	1 540			3
ĺ	1.4	4	.02	10			218	22	13	14	973	5,650	14600	· ·	1050	1310			23
L	4.7	5	. 06	55		2		20		15	1,150	4.600	20200	434	910	1070		288	
ł	61	4	o7,	42		*		17		3.6	1,220				120	1010			2 1
		4	70	37				10	3.6	17	1,270	4950			110	741			3
1	6.1	ع ا	0 5	33		.14	117	10	3.2 54	1.	1420	¥ 50		427	780	\$\$5 1,010		245	1
l	61.	2	03 03	30 21				/5	5 461	30	1.500	2450	34000	462	110	1250	· ·	276	2
F	<u> </u>	1	.01	33		01			259.000	33	1,260			4 90	(1320		\$15	2
i	6.1		02	42						11	6030			440	170	1,310		415	34
[61.		02	39			حص بک ا	463-00	2,10	23			\$154000		1 50	1.020		410	3
ļ	66,	. /	02	42		//	معور ج	34,600	1,020	34		14,400	5119,000	406	600	45E		401	31
L		+/	02	371	369	\$ 561			346.000	25	1.470	1)	32900	316	<u> 510</u>	603		110	2
ł		2	. 04	2 32					267-	27	1,150	6. 9 00	21400	378 755	570		274	315	2
ł	90	2		148		5 /, 330			135,000	28	962	6,150 5360	12,600	554	3340	2,340	2.57	3 lo 168	
	19 10	226	, 15 5 60	41 6 595		5 723			44 9	20	900	5,160	12,500	\$17	(140	1590	240	2 77	
ł	48	66	36	315					34,100	30	1,170	5960	18,500				257	212	
Ĺ	22	10	.59				1,070	1950	13,000	31	1290	5,960	20,700				274	301	1
	26070		69.88	4,479	1	5.563.43	\$1 ,257		9362012	Total	37,512		910,310	17,672		124185	4838		11,3
T		April			May			June			Γ.	July			August		1	September	
t	214	83a	253	594	1350	\$/70	246		73	1	53	/2	1.7	7.0 :	\$	0.10			
	320	456		504		137-	223	41	59	2		,• ` •	1.3	. 5.7	+	.06	-		
	274	99 0	244	4 6 2	1 1 60	1060	214		. 52	,	46,	- 7	.17	57.	4	.06	i .		
ŀ	279	344		441		, 174	-	7.	31	1	46.	· .	87	61, 6.2	. +	. 10			
ł	290	330		494	414		<u>196</u> , 92	<u>, n</u>	21	1.	42	<u> </u>	<u>. 9/</u>			.06	.		<u> </u>
1						449													
L	274	120	163	448				54	1 1	,	4/	6				04			
	274 284	→ ¹²⁰	163 230	462	. 310	387		54 18	1 1	,	4/ 33. 30.	، د	53	10.	• *	03			
	274	320 4/10	163	462	310 340 355	418	184- 156 (4)	48 36	27	,	33. 30. 28.	6 7	53 .44_	10 14 21	▲ ` 3	I I	• • • •		
	274 284 1010 702	320 300 4/10 2730	163 230 5 12,100 5 3,640	462 485	310 140 356 380	418	184- 156 (4)	48	27	,	33. 30. 28. 28.	د • ۲	53 .44	10 14 21 11	•	03 04	• • • •		
	274 284 1910 702 434 357	220 4,110 4,720 	163 230 \$ 12,100 \$ 3,640 773 511	462 488 462 <u>476</u> 514	310 140 355 380 430	418 445 478 593	184- 156 (4) 127; 123	48 93 - 34 - 31	27 19 14 11	, * 10	33. 30. 28. 28. 21.	• • 7 • •	53 .44 .53 .31 .30	30, 1,4, 28, 1,7,	* 3 1 2	03 04 02 01	• •		
	274 284 1910 292 434 357 427	220 300 4/10 4730 610	163 230 5 12,800 5 3,640 773 511 703	442 488 442 <u>476</u> 514 531	310 340 356 310 410 445	418 443 478 593 646	184- 156 (4) 	95 93 93 91 91	27 19 14 11 10 1 7	7 8 10 11 12	33. 30. 28. 28. 28. 28. 21.	s 7 4 4	53 .44 .53 .31 .30 .17	30, 28, 28, 1,7, 1,1,	4 3 1 1	03 04 02 01 01 01	· · ·		
	274 284 1910 102 434 357 427 448	220 4/10 2730 640 630 610	163 230 \$ 12,800 \$ 3,640 773 511 703 761	462 485 442 <u>476</u> 5/4 538 578	310 340 355 310 430 445 445	418 443 488 593 646 702	184- 156 (4) 	48 93 34 37 26	27 19 14 11 10 87 68	7 4 10 11 12 13	33. 30. 28. 28. 28. 28. 28. 27. 28.	L 3 4 4 4 4	53 .44 .53 .38 .30 .17 .30	50, 2,4, 28, 22, 1,7, 1,1, 1,1, 1,1,	* 1 2	03 04 01 01 01 01	· · ·		
	274 284 1010 102 434 357 437 437 418 476	320 4/10 4/10 4/20 630 630 700 700	163 230 5 12,800 5 3,440 773 511 703 768 400	442 485 442 5/4 5/4 538 578 572	310 340 356 310 410 445	418 443 488 583 646 702 602	184 156 (4) 127, 123 114 97 9/*	vs 55 57 51 57 26	27 19 14 14 11 10 87 68 59	7 8 10 11 12	33. 30. 28. 28. 28. 28. 27. 28. 30.	L 3 4 4 4 4	53 .44 .53 .38 .70 .27 .30 .24	50 214 28 22 1.7 1.7 1.1 1.1 1.1 50	4 2 1 2 1 3	03 04 02 01 01 01	· · ·		
	274 284 1010 102 434 357 437 437 437 476 892	220 300 4/10 4730 530 530 530 530 530 530 530 530 530 5	163 230 \$ 12,800 \$ 3,640 773 511 703 761	442 +#5 442 476 5/4 5/4 578 578 572 54	310 340 355 310 430 445 390	418 445 498 593 646 702 602 568	184 156 (4) 127; 123; 124 47 47 47; 47; 77	15 36 37 31 14 26 12 20	27 19 14 10 10 17 68 54	7 8 10 11 12 13 14	33. 30. 28. 28. 28. 28. 27. 28. 37. 28. 30. 30.	L 7 4 4 4 4 4 4	53 .44 .53 .38 .40 .24 .24 .24 .24 .23	30 . 28 . 22 . 1.7 . 1.1 . 1.1 . 91 . 80 .	4 1 1 1 1 2 1 2 1 2 1 2	03 04 02 01 01 01 01	· · ·		
	274 284 1010 102 434 357 427 427 427 448 476 892 1,390 900	220 500 (//0 (/720 630 700 	163 230 3 12,200 5 3,640 773 511 703 768 900 8 11,000	442 +#5 442 476 5/4 5/4 578 578 572 54	310 340 430 430 440 370 370 378	418 445 498 593 646 7-2 602 568 4 <i>a</i> / 470	184 156 (4) 127, 123 114 97 97 9/*	+6 53 54 51 51 51 51 51 51 51 51 51 51	27 19 14 11 10 19 68 54 54 42 36	7 4 10 11 12 13 14 15 16 17	33. 30. 28. 28. 28. 37. 28. 30. 30. 30. 28. 24.		53 .44 .53 .38 .29 .24 .24 .24 .23 .19	3 0 , 2 8 , 2 2 , 1.7 , 1.1 , 91, 80 , 10, 10, 10, 10, 10, 10, 10, 10	* ³ 2 1 3 2 2 1 3 1 2	03 04 02 01 01 01 01 01	· · ·		
	274 284 1010 702 357 427 427 427 427 427 427 427 427 427 42	220 800 (//0 (720 610 780 	163 230 3 12,800 5 3,440 773 511 703 768 400 \$ 11,000 \$ 7800 \$ 38000	462 485 442 476 574 574 572 587 580 572 580 572	310 355 310 410 410 410 310 310 310 310 244	418 445 498 593 646 7-2 602 568 44/ 470 573	184- 156 (4) 127, 123 124 97 9/- 97 72 72 77,	46 93 34 37 37 31 24 26 ,7 20 ,7 20 ,7	27 19 14 11 10 89 54 54 42 35 31	7 4 10 11 12 13 14 15 16 17 18	33. 30. 28. 28. 28. 27. 28. 30. 30. 30. 28. 24. 19.	5 3 4 4 5 3 3 3	53 .44 .53 .38 .39 .24 .24 .24 .24 .17 .17	50, 28, 21, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1	* 3 2 2 2 3 3 2 3 1 1	03 04 02 01 01 01 01 0 0 0 0 0 0 0	· · ·		
	274 284 1010 702 357 427 427 427 427 427 427 427 427 427 42	220 800 4/10 4720 610 720 720 720 	163 230 3 12,100 5 3,640 773 511 703 511 703 900 8 11,000 8 11,000 8 71,000 510 00 510 00 (03,000	462 405 442 5/4 5/4 572 604 572 872 872 872 872 872 872 872 872 872 8	510 340 355 310 445 445 390 370 370 370 244 250	418 445 438 593 646 7-2 602 568 447 470 373 336	1844 156 (4) 127, 123 114 97 97 77 72 77 72 77, 12,	48 31 31 31 31 31 31 31 32 32 ,3 26 ,3 20 ,3 20 ,3 20 ,3 20 ,3 20 ,3 20 20 20 20 20 20 20 20 20 20	27 19 14 10 17 68 54 54 31 31 25	7 4 10 11 12 13 14 15 16 17 18 19	33. 30, 28. 28. 28. 37. 28. 30. 30. 28. 24. 19. 14.		53 ,44 ,53 ,38 ,70 ,24 ,24 ,24 ,24 ,19 ,19 ,15 ,13	50, 24, 23, 1,1, 1,1, 1,1, 1,1, 1,1, 1,1, 1,1,	* 3 2 2 2 3 3 2 1 3 2 1 1	03 04 02 01 01 01 01 0 0 0 0 0 0 0 0 0	· · ·		
	274 284 1010 7521 434 357 434 476 476 476 572 448 476 572 566 2,530	320 5=== (//0 /72e (73= (73= (73= (74) (74) (74) (74) (74) (74) (74) (74)	163 230 5 12,200 5 3,240 773 511 703 768 900 8 12,000 8 12,000 8 12,000 5 10,000 5 20,000 5 20,000	462 405 442 5/4 5/4 578 578 578 572 580 5/8 572 580 5/9 498 498	310 340 356 310 410 445 370 370 370 370 370 370 370 370 370 370	418 445 488 583 646 7-2 568 48/ 470 373 336 524	1844 156 (4) 127, 123 1)4 97 97 77 77 77 77 77 77 77 77	ve 34 31 37 31 37 31 36 32 ,7 20 ,7 20 ,7 20 ,7 20 ,7 20 ,7 20 ,7 20 ,7 20 ,7 20 ,7 20 20 20 20 20 20 20 20 20 20	27 19 14 10 87 68 54 41 38 31 25 25	7 4 10 11 12 13 14 15 16 17 18	33. 30, 28. 28. 37. 28. 30. 30. 30. 28. 30. 28. 44. 14. 14. 14. 15.	6 7 4 3 3 3 3 3 3 3 3 3	53 .44 .53 .38 .29 .24 .24 .24 .23 .17 .15 .15 .13 .12	50, 24, 28, 22, 1,7, 1,1, 41, 50, ,72, 91, 54, 54, 54, ,21,	* 3 2 2 3 2 3 2 2 1 1 1 2	03 04 02 01 01 01 0 0 0 0 0 0 0 0 0 0	· · ·		
	274 284 1010 7521 434 357 434 476 476 476 576 900 (910 9,660 2,530 2,180	220 5=== (1/10 (72+ (72+ (72+)	143 230 3 12,800 5 2,440 773 511 703 751 703 6 7100 6 7100 7 5000 7 5000 6 7,000 6 7,000 6 7,000 7 5000 7 6 0,000 6 7,000 7 6 0,000 7 6 0,000 7 7 7 7	462 462 462 475 574 578 578 572 580 572 580 572 580 572 580 572 580 572 580 572 580 574 574 574 574 574 574 574 574	310 340 355 380 445 445 445 390 370 370 370 370 370 370 370 370 370 37	418 445 438 593 646 7-2 602 568 447 470 373 336	1844 156 (4) 127, 123 1)4 97 97 77 77 77 77 77 77 77 77	45 31 37 31 37 26 12 ,7 20 ,7 20 ,7 20 ,7 20 ,7 20 ,7 20 ,7 20 ,7 20 ,7 20 ,7 20 12 12 12 12 12 12 12 12 14 12 14 14 14 14 14 14 14 14 14 14	27 19 14 11 10 19 41 54 54 54 54 54 54 54 54 55 56 56 56 56 56 56 56 56 56	7 8 10 11 12 13 14 15 16 17 18 19 20	33. 30, 28. 28. 37. 28. 30. 30. 28. 24. (9. 14. 15. 15.	6 7 4 4 3 3 3 3 3 3 3 3 3	53 .44 .53 .31 .29 .24 .24 .24 .23 .19 .15 .13 .12 .12	50, 24, 23, 1,1, 1,1, 1,1, 1,1, 1,1, 1,1, 1,1,	* 3 2 2 3 2 3 2 2 1 1 1 2	03 04 01 01 01 01 01 0 0 0 0 0 0 0 0 0 0 0	· · ·		
	274 284 1010 7521 434 357 434 476 476 476 572 448 476 572 490 (900 900 (910 9560 2,530	220 800 4/10 2720 630 780 	143 230 3 12,800 5 2,440 773 511 703 751 703 6 7100 6 7100 7 5000 7 5000 6 7,000 6 7,000 6 7,000 7 5000 7 6 0,000 6 7,000 7 6 0,000 7 6 0,000 7 7 7 7	462 462 462 475 574 578 572 572 572 572 572 572 572 572	310 340 355 380 445 445 445 390 370 370 370 370 370 370 370 370 370 37	418 445 488 583 644 7-2 602 568 470 573 336 514 280	1844 156 (4) 127, 123 1)4 97 97 77 77 77 77 77 77 77 77	ve 31 31 31 31 26 12 ,7 20 ,8 16 ,4 25 .2)	27 19 14 10 10 17 68 54 54 54 54 54 56 50 47	7 4 10 11 12 13 14 15 16 17 18 19 20 21	33. 30, 28. 28. 37. 28. 30. 30. 30. 28. 30. 30. 28. 44. 14. 14. 14.	6 7 4 4 3 3 3 3 3 3 3 3 3	53 .44 .53 .38 .29 .24 .24 .24 .23 .17 .15 .15 .13 .12	50, 2,4, 2,2, 1,7, 1,1, 1,1, 1,1, 1,1, 1,1, 50, 1,1, 1,1,	* 3 2 2 3 2 3 2 2 1 1 1 2	03 04 02 01 01 01 0 0 0 0 0 0 0 0 0	· · ·		
	274 284 1010 102 434 3577 427 448 476 892 1,390 2,600 2,500 2,500 1,300 1,300 1,300	220 800 4/10 2720 630 780 	143 230 3 12,100 3 3,400 773 511 703 716 400 8 11,000 8 11,000 5 1,000 5 1,0000 5 1,0000 5 1,0000 5 1,0000 5 1,0000 5 1,0000 5 1,0000 5 1,0000	442 467 442 574 578 578 578 578 578 578 578 578 578 578	310 340 310 410 410 310 110 110 110 244 241 241 241 241 110	418 445 488 583 646 792 692 692 694 49/ 470 573 336 514 290 [94 134	184- 156 (4) 127, 123, 14 97 97, 97, 72, 77, 12, 67, - 69, - 70, - 70,	ve 32 34 37 26 12 ,7 20 ,8 ,6 ,7 ,6 ,7 ,6 ,7 ,7 ,7 ,7 ,7 ,7 ,7 ,7 ,7 ,7	27 19 10 87 68 54 41 38 54 25 30 47 31 35	7 8 10 11 12 13 14 15 16 17 18 19 20 21 21 22	33. 30. 28. 28. 37. 28. 30. 30. 24. 14. 14. 15. 14. 14. 14. 14. 14. 14. 14. 14.	6 7 4 4 3 3 3 3 3 3 3 3 3 3 3 3 3	53 .44 .53 .38 .70 .24 .24 .24 .23 .19 .15 .13 .12 .12 .12	50, 24, 28, 21, 1,1, 1,1, 1,1, 50, 10, 12, 54, 54, 54, 54, 54, 54, 54, 54	* 3 2 2 3 2 3 2 2 1 1 1 2	03 04 01 01 01 01 01 01 01 01 0 0 0 0 0 0	· · ·		
	274 284 1010 702 4367 437 437 448 476 892 1,390 2,560 2,550 1,350 1,350	220 bee 4/10 472p 630 630 630 700 700 700 700 700 7100 7100 7100 71	143 230 3 12,100 5 12,100 5 3,240 773 511 703 768 4 72,00 5 7800 5 7800 5 7800 5 7800 5 7800 5 7800 7 8 7800 7 8 7800 7 8 7800 7 8 7800 7 8 7800 7 9 78000 7 9 780000000000000000000000000000000000	442 462 462 476 578 578 578 578 578 578 578 578	310 340 310 410 410 370 370 370 370 370 370 370 370 370 37	418 445 488 583 646 792 692 692 694 49/ 470 573 336 514 290 [94 134	184- 156 (4) 127, 123, 14 97, 91, 77 72 77, 12, 67, 69, 69, 69,	ys 3L ys <	27 19 14 10 87 68 54 54 54 54 54 55 25 25 25 25 30 47 57 31 31	7 8 10 11 12 13 14 15 16 17 18 19 20 21 22 23	33. 30. 28. 28. 37. 28. 30. 30. 24. 14. 14. 15. 14. 14. 14. 14. 14. 14. 14. 14.	6 7 4 4 3 3 3 3 3 3 3 3 3 3 3 3 3	53 ,44 ,53 ,38 ,30 ,24 ,24 ,24 ,23 ,17 ,15 ,13 ,12 ,12 ,12 ,11	50 24 28 1.1 1.1 50 .12 .12 .12 .17 .21 .17 .21 .17 .21 .17 .21 .17 .21 .17 .17 .17 .17 .17 .17 .17 .1	* 3 1 2 3 2 2 3 2 1 1 1 2 1 1 2 1 1 1 1 2 1 1 1 2 1 2	03 04 01 01 01 01 0 0 0 0 0 0 0 0 0 0 0 0	· · ·		
	274 284 1010 102 434 357 427 427 427 476 970 9,00 9,560 2,530 2,550 1,340 496 1,050	220 bee 4/10 2720 730 730 740 740 4/00 740 9300 2150	143 230 312,000 312,000 773 768 773 768 970 87,000 87,000 87,000 97,000 16,000 1/400 1/400	442 442 445 5/4 5/4 5/4 5/4 5/4 5/7 5/7 5/7 5/7 5/7 5/7 5/7 5/7) 0 0 3 0 0 3 0 0 5 10 5 10 5 10 5 10 1 0 1 0 1 0 1 0 1 0 1 0 1 0	418 433 445 792 641 792 647 470 373 336 514 250 194 134 250 194 134 134 193 97	184- 156 (4) 127, 127, 124 97, 77 72 77, 72 77, 72 77, 72 77, 72 77, 72 77, 72 77, 72 77, 72 77, 72 77, 67, 67, 67, 67, 67, 67, 67,	ve ,, 3c ,, 3c ,, 2c ,, 2c	27 19 11 10 57 54 54 54 54 54 54 54 54 54 54	7 4 10 11 12 13 14 15 16 16 17 16 20 21 22 23 24 23 24 25 34	33. 30, 28. 28. 27. 28. 30. 30. 28. 24. 14. 15. 14. 15. 14. 15. 14. 15. 14. 15. 14. 14. 14. 14. 14. 14. 14. 14. 14. 14	6 7 4 4 3 3 3 3 3 3 3 3 3 3 4 5 4	53 ,49 ,31 ,30 ,20 ,24 ,24 ,24 ,24 ,24 ,24 ,24 ,17 ,15 ,15 ,13 ,12 ,12 ,12 ,12 ,12 ,12 ,12 ,12 ,12 ,12	50 24 21 1.1 1.1 10 10 .12 .12 .12 .12 .12 .12 .12 .12	* 3 2 2 3 3 2 3 2 3 2 3 2 3 2 3 2 3 2 3 2				
	274 284 1010 102 367 436 436 448 476 892 1,970 9,560 2,180 1,360 1,360 1,360 1,2	2120 500 (1/10) 2730 730 730 730 730 730 730 730 730 730	143 230 312,000 312,000 773 51/ 703 768 81/2000 87,000 87,000 89,000 1/400 1/400 1/400 1/400 7,150 8,1/0	442 462 465 574 574 572 572 572 572 572 572 572 572	510 340 350 410 410 410 370 370 370 370 370 370 370 37	418 445 478 646 7-2 646 7-2 646 447 470 573 536 514 280 194 136 103 97 42 10	184- 156 (41) 123 114 47 47 77 72 77 72 77 72 77 72 77 72 77 4/ 47 67 47 67 47 67 47 57 57 57 57 57 57 57 57 57 5	ys 31 37 31 31 31 31 31 31 31 31 31 31	27 19 14 10 87 68 54 54 54 54 54 54 54 54 54 54	7 4 10 11 12 13 14 15 16 16 17 16 20 21 22 23 24 23 24 23 24 27	33. 30, 28. 37. 30. 30. 30. 30. 30. 30. 30. 30. 30. 30	6 7 4 4 3 3 3 3 3 3 3 3 3 3 3 3 3	53 ,49 ,31 ,20 ,24 ,24 ,24 ,24 ,24 ,24 ,24 ,24 ,24 ,24	5 0 2 24 2 2 1.1 1.1 1.1 1.1 1.1 .1 .1 .1	* 3 2 2 3 2 3 2 3 2 3 2 3 2 3 2 3 2 3 2 3	03 04 01 01 01 0 0 0 0 0 0 0 0 0 0 0 0 0 0			
	274 284 1010 752 434 387 438 438 476 872 1,390 5,60 2,530 7,340 5,60 2,530 1,970 495 1,070 495 1,070 495	220 500 (1/10) (2720) (720) (730) (730) (730) (730) (730) (740) (7	143 312,600 5 3,440 773 703 703 703 703 703 703 703 703 70	442 442 445 446 5/4 5/4 5/4 5/4 5/4 5/4 5/4 6/4 6/4 4/4 4/4 4/4 4/4 5/4 5/4 5/4 5	510 140 350 410 410 410 410 10 10 10 10 10 10 10 10 10	418 443 443 644 7-2 646 7-2 646 7-2 647 470 336 513 447 470 336 514 290 134 103 914 103 91 410 103 91	$\begin{array}{c} 184 \\ 156 \\ (41) \\ 123 \\ 114 \\ 97 \\ 97 \\ 97 \\ 97 \\ 77 \\ 72 \\ 77 \\ 72 \\ 77 \\ 72 \\ 67 \\ 69 \\ 69 \\ 69 \\ 69 \\ 69 \\ 69 \\ 69$	ys 31 17 14 12 14 12 13 14 12 13 14 12 13 14 15 16 14 15 16 17 18 19 10 10	27 19 11 10 87 687 54 54 54 54 54 54 54 54 54 54	7 4 10 11 12 13 14 15 16 16 17 16 20 21 22 23 24 23 24 25 34	33. 30. 28. 28. 37. 28. 30. 30. 30. 30. 30. 30. 30. 30	6 7 4 4 3 3 3 3 3 3 3 3 3 3 3 3 3	53 , 44 , 53 , 38 , 27 , 20 , 24 , 24 , 24 , 24 , 19 , 15 , 12 , 12 , 12 , 12 , 12 , 12 , 12 , 12	5 0 , 2 14 , 2 12 , 1,1 ,	* 3 2 3 3 2 3 2 1 2 1 2 1 2 1 2 1 2 1 2 1				
	274 284 1910 702 357 424 426 476 476 476 476 5,60 2,500 1,900 1,500 1,500 1,500 1,500 1,500 495 (,010 1,20 2,500 2	210 500 4/10 2730 780 780 780 780 780 780 780 780 780 78	1230 312100 3240 3240 371 703 718 708 400 51/0000 51/0000 51/0000 51/0000 51/0000 51/0000 51/0000 51/0	442 +45 445 5/4 53 57 580 572 580 572 580 572 580 572 580 572 580 572 580 572 580 572 580 572 580 572 570 575 570 575 570 575 570 575 570 575 570 575 570 575 570 575 570 577 570 577 570 577 570 577 570 577 570 577 570 577 570 577 570 577 570 577 570 577 570 577 570 577 570 577 570 570	210 340 310 410 410 410 100 100 241 210 241 210 110 110 110 110 110 110 11	418 418 593 644 563 471 563 470 573 334 290 194 103 91 425 109 110 125 125		yF y2 y3	27 19 11 10 87 687 54 54 54 54 54 54 54 54 54 54	7 4 10 11 13 14 15 16 17 16 17 18 19 20 21 22 23 24 23 24 23 24 25 24 27 28	33. 30, 28. 28. 37. 28. 390 30 30 30 30 30 14 14 15 15 15 15 15 14 24 14 24 14 24 14 24 24 24 24 24 24 24 24 24 24 24 24 24	6 7 4 4 3 3 3 3 3 3 3 3 3 3 3 3 3	53 , 44 , 53 , 38 , 30 , 24 , 23 , 24 , 24 , 23 , 14 , 15 , 12 , 12 , 11 , 12 , 11 , 12 , 12 , 14 , 12 , 14 , 12 , 14 , 15 , 12 , 14 , 15 , 14 , 15 , 14 , 15 , 15 , 12 , 12 , 12 , 12 , 12 , 12 , 12 , 12	30 24 23 11 11 11 11 10 <td></td> <td></td> <td></td> <td></td> <td></td>					
	274 284 1010 752 434 387 438 438 476 872 1,390 5,60 2,530 7,340 5,60 2,530 1,970 495 1,070 495 1,070 495	210 500 4/10 2730 780 780 780 780 780 780 780 780 780 78	143 312,600 5 3,440 773 703 703 703 703 703 703 703 703 70	442 442 445 446 5/4 5/4 5/4 5/4 5/4 5/4 5/4 6/4 6/4 4/4 4/4 4/4 4/4 5/4 5/4 5/4 5	510 140 355 340 410 410 410 410 150 150 100 100 100 100 100 1	418 418 593 644 7-22 6-2 563 470 373 336 514 100 100 110 110 143 127	$\begin{array}{c} 184 \\ 156 \\ (4) \\ (4) \\ 127 \\ 123 \\ 1)4 \\ 97 \\ 97 \\ 97 \\ 77 \\ 72 \\ 77 \\ 72 \\ 77 \\ 72 \\ 77 \\ 72 \\ 67 \\ 69 \\ 69 \\ 69 \\ 69 \\ 69 \\ 69 \\ 69$	yF y2 y3	27 19 14 10 59 59 50 50 50 50 50 50 50 50 50 50	7 4 10 13 13 14 15 14 15 14 15 14 15 14 15 14 17 20 21 21 23 24 23 24 23 24 25 23 24 25 23 24 25 23 24 25 25 25 25 25 25 25 25 25 25 25 25 25	33. 30, 28. 28. 30, 28. 30, 30, 30, 30, 30, 30, 30, 30, 30, 30,	6 7 4 4 3 3 3 3 3 3 3 3 3 3 3 3 3	53 , 44 , 53 , 38 , 27 , 20 , 24 , 24 , 24 , 24 , 19 , 15 , 12 , 12 , 12 , 12 , 12 , 12 , 12 , 12	50 24. 23 1.7 1.7 1.7 10 10 10 10 10 10 10 10 10 10					
	274 1910 1910 192 192 192 193 193 193 193 193 193 193 193	110 6(1)0 (1)720 500 500 700 700 700 700 7100	163 312100 512100 53240 511 703 710 511 703 7100 51100 51000 5000000	444 407 402 402 402 404 404 404 404 404 404 404	510 140 356 350 410 410 370 370 370 370 370 370 370 37	418 418 513 644 7-22 602 573 334 447 470 524 124 144 144 103 103 103 103 100 100 10,466	184- 156 (4) 127- 97 97 77 72 67 - 69 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 53 - - - - - - - - - - - - - - - - <td>48 31 31 31 31 31 31 32 32 32 32 32 32 32 32 32 32</td> <td>27 19 14 11 10 68 54 41 38 25 25 47 38 25 47 37 47 31 25 47 47 51 -31 -31 -31 -31 -31 -31 -31 -3</td> <td>7 4 10 11 12 13 14 15 14 15 14 15 16 19 20 21 21 22 23 24 25 24 27 26 27 28 29 20 21 21 21 21 21 21 21 21 21 21 21 21 21</td> <td>53. 50. 28. 28. 28. 30. 30. 30. 30. 30. 30. 30. 30. 30. 30</td> <td><pre>c</pre></td> <td>53 ,</td> <td>50 24. 23 1.7 1.7 1.7 10 10 10 10 10 10 10 10 10 10</td> <td>4 3 2 2 3 3 2 3 3 1 1 2 1 1 </td> <td></td> <td></td> <td></td> <td></td>	48 31 31 31 31 31 31 32 32 32 32 32 32 32 32 32 32	27 19 14 11 10 68 54 41 38 25 25 47 38 25 47 37 47 31 25 47 47 51 -31 -31 -31 -31 -31 -31 -31 -3	7 4 10 11 12 13 14 15 14 15 14 15 16 19 20 21 21 22 23 24 25 24 27 26 27 28 29 20 21 21 21 21 21 21 21 21 21 21 21 21 21	53. 50. 28. 28. 28. 30. 30. 30. 30. 30. 30. 30. 30. 30. 30	<pre>c</pre>	53 ,	50 24. 23 1.7 1.7 1.7 10 10 10 10 10 10 10 10 10 10	4 3 2 2 3 3 2 3 3 1 1 2 1 1 				
	274 214 1010 102 102 102 102 102 102 10	130 4(1)0 4(1)0 4(1)0 4(1)0 4(1)0 4(1)0 50 50 50 50 50 50 50 50 50 5	163 312,000 5 3,400 5 3,400 5 7,73 703 703 703 703 703 703 703 703 703 7	444 407 402 476 571 571 572 572 572 572 572 572 572 572 572 572) 310 140 355 360 445 440 340 340 340 340 244 250 244 250 100 100 100 100 100 100 100 1	418 418 418 513 644 702 602 563 470 573 373 336 210 102 91 102 102 102 102 102 102 103 103 103 103 103 103 103 103	$\begin{array}{c} 184 \\ 156 \\ (4) \\ 127 \\ 128 \\ 100 \\$	48 31 31 31 31 31 31 31 32 32 32 32 33 32 32 32 32 32	27 19 14 19 19 19 19 19 19 19 19 10 19 10 11 26 26 26 26 26 26 26 26 26 26	7 4 10 12 13 14 15 14 15 14 15 14 17 18 19 20 21 24 20 23 24 25 24 26 26 26 27 28 26 30 31 Trent	33. 30. 28. 28. 28. 30. 30. 30. 28. 28. 28. 28. 28. 28. 28. 28. 28. 29. 29. 29. 29. 29. 29. 29. 29. 29. 29	C 7 A 4 3 3 3 3 3 3 3 3 3 3 3 3 3 3 3	53 , 49 , 53 , 29 , 20 , 29 , 20 , 24 , 23 , 19 , 12 , 19 , 12 , 19 , 12 , 19 , 12 , 19 , 12 , 19 , 10 , 19 , 10 , 10 , 10 , 10 , 10 , 10 , 10 , 10	5 0 2 4 2 3 2 3 2 3 2 3 2 3 1,7 ,17 ,17 ,17 ,17 ,17 ,17 ,17	4 3 2 2 3 3 2 3 3 1 1 2 1 1 			٦٩,20	3 96
	274 1610 1610 162 162 163 163 163 164 164 164 164 164 164 164 164	1100 4/10 4/10 4/10 4/10 4/10 5 5 5 5 5 6 1/20 5 1/20 1/2	163 312100 512100 53240 511 703 710 511 703 7100 51100 51000 5000000	444 +05 514 514 514 518 518 518 518 518 518 518 518 518 518	510 140 356 350 410 410 370 370 370 370 370 370 370 37	418 418 418 513 644 702 602 563 470 573 373 336 210 102 91 102 102 102 102 102 102 103 103 103 103 103 103 103 103	184- 156 (4) 127- 97 97 77 72 67 - 69 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 67 - 53 - - - - - - - - - - - - - - - - <td>48 31 31 31 31 31 31 31 32 32 32 32 33 32 32 32 32 32</td> <td>27 19 14 11 10 68 54 41 38 25 25 47 38 25 47 37 47 31 25 47 47 51 -31 -31 -31 -31 -31 -31 -31 -3</td> <td>7 4 10 12 13 14 15 14 15 14 15 14 17 18 19 20 21 24 20 23 24 25 24 26 26 26 27 28 26 30 31 Trent</td> <td>33. 30, 28. 28. 28. 28. 30. 30. 30. 30. 30. 30. 30. 30. 30. 30</td> <td>C 7 A 4 3 3 3 3 3 3 3 3 3 3 3 3 3 3 3</td> <td>53 , 44 , 53 , 38 , 30 , 24 , 25 , 17 , 19 , 15 , 12 , 19 , 11 , 10 , 11 , 12 , 10 , 11 , 12 , 10 , 11 , 12 , 10 , 11 , 12 , 12 , 12 , 12 , 12 , 12 , 12</td> <td>5 0 2 4 2 3 2 3 2 3 2 3 2 3 1,7 ,17 ,17 ,17 ,17 ,17 ,17 ,17</td> <td>4 3 2 2 3 3 2 3 3 1 1 2 1 1 </td> <td></td> <td></td> <td></td> <td></td>	48 31 31 31 31 31 31 31 32 32 32 32 33 32 32 32 32 32	27 19 14 11 10 68 54 41 38 25 25 47 38 25 47 37 47 31 25 47 47 51 -31 -31 -31 -31 -31 -31 -31 -3	7 4 10 12 13 14 15 14 15 14 15 14 17 18 19 20 21 24 20 23 24 25 24 26 26 26 27 28 26 30 31 Trent	33. 30, 28. 28. 28. 28. 30. 30. 30. 30. 30. 30. 30. 30. 30. 30	C 7 A 4 3 3 3 3 3 3 3 3 3 3 3 3 3 3 3	53 , 44 , 53 , 38 , 30 , 24 , 25 , 17 , 19 , 15 , 12 , 19 , 11 , 10 , 11 , 12 , 10 , 11 , 12 , 10 , 11 , 12 , 10 , 11 , 12 , 12 , 12 , 12 , 12 , 12 , 12	5 0 2 4 2 3 2 3 2 3 2 3 2 3 1,7 ,17 ,17 ,17 ,17 ,17 ,17 ,17	4 3 2 2 3 3 2 3 3 1 1 2 1 1 				
	274 1610 16200 1620 1620 1620 1620 1620 1620 1620 1620 1620 1620	1100 4/10	1230 312100 3240 3240 311 703 713 713 716 710 710 710 710 710 710 710 710	444 +47 462 476 514 514 513 518 518 518 518 518 518 518 518 518 518	310 355 30 30 30 30 30 30 30 30 30 30	418 418 513 644 7-22 602 524 602 524 103 524 103 103 110 103 127 100 104 105 107 100 100 100 100 100 100 100	$\begin{array}{c} 184 \\ 156 \\ (4) \\ 127 \\ 123 \\ 127 \\ 97 \\ 97 \\ 97 \\ 97 \\ 97 \\ 97 \\ 97 \\ $	48 36 31 31 31 31 32 36 32 36 33 16 34 36 35 16 36 16 37 16 38 16 39 16 31 16 32 16 33 16 34 16 35 17 36 16 37 17 38 17 39 17 39 17 39 17 39 17 39 17 39 18 39 19 39 19 39 19 39 19 39 19 39 19 39 19 39 19 39 <	27 19 11 10 68 54 41 38 54 54 54 54 54 54 54 54 54 54	7 4 10 11 12 13 14 15 16 17 18 19 20 20 21 22 23 24 25 23 24 25 23 24 27 23 24 27 20 30 11 17 17 17 17 17 17 17 17 17 17 17 17	33. 30. 28. 28. 28. 30. 30. 30. 30. 30. 30. 30. 30. 30. 30	C 7 3 4 3 3 3 3 3 3 3 3 3 3 3 3 3 3	53 , 44 , 53 , 32 , 30 , 24 , 24 , 23 , 14 , 15 , 12 , 12 , 12 , 12 , 12 , 12 , 12 , 12	50, 24, 21, 21, 21, 21, 21, 21, 21, 21, 21, 21	4 3 2 1 3 2 1 3 2 1 3 2 1 3 			٦٩,20	3 ¶6
	274 214 100 102 102 102 102 102 102 102	1100 4/10 4/10 4/10 4/10 4/10 4/10 50 70 70 70 70 70 70 70 70 70 7	163 312,100 5 3,400 5 3,400 5 773 703 703 703 703 703 703 703 703 703	444 +47 462 476 514 514 513 518 518 518 518 518 518 518 518 518 518	310 356 300 310 310 310 310 310 310 310	418 418 513 644 7-22 602 563 563 543 513 514 210 144 103 103 110 110 103 103 103 103	$\begin{array}{c} 184 \\ 156 \\ (4) \\ 127 \\ 123 \\ 127 \\ 97 \\ 97 \\ 97 \\ 97 \\ 97 \\ 97 \\ 97 \\ $	ys 3c 31 31 31 26 32 20 13 26 14 26 12 20 13 16 14 25 17 13 18 10 19 10 10 10 11 10 12 10 13 10 14 10 15 10	27 19 14 19 19 66 54 54 54 54 54 54 54 54 54 54	7 4 4 10 11 12 13 14 15 15 14 17 16 18 19 20 21 21 20 21 22 23 24 24 25 24 24 25 24 20 30 0 11 Total Total Total Total	33. 30. 28. 28. 28. 30. 30. 28. 28. 28. 28. 28. 28. 28. 28. 28. 29. 28. 29. 28. 29. 29. 29. 29. 29. 29. 29. 29. 29. 29	c 7 s 4 s 4 s 4 s 3 s 3 s 3 s 3 s 3 s 3 s 3 s 3	53 , 49 , 53 , 29 , 20 , 29 , 20 , 24 , 23 , 19 , 12 , 19 , 12 , 19 , 12 , 19 , 12 , 19 , 12 , 19 , 10 , 19 , 10 , 10 , 10 , 10 , 10 , 10 , 10 , 10	50, 24. 25. 22. 21. 21. 21. 21. 21. 21. 21				٦٩,20	3 ¶6

Figure 38.—Worksheet for annual suspended-sediment discharge.

1

- (a) Copy values of concentration from graph.
 - (1) During low-flow periods, if samples are not collected daily, tabulate concentrations for days actually sampled, and estimate missing concentrations on basis of values for adjoining days and the waterdischarge hydrograph. A concentration will be determined for each day of flow, if possible. On the worksheet the sampled concentrations are tabulated near the right-hand margin of the concentration column; the interpolated, or estimated, concentrations are tabulated near the left-hand margin of the column. Separation of the sampled and estimated values is for convenience only and aids in checking and reviewing the record. Punched cards for automatic data processing must contain both water discharge and concentration.
 - (2) During periods of changing concentration, a concentration graph may have been estimated on the basis of adjacent records and other data. The daily mean concentrations from the estimated graphs should be copied in the concentration column and the footnote A entered in the tons per day column. If a concentration graph cannot be estimated and sediment discharge determined by indirect is means such as a sedimenttransport curve, then enter a leader (...) in the concentration column and the footnote E in the tons per day column. If 6 or more days require the footnote E. then delete E and use a headnote.

(b) Show concentrations less than 0.5 mg/l either as a leader (..) or as zero. If the water discharge is zero, a leader will be placed in the concentration column and a zero in the tons per day column:

Mean discharge	Suspended sediment									
(cfs)	Mean concen (mg/l)									
3.7		0.01								
1.8	1	0								
.18		0								
.18		Ó								
0		Ō								

- (c) Report concentrations in whole numbers only. Values are rounded as instructed in the section on "Significant figures."
- (d) Omit footnotes, with the exception of C, in the concentration column. Avoid the use of composite values. Daily values should be shown whenever possible.
- (e) Copy footnotes recorded on gageheight chart that explain the concentration or sediment-discharge computation in the tons per day column.
- (f) Copy sediment discharge (tons per day) and footnotes for subdivided days at the time concentration values are copied from chart.
- 3. Tons per day:
 - (a) Enter sediment discharge for subdivided days and footnotes at the time concentration values are copied from the chart.
 - (b) Compute sediment discharge for days where average concentration and water values are given. The daily sediment discharge is mean discharge \times mean concentration \times 0.0027.
 - (c) Round sediment-discharge values according to the instructions in the section on significant figures.
 - (d) Estimate sediment discharge for days where concentration is not shown.
 - (e) Plot all values of water and sediment discharge on logarithmic paper in the form of a sediment-transport curve (fig. 39). Check values that

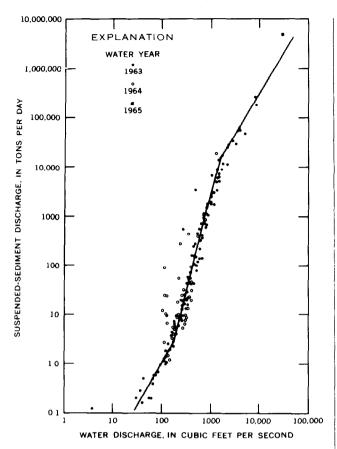


Figure 39.—Relation between daily suspended-sediment discharge and water discharge, Thomes Creek at Paskenta, Calif., 1963-65 water years.

deviate significantly from the average curve for decimal point error, incorrect concentration graph, error in computation, or error in water-discharge values. This plot is useful particularly on large streams having storm runoff events that last several days and where seasonal changes may be significant.

(f) Make revisions as necessary; however, do not try to revise a sediment-discharge value because the value is not consistent with an assumed pattern. The value may not agree because of unusual circumstances, which occur infrequently. These infrequent occurrences often may be verified by comparing them with log plots of water and sediment discharge made for previous years.

- (g) Compute and check monthly and annual totals.
- (h) Insert and check footnotes.
- (i) Compute, record, and check maximum and minimum concentration and sediment discharge.

Station Analysis

A station analysis is a concise summary of (1) equipment used at the station, (2) stream conditions that may affect the quality of the sediment record with respect to particle-size distribution, concentration, and sediment transport, (3) the sampling procedures used, and (4) special methods used to compute sediment discharge.

The analysis of the sediment record may include a summary of chemical-quality sampling procedures and frequency, constituents determined, and special analytical methods. If the sediment records are prepared by the same office preparing the surface-water records, the analyses for the water records and sediment records may be combined. A consolidated analysis including all records collected at one station is preferred because the basic information concerning location, physiography, land use, and climate are common to each analysis. An individual analysis may be necessary when the records are not prepared by the same office or because of conflicts with local needs.

Items that should be considered in a consolidated analysis of sediment discharge, chemical quality, temperature, and rainfall records include:

Equipment type and location.

- Equipment changes and date of change. Important land-use changes that will probably affect the quantity and quality of runoff, such as burn areas, logging, construction, urbanization, and mining.
- Temperature record: completeness, accuracy, and reliability.
- Rainfall record: completeness, accuracy, and reliability, if obtained in connection with gaging station.

56

- Sediment-sampling methods: regular or special techniques, number of sediment-discharge measurements made, and reliability of observer.
- Gage-height record including the datum corrections.
- Water-discharge record including the stage-discharge rating.
- Concentration graph and sediment-discharge computations, and coefficients applied.
- Particle-size analyses: number, methods of analysis, and average percentage of clay, silt, and sand.
- Chemical quality: sampling procedures and frequency, constituents determined, and special analytical methods.
- Remarks concerning other aspects that affect the record.

An example of a station analysis associated with water, sediment, temperature, and rainfall records is given in figure 40, an analysis for chemical quality and sediment records is given in figure 41, and an analysis for sediment records is given in figure 42.

Station Description Heading

A complete station description heading is prepared for each new station when a chemical or sediment data-collection program is started. It is revised when changes in location or operation warrant. Otherwise, no changes are made in the heading on an annual basis except the dates and figures for records available and extremes. A station description heading as published in the water-qualitydata report is shown in figure 43.

Station headings may be updated and typed by the originating office or, once prepared, may be put on punched tape and filed at the records-processing center. This tape is used as input to an automatic typewriter which produces a final copy for offset printing. A major advantage of the taped heading is that the period of records and extremes can be updated without rewriting the remainder of the heading. This eliminates the need for copying, typing, checking, and reviewing the major part of the heading each year.

A copy of the heading (fig. 44) is furnished the office that computed the record. This copy will be used to update the heading for the next water year. The procedure is as follows:

- 1. Use red pencil to update heading.
- 2. Location should be changed only when major change occurs in sampling point or gaging station.
- 3. Drainage area is the value used in the surface-water-data report for the same water year.
- 4. Records available: Update the final data under "Sediment records" to include current records. Example: For 1968 records, line out 1967 in red, and write in 1968.
- 5. First "Extremes" paragraph lists extremes for current water year only. Thus, all dates and extremes probably will be changed.
- 6. Second "Extremes" paragraph shows extremes for period of record, including current year.

The first date, 1961-67, is the period which includes all other dates in the paragraph. In the example, water-temperature records extend from 1961-67, thus the date is not repeated after water temperatures. Sediment-concentration record is 1962-67, which is included in, but different from, the first date; thus, it is placed after sediment concentrations and sediment loads. The last date must be lined out and updated, as shown in the example.

Maximum and minimum values and the dates of occurrence may or may not need to be changed.

Line out and update extremes and dates that change. Check all extremes and dates that remain the same. All dates and figures must be updated or checked.

7. Remarks: Revise this paragraph as required. This paragraph includes information regarding changes in sampling

Observer samples consisting mostly of two bottles were taken at the center of the strame cross section. The samples were analysed separately and compared. It most cases the individual samples afteed fairly well. Several bottles differed considerably, and the unreasonable value was not used in this analysis. Concentrations were plotted on a print of the gage height graph and the concentration curve developed. There ware aufficient samples on the rises to adequately define the concentration curve. No coefficient was needed for the observer samples. Standard procedures were used to defrain the daily sediment 2 Concentration graph and sediment load.--Eleven sediment measurements, depth-integrated at 3 to 5 varticals in the cross section, were made during the year. Particle size.--Ten samples were analyzed. They ranged in concentration from 122 11,600 ppm and in discharge from 780 to 50,000 cfs. Average results of the analysis show 65 percent of material transported to be finer than 62 alterons. All particle-size analyses in the silt-clay range were ande using distilled vater, with a dispersing agent added, as a sattling medium. Discharge.--Rating No. 4 was used with small low-water shifts to the peak of January 12. Rating No. 5 was used direct for the remainder of the year. January 12. Rating No. 5 was used direct for the remainder of the year. Juscharge for periods of no gage height record were estimated on basis of observer's staff readings and hydrographic comperiods with other stations in Remarks.--Water discharge records are good. Suspended sediment records are fair. James L. Cook January 26, 1968 the basin. discharge. once ALI Equipment.--Stevens A-35 recorder with manometer, and tipping hucket rain gage Weisher Thermograph, in steelow shelter. D-49 sediment sampler located on upstream side of bridge 1.2 miles downstream from gage. Measuring cable 0.3 mile below gage Observer equipped with DH-48. OF RED RIVER NEAR CLENWOOD, CALIF. 11-3789.00 (Vater, sediment, temperature, and rainfall) 1966 Water Year STATION ANALYSIS

NORTH FORK

- <u>Hydroiogic changes</u>.--The area upstream varies in elevation from 900 ft to 7,600 ft. Tot's composed chiefly of sandsrone, shile, and greenstones intruded by ultramafic rocks largely altered to serpentine. The combination of sheared rocks, steep slopes and heavy precipitation produces the landslides common to the area. The cover varies from pastures and sagebrush to heavy timber. The timber has been and is continuing to be heavily pogged. Owing to the impermeability of the soil and manife rock, base-flow is poorly sustained and the bulk of the runoff in both water and sediment is storm produced.
- Temperature record, --Record complete except period November 12 to December 4 when Veksler clock stopped. Water temperature observations made at time of sediment samples were available for part of period.

Rainfall record. -- Record complete for the year.

- Sediment record.--Samples were taken about three times a week during low-flows, a day during medium-flow, and one or more times a day during high-flows. Al samples were depth-integrated and considered of fair to good quality.
- Gage-height record,--Record complete except periods January 4-6, January 20 to February 2, and June 19 to July 1 when manometer malfunctioned. Observer's staff-gage readings were available for some of these periods.
- Datum corrections.--Levels of June 20, 1966, found datum of gage to be 901.58 ft above mean sea level, datum of 1929, supplementary adjustment of 1956. The orifice elevation was found to be 3.26 ft and all outside staffs to be correct within 0.01 ft. No corrections applied.
- Rating.--The streambed at the station is composed of sand, gravel, cobblestomes, and a reviarge boulders. A gravel riffie just downstream forms the low-water control. The channel, with banks which are steep and of fairly stable base rock, forms the medium- and high-water control. There is extensive gravel mining 0.5 mile above the station which at times affects the discharge rating and concentration curve.

Sixteen discharge measurements, Nos. 115-130, were made during the current year. Neasurement No. 123 made by an inexperienced hydrographer during a period of rapidly changing stage was not used in the analysis. Sating No. 4 in use at the close of the previous water veat was continued in use until the high-water of Jamuary 12. Neasurements after this date indicate an extensive scour and partial return to conditions similar to that existing prior to the Pecember 1964 flood. Therfore, rating No. 5 was developed for use during the remainder of the year. This rating was based on meaurements and effer the Jamuary flood, is well defined and is the same as all previous ratings above 12,000 cfs.

Page 2 of 2

Figure 40.—Station analysis (water, sediment, temperature, and rainfall) showing format and content.

Page 1 of 2

RIO CHAMA NEAR ABIQUIU, N. MEX. 8-2875.	<u>Chemical quality methods</u> The midmection samples collected at this station are considered to give a reprementation of quality as there is no inflow immediately
STATION ANALYSIS (Chemical Quality and Sediment) 1960 Water Year	above the station. The stream is quite uniform in composition both in depth and width as reflected in three sets of cross-section samples taken throughout the year. Conductance was run on all daily samples prior to compositing. Composite samples are made by discharge-weighted method. Composite analyses during the
Sampling programCollection of chemical quality and sediment samples on a daily busis was begun on January 15, 1959. Temperature observations are made on samples voidected for chemical analyses only. The observar collects sediment samples twice daily from cable car at gating station during normal flow periods and at more frequent intervals during periods of high flow. All samples are depth intervated and taken with either a D-40 or DH-48 sampler. (Estimate percentage of reported data collected by such.) Three samples were for each set, at 1/3, 1/2, and 2/3 of distance across the stream. An integrated samples were obtained during several intervals in becceion about 0830 daily. No samples were obtained during several intervals in becember and January. The collection of samples and records maintained by observer, John H. Jones, is satisfactory.	year are equivalent to a 1.1 adjuated complete analysia. MeanrikaSediamit records at this station are considered fair during the year except during the period of December and January when many samples were missing and it was necessary to estimate sediamit dischring those 2 worths are considered poor. At times the observer was unable to make more than one trip a day though his interpolation. The records during those 2 months are considered poor. At times the observer was unable to make more than one trip a day though his interpolations were to obtain two sets of samples. This station is on a requisted stream belicons were to obtain two fact and all sediment originated downstream from this reservoir. The channel is sandy, but banks are fairly stable. No diversions or tributary inflow for several wides uptream from sampling site. The chemical quality was quite uniform throughout the year because of leveling effects of the reservoir.
EquipmentU.S. D-43 sediment sampler installed in metal shelter on bridge 100 yards upptream from gaging station. A cableway is available for sediment- discharge measurements at normal gaging section one quarter of a mile upstream from water-staga recorder.	Published and unpublished data for the 1960 water year are on file at the Albuquerque office.
Special sediment memourgements,Sixteen sediment discharge measurements vere made during the year. Each measurement constituted a cross section, consisting of deph-integrated samples from five to seven verticals. No coefficients were applied as the three regular samples showed virtually no change with any discharge measurement.	John Doe December 12, 1960
Sedimmer-discharge computationsThe mean daily concentrations obtained at the station were plotted on a print of the gage-height chart and a smooth concen- tration curve dram on a basin of the points. Daily mean concentrations were computed from the average of the cross-meticion sample. Daily sediment discharge was computed by multipying product of adjusted daily mean discharge and aman computed by multipying product of adjusted daily mean discharge and aman concentration by factor 0.0027. On 7 days of rapidly changing water discharge and concentration the graphs were subdivided and total sediment discharge for the day was computed by summing the sediment discharge for the day.	
<u>Size analysis</u> Thirty-three particle-size analyses were made on samples taken at this station during the year, of which 25 were tabulated for publication. The sample for April 10 was analysed fits trans in different ways and at different concentrations. The botcon withdrawal results were unsatisfactory and only one decontation was listed for this date. Three bed material samples were taken August 12, September 10, and September 20. Analyses were made, but the data were not listed.	

Page 1 of 2

Page 2 of 2

Figure 41.---Station analysis (chemical quality and sediment) showing format and content.

SACRAMENTO RIVER BASIN

11-4525.00 CACHE CREEK AT YOLO, CALIF.

STATION ANALYSIS (Sediment) 1959 Water Year

- Equipment.--A U.S. D-49 sediment sampler is stored at the station and used from the cable provided for stream measuring. A Stevens A-35 water-stage recorder is located on the left bank upstream from U.S. Highway 99W bridge.
- Sampling frequency.--Samples were collected once a day during low flow and two or more times a day on most days of discharge. All samples were depth integrated and were collected with a D-49 or USDH-48 sampler. The collection of samples and records maintained by the observer, Herman Gonzales, was fair to poor. Water-temperature observations were generally made at the time sediment samples were collected.
- <u>Concentration graph.</u>--Most of the suspended-sediment samples consisted of three bottles, each taken at a separate vertical in the stream cross section. The samples were analyzed separately to facilitate comparison. If any one bottle seemed to be considerably different, more weight was given to the ones that seemed more reasonable. In most cases the individual bottles agreed fairly well with each other. All concentrations were plotted on a print of the gage-height graph, and a smooth graph was drawn on the basis of the plotted concentrations. The two major rises were sampled often enough to define adequately the concentration curve.
- Suspended-sediment discharge computation.--The suspended-sediment discharge, in tons per day, was computed from the daily mean water discharge as determined at the stream-gaging station and the daily mean sediment concentration as determined at the sediment-measuring station. On days of rapidly changing flow conditions the graphs were subdivided and the total suspended-sediment discharge for the day was computed by averaging the suspended-sediment discharge for appropriate small intervals of the day. Streamflow data were furnished by the Sacramento Surface Water area office.
- Particle size.--Twelve samples were analyzed for particle size resulting in 12 observations of particle-size gradation in the sand, silt, and clay range. The samples ranged in concentration from 12 to 7,970 ppm and water discharge from 122 to 14,000 cfs. Average results of the analyses show 16 percent of the material transported to be coarser than 62 microns and 41 percent to be finer than 4 microns. All particle-size analyses in the silt-clay range were made using distilled water, with a dispersing agent added, as the settling medium.
- Remarks.--Suspended-sediment records for the year are considered fair. Water discharge records for the year are considered good. The streambed at the station is composed of sand, gravel, and cobblestones. The stream is leveed and has stable banks. Station is located on an intermittent stream.

Harry Mudboy July 25, 1960

Figure 42.-Station analysis (sediment) showing format and content.

SACRAMENTO RIVER BASIN

11-3820. THOMES CREEK AT PASKENTA, CALIF.

LOCATION (revised).--Lat 39°52'57", long 122°33'03", in SW1NW1 sec.4, T.23 N., R.6 W., Tehama County, at gaging station 0.25 mile uspstream from Digger Creek, and 0.3 mile upstream from highway bridge at Paskenta. DRAINAGE AREA. -- 194 sq mi.

RECORDS AVAILABLE. --Chemical analyses: October 1958 to September 1967. Water temperatures: October 1961 to September 1967.

Sediment records: October 1962 to September 1967. EXTREMES, 1966-67.--Water temperatures: Maximum, 94°F Aug. 18, 23; minimum, freezing point Dec. 27, 28, Jan. 6. Sediment concentrations: Maximum daily, 16,100 ppm Jan. 29; minimum daily, 1 ppm on several days.

CHEMICAL ANALYSES IN MILLIGRAMS PER LITER, WATER YEAR OCTOBER 1967 TO SEPTEMBER 1968

DATE	DIS- Charge (CFS)	SILICA (SI02)	D1S- SCLVED IPON (FE)	CAL- CIUM (CA)	MAG- NE- SIUM (MG)	SODIUM (NA)	PO- TAS- SIUM (K)	BICAR- BONATE (HC93)	CAR- Bonate {(03)	SULFATE (SU4)	CHLU- RIGE (CL)
ост.									_		
28	2.2	19	.01	20	1.3	6.5	1.1	74	0	5.0	2.7
MAR. 15		18				<i>.</i> .	-		•		
APF.	3.1	18	.00	19	1.5	6.2	•7	70	n	7.0	1.9
16 JUNE	7.2	18		18	1.5	6.1	.8	65	0	7.0	1.9
20 JUL Y	8.8	2C		11	1.0	4.8	.5	47	ç	4.0	1.2
23 AUG.	2.4	3.4		1.3	.1	1.1	• 2	5	C	2.0	1.0
15 SEPT.	3.1										
12	1.2	21		18	1.5	6.8	۹,	72	e	7.0	1.3

SACRAMENTO RIVER BASIN

11-3820. THOMES CREEK AT PASKENTA, CALIF .-- Continued

,

EXTREMES, 1966-67--Continued

EXTREMES, 1966-67--Continued Sediment loads: Maximum daily, 273,000 tons Jan. 29; minimum daily, less than 0.05 ton on many days. EXTREMES, 1961-67.--Water temperatures: Maximum, 94°F Aug. 18, 23, 1967; minimum, freezing point on several days in December and January of most years. Sediment concentrations (1962-67): Maximum daily, 57,900 ppm Dec. 22, 1964; minimum daily, no flow Oct. 4, 1964. Sediment loads (1962-67): Maximum daily, 5,070,000 tons Dec. 22, 1964; minimum daily, 0 ton Oct. 4, 1964. REMARKS.--Clock stopped Apr. 14-12; temperature range, 38°F to 53°F.

CHEMICAL ANALYSES IN MILLIGRAMS PER LITER, WATER YEAR OCTOPER 1967 TO SEPTEMBER 1968

DATE	FLUO- KIDE (F)	N I TRATE (NU3)	BUR CN (B)	DIS- SOLVED SOLIDS (RESI- DUE AT 180 CI	DIS- SOLVED SOLIDS (TONS PER AC-FT)	DIS- SOLVED SOLIDS (TONS PER DAY)	HARD- NESS (CA,MG)	NI)N- CAR- BONATE HARD- NESS	SODIUM AD- SORP- TION RATIO	SPECI- FIC COND- UCTANCE (MICRO- MHOS)	рн
001.											
28	• 2	•C	• 02	97	•13	.58	56	0	•4	134	7.6
MAK. 15 APK.	• 2	• 1	.00		•12	.74	54	¢	•4	137	7.9
16 JUNE	• 2	• 4	•07		•12	1.67	51	c	•4	126	7.7
2C	• 1	•1	•00		.09	1.57	32	0	•4	92	7.4
23 AUG.	•0		.00		•01	.07	4	o	•3	14	6.4
15 SEPT.		•0									
12	•1	• 5	.00		•13	.30	51	0	.4	133	7.5

Figure 43.---Heading and chemical-quality tabulation in annual water-quality-data report. (Heading and tabulation are on two pages that are published face to face or back to back.) SACRAMENTO RIVER BASIN

11-3820. THOMES CREEK AT PASKENTA, CALIF.

LOCATION (rovined) .-- Lat 39°52'57", long 122°33'03", in SW}NW sec.4, T.23 N., R.6 W., Tehama County, at gaging

station 0.25 mile upstream from Digger Creek, and 0.3 mile upstream from highway bridge at Paskenta. DRAINAGE AREA. -- 194 sq m1.

RECORDS AVAILABLE, --- Chemical analyses: October 1958 to September 1967

1968 Water temperatures: October 1961 to September 1967

1968 Sediment records: October 1962 to September 1967.

1967-66 EXTREMES, 1968-67.--Water temperatures: Maximum, 6447 Aug. 15, 25; minimum, freezing point, December and January. 10,700 mg/2 14 Sediment concentrations: Maximum daily, 15,700 ppm Jan. 25; minimum daily, 1 ppm on several days. 149,000 149,000 Sediment loads: Maximum daily, 272,000 tons Jan. 29; minimum daily, less them 0.01 ton on many days. 534°C EXTREMES, 1961-67.--Water temperatures: Maximum, 94°F Aug. 18, 23, 1967; minimum, freezing point on several days in December and January of most years.

in December and January of most years. 60,200 mg/L Sediment concentrations (1962-57): Maximum daily, 57,900 ppm Dec. 22, 1964; minimum daily, no flow Oct. 4, 1964. Sediment loads (1962-57): Maximum daily, 5,070,000 tons Dec. 22, 1964; minimum daily, 0 ton Oct. 4, 1964. REMARKS .-- Clock stopped Apr. 14-12; tempocature range, 3897 to 53 %. Chew cal- quality records furnished by California d'epartine, 1 of Water Poscine. Where no maximum or minimum is showing ten perature is oncé-daily obscivition.

(make updating in red)

Figure 44.—Updated station heading.

location during the water year, water temperatures, sampling procedures, and reference to discharge records if (1) sampling site and gaging stations were at different locations or (2) discharge records were furnished by other agencies. Reference may also be made to upstream factors affecting water and sediment movement and to periods of affected or estimated record.

Periodic Observations

Sediment samples are obtained at some stations on a periodic or infrequent basis. and at a few stations, samples are obtained monthly or more frequently during the periods of storm runoff. Data obtained on an infrequent basis that cannot be used to compute daily sediment-discharge records as described in the procedure used for daily stations are tabulated on form 9-1539D as illustrated in figure 45.

Periodic observations of suspended-sediment discharge are published as daily suspended sediment, illustrated in figure 30, or as instantaneous suspended sediment, illustrated in figure 31. They also may be combined with particle-size distribution and published as illustrated in figure 34.

The data for periodic observations may be tabulated on form 9-265b in the same format as that for particle-size distribution shown in figure 2 and described in the section on "Tabulation of Size Data."

Footnotes used for periodic tabulations are standard and are the same as those used for daily tabulations. (See section on "Footnotes.") Periodic or miscellaneous concentration and particle-size data are tabulated on form 9-1593D as described in the section on "Tabulation of Size Data" (fig. 2).

Checklist for periodic records

A checklist for tabulation of periodic samples follows. The editorial rules and computation procedures are the same as those described for the similar step for daily stations.

- 1. Size analysis: Compute and review for proper procedures and accuracy.
- 2. Concentration notes: Compute and evaluate for validity and sampling technique.

Form
9-1539D
(10-69)

UNITED STATES DEPARTMENT OF THE INTERIOR - GEOLOGICAL SURVEY - WATER RESOURCES DIVISION Coding Form for Instantaneous Suppended Sediment and Particle Size

Type Station ident number 1 0000000000000000 Stream and <u>Sample Creek near Somewhere</u> , 1.5.A. Methods of analysis (74-77): B, bottom withdrawal tube: C, chemically dispersed: D, decantation: N, in native water: P, pipet: S, sieve, W, in distilled water: V, visual accumulation tube.													
Please do not punch the decimal point unless data appears in the field.													
lime	Time Discharge Fardere ster, perceat mez (a animiteters) dan size indicated												Method of
Yr Mo Day (24 hour) (C)	(cfs)	(mg/1)	0 002	0 004 0	008 0.016	0 031	0 062	0 125	0 250	0 500	1 200	2 300	analysis
1 11 12 13 14 15 16 17 14 14 22 21 22 23	26 25 26 27 28 29 30 31 32 33	34 35 36 37 38 34 40	61 62 63 66	45 45 47 4	8 49 50 51 5:	53 54 55	56 57 58	59 60 61	62 63 64	65 66 67	68 69 70	<u>n n n</u>	14 75 76 77
691011152014.5	1.1.1.1.1.1.9.1.1			. Li	Lin		7.01	74.	7.7.	96	100		S
				-					•				
1,21,11,400 4.0	1 1 1/14211	1.1.1.0											
					<u> </u>	*							
7.00/15/0005	1.1.1.08.1.1			· · · · · ·	· · · · ·								
Telev Tist Koop				<u></u>	┹╌┫╌╴┫╺╌┙┩╍╸						أساسط		أحصاف والمساقية
	· · · · · · · · · · · · · · · · · · ·	· · · · · · · · · · · · · · · · · · ·				T							
02191440 7:0	1 1 2 2 4 1	44			<u>. L. L</u>	لمبيل	لتب				لىب	نحب	أحصحما
	· · · · · · · · · · · · · · · · · · ·	·	·					,					
1 93,30 11,05 6.0		4/2	23 2	8 3.9	1 48	63	7.1	8.9	96	9.9	100	سیاسیا	VPWC
05/B130011.0	595.			[.			60	7.4.	9.1.	100		1 1	5
		· · · · · · · · · · · · · · · · · · ·			<u> </u>	4							*****
07/6/32029.0	, , ,2,0,.				1								
<u> </u>				L. L. L. L.	<u> </u>		الم الم الم				لىبى	ليسابعه المساسم	لمسلسا
[<u>F</u> <u>F</u>	r	· · · · · · · · · · · · · · · · · · ·	· 7		·	· · · · · · · · · · · · · · · · · · ·							
	<u></u>			чĻт			أحياب		سلبلہ			أساساسا	
		·	···· ·			,							
						L						لللله	
		M							1.1		1.1	11	
	Co	P		Date 29	19/20		ked by	RD	B			101	12/70
	Compiled by	* !				Chec	ked by				- Del	ie ~~~	
	Punched by			Date		Venf	led by	<u> </u>			_ Det	te	

Figure 45.—Tabulation of periodic sediment data.

- 3. Particle-size analyses: Compile in chronological order on form similar to that illustrated in figure 45.
- 4. Water discharge:
 - (a) Report water discharge for time each sample was collected if instantaneous values are published (figs. 31-35).
 - (b) Report daily mean discharge if daily mean concentrations are published (figs. 29, 30).
- 5. Sediment discharge: Compute sediment discharge, in tons per day. Sediment discharge is equal to water discharge, in cubic feet per second, \times sediment concentration, in milligrams per liter, \times 0.0027. Sediment discharge is not computed if data are punched for computer processing.

- 6. Review for editorial and technical correctness.
- 7. Station description heading is not prepared for periodic records.

Combined periodic and seasonal observations

A sediment record for a station where sediment samples were collected on a periodic basis for part of the year and on a daily or more frequent basis for the remainder of the year may be compiled in two parts. This type of record may occur because of a change in station operation or because of seasonal effect on the stream discharge and sediment concentration. The periodic part of the record consists of a tabulation of samples which are collected infrequently during routine visits and which do not provide sufficient information for computation of daily sediment discharge. These samples are combined with particle-size data on the form illustrated in figure 45.

The second part of the record consists of daily discharges computed by the same procedures described previously, and the sediment-discharge worksheet (fig. 38) is completed for as many complete, consecutive months as data are available. A station description and an analysis are prepared as described for computation of records for daily stations.

Transmittal of Completed Data

The completed compilation of water-quality data includes the following items:

- 1. Station heading (fig. 44).
- 2. Chemical analyses (fig. 43).
- 3. Temperature tabulation (figs. 3, 4).
- 4. Sediment discharge:

Daily records (fig. 38).

Periodic records (fig. 45).

- 5. Particle-size distribution of suspended sediment (fig. 2).
- 6. Particle-size distribution of bed material (fig. 35).

Each originating office should forward the records for which they are responsible through appropriate channels to be reviewed, consolidated, and forwarded to the records center for processing. Each record submitted for publication should be designated as to publication format desired (figs. 29-35).

Selected References

- Buchanan, T. J., and Somers, W. P., 1965, Discharge measurements at gaging stations: U.S. Geol. Survey Surface Water Techniques, book 1, chap. 11, 67 p.
- Carter, R. W., and Davidian, Jacob, 1965, Discharge ratings at gaging stations: U.S. Geol. Survey Surface Water Techniques, book 1, chap. 12, 36 p.
 - ----- 1968, General procedure for gaging streams:

U.S. Geol. Survey Techniques of Water-Resources Inv., book 3, chap. A6, 13 p. file report, 170 p.

Colby, B. R., 1956a, Relationship of sediment discharge to streamflow: U.S. Geol. Survey open-

- 1957, Relationship of unmeasured sediment discharge to mean velocity: Am. Geophys. Union Trans., v. 38, no. 5, p. 708-717.
- ------ 1963, Fluvial sediments-A summary of source, transportation, deposition, and measurement of sediment discharge: U.S. Geol. Survey Bull. 1181-A, 47 p.
- 1964a, Discharge of sands and mean-velocity relationships in sand-bed streams: U.S. Geol. Survey Prof. Paper 462-A, 47 p.
- ------- 1964b, Scour and fill in sand-bed streams: U.S. Geol. Survey Prof. Paper 462-D, 32 p.
- Colby, B. R., and Hembree, C. H., 1955, Computations of total sediment discharge, Niobrara River near Cody, Nebraska: U.S. Geol. Survey Water-Supply Paper 1357, 187 p.
- Colby, B. R., Hembree, C. H., and Rainwater, F. H., 1956, Sedimentation and chemical quality of surface waters in the Wind River Basin, Wyoming: U.S. Geol. Survey Water-Supply Paper 1373, 336 p.
- Colby, B. R., and Hubbell, D. W., 1961, Simplified methods for computing total sediment discharge with the modified Einstein procedure: U.S. Geol. Survey Water-Supply Paper 1593, 17 p.
- Colby, B. R., and Scott, C. H., 1965, Effects of water temperature on the discharge of bed material: U.S. Geol. Survey Prof. Paper 462-G, 25 p.
- Corbett, D. M., and others, 1943, Stream-gaging procedure, a manual describing methods and practices of the Geological Survey: U.S. Geol. Survey Water-Supply Paper 888, 245 p.
- Dawdy, D. R., 1961, Depth-discharge relations of alluvial streams—discontinuous rating curves: U.S. Geol. Survey Water-Supply Paper 1498-C, 16 p.
- Einstein, H. A., 1950, The bedload function for sediment transportation in open channel flows: U.S. Dept. Agriculture, Soil Conserv. Service Tech. Bull. 1026, 70 p.
- Einstein, H. A., Anderson, A. G., and Johnson, J. W., 1940, A distinction between bedload and suspended load in natural streams: Am. Geophys. Union Trans., v. 21, pt. 2, p. 628-633.
- Fahnestock, R. K., 1963, Morphology and hydrology of a glacial stream—White River, Mount Rainier, Washington: U.S. Geol. Survey Prof. Paper 422-A, 70 p.

- Federal Inter-Agency River Basin Committee, 1948, Proceedings of the Federal Inter-Agency Sedimentation Conference, Denver, May 6-8, 1947: Washington, U.S. Bur. Reclamation, 314 p.
- Gilbert, G. K., 1914, The transportation of debris by running water: U.S. Geol. Survey Prof. Paper 86, 263 p.
- Gottschalk, L. C., 1952, Measurement of sediment in small reservoirs: Am. Soc. Civil Engineers Trans., v. 117, p. 59-71.
- Guy, H. P., 1957, The trend of suspended-sediment discharge of the Brandywine Creek at Wilmington, Delaware, 1947-55: U.S. Geol. Survey openfile report, 55 p.

- 1969, Laboratory theory and methods for sediment analysis: U.S. Geol. Survey Techniques Water-Resources Inv., book 5, chap. Cl, 58 p.
- Guy, H. P., and Norman, V. W., 1970, Field methods for fluvial sediment measurements: U.S. Geol. Survey Techniques Water-Resources Inv., book 3, chap. C2, 59 p.
- Guy, H. P., and Simons, D. B., 1964, Dissimilarity between spatial and velocity-weighted sediment concentrations: U.S. Geol. Survey Prof. Paper 475-D, p. D134-D137.
- Heidel, S. G., 1956, The progressive lag of sediment concentration with flood waves: Am. Geophys. Union Trans., v. 37, no. 1, p. 56-66.
- Hubbell, D. W., 1960, Investigations of some sedimentation characteristics of sand-bed streams, Progress report 2: U.S. Geol. Survey open-file report, 78 p.
- 1964, Apparatus and techniques for measuring bedload: U.S. Geol. Survey Water-Supply Paper 1748, 74 p.
- Langbein, W. B., and Schumm, S. A., 1958, Yield of sediment in relation to mean annual precipitation: Am. Geophys. Union Trans., v. 39, no. 6, p. 1076-1084.
- Leopold, L. B., and Maddock, Thomas, 1953, The hydraulic geometry of stream channels and some physiographic implications: U.S. Geol. Survey Prof. Paper 252, 57 p.
- Leopold, L. B., Wilman, M. G., and Miller, J. P., 1964, Fluvial processes in geomorphology: San Francisco and London, Freeman, 522 p.
- Meyer-Peter, E., and Muller, R., 1948, Formulas for bedload transport: Int. Assoc. for Hydraulic Structures Research, 2d mtg., Stockholm, p. 39-64.

- Miller, C. R., 1951, Analysis of flow-duration, sediment rating curve method of computing sediment yield: Denver, U.S. Bureau of Reclamation, 15 p.
- Ritter, J. R., and Helley, E. J., 1969, Optical method for determining particle sizes of coarse sediment: U.S. Geol. Survey Techniques Water-Resources Inv., book 5, chap. C3, 33 p.
- Schroeder, K. B., and Hembree, C. H., 1956, Application of the modified Einstein procedure for computation of total sediment load: Am. Geophys. Union Trans., v. 37, no. 2, p. 197-212.
- Schumm, S. A., 1961, Effect of sediment characteristics on erosion and deposition in ephemeral-stream channels: U.S. Geol. Survey Prof. Paper 352-C, p. 31-70.
- Shulits, Sam, 1935, The Schoklitsch bedload formula: Engineering, v. 139, p. 644-646, 687.
- Simons, D. B., Richardson, E. V., and Haushild, W. L., 1963, Some effects of fine sediment on flow phenomena: U.S. Geol. Survey Water-Supply Paper 1498-G, 47 p.
- Twenhofel, W. H., 1950, Principles of sedimentation [2d ed.]: New York, McGraw-Hill Book Co., 610 p.
- U.S. Bureau of Reclamation, 1960, Investigation of Meyer-Peter and Muller bedload formulas: Proj. Inv. Div., Denver, 22 p.
- U.S. Inter-Agency Committee on Water Resources, Subcommittee on Sedimentation, 1940, Field practice and equipment used in sampling suspended sediment, Report 1 of A study of methods used in measurement and analysis of sediment loads in streams: 175 p.
 - —— 1940, Equipment used for sampling bedload and bed material, Report 2 of A study of methods used in measurement and analysis of sediment loads in streams: 57 p.
 - ——— 1941, Analytical study of methods of sampling suspended sediment, Report 3 of A study of methods used in measurement and analysis of sediment loads in streams: 82 p.
 - 1941, Methods of analyzing sediment samples, Report 4 of A study of methods used in measurement and analysis of sediment loads in streams: 203 p.

 - ------ 1943, A study of new methods for size analysis of suspended-sediment samples, Report 7 of A study of methods used in measurement and analysis of sediment loads in streams: 102 p.
 - 1943, Density of sediments deposited in reservoirs, Report 9 of A study of methods used in measurement and analysis of sediment loads in streams: 60 p.

- 1952, The design of improved types of suspended-sediment samplers, Report 6 of A study of methods used in measurement and analysis of sediment loads in streams: 103 p.
- 1953, Accuracy of sediment size analyses made by the bottom-withdrawal-tube method, Report 10 of A study of methods used in measurement and analysis of sediment loads in streams: 115 p.
- <u>1957</u>, The development and calibration of the visual-accumulation tube, Report 11 of A study of methods used in measurement and analysis of sediment loads in streams: 109 p.
- 1957, Some fundamentals of particle-size analysis, Report 12 of A study of methods used

in measurement and analysis of sediment loads in streams: 55 p.

- 1961, Single-stage sampler for suspended sediment, Report 13 of A study of methods used in measurement and analysis of sediment loads in streams: 105 p.
- 1963, Determination of fluvial sediment discharge, Report 14 of A study of methods used in measurement and analysis of sediment loads in streams: 151 p.
- Vetter, C. P., 1953, Sediment problems in Lake Mead and downstream on the Colorado River: Am. Geophys. Union Trans., v. 34, no. 2, p. 249-256.
- Waananen, A. O., Harris, D. D., and Williams, R. C., 1971, Floods of December 1964 and January 1965 in the Far Western States: Part 1, Description: U.S. Geol. Survey Water-Supply Paper 1866-A, 265 p.