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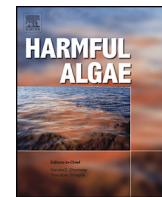
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Review

How rising CO₂ and global warming may stimulate harmful cyanobacterial blooms[☆]



Petra M. Visser ^{a,*}, Jolanda M.H. Verspagen ^a, Giovanni Sandrini ^a, Lucas J. Stal ^{a,b}, Hans C.P. Matthijs ^a, Timothy W. Davis ^c, Hans W. Paerl ^d, Jef Huisman ^a

^a Department of Aquatic Microbiology, Institute for Biodiversity and Ecosystem Dynamics, University of Amsterdam, P.O. Box 94248, 1090 GE Amsterdam, The Netherlands

^b Department of Marine Microbiology, Royal Netherlands Institute for Sea Research (NIOZ), P.O. Box 140, 4400 AC Yerseke, The Netherlands

^c NOAA Great Lakes Environmental Research Laboratory, Ann Arbor, MI 48108, USA

^d Institute of Marine Sciences, University of North Carolina at Chapel Hill, 3431 Arendell Street, Morehead City, NC 28557, USA

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ABSTRACT

Climate change is likely to stimulate the development of harmful cyanobacterial blooms in eutrophic waters, with negative consequences for water quality of many lakes, reservoirs and brackish ecosystems across the globe. In addition to effects of temperature and eutrophication, recent research has shed new light on the possible implications of rising atmospheric CO₂ concentrations. Depletion of dissolved CO₂ by dense cyanobacterial blooms creates a concentration gradient across the air-water interface. A steeper gradient at elevated atmospheric CO₂ concentrations will lead to a greater influx of CO₂, which can be intercepted by surface-dwelling blooms, thus intensifying cyanobacterial blooms in eutrophic waters. Bloom-forming cyanobacteria display an unexpected diversity in CO₂ responses, because different strains combine their uptake systems for CO₂ and bicarbonate in different ways. The genetic composition of cyanobacterial blooms may therefore shift. In particular, strains with high-flux carbon uptake systems may benefit from the anticipated rise in inorganic carbon availability. Increasing temperatures also stimulate cyanobacterial growth. Many bloom-forming cyanobacteria and also green algae have temperature optima above 25 °C, often exceeding the temperature optima of diatoms and dinoflagellates. Analysis of published data suggests that the temperature dependence of the growth rate of cyanobacteria exceeds that of green algae. Indirect effects of elevated temperature, like an earlier onset and longer duration of thermal stratification, may also shift the competitive balance in favor of buoyant cyanobacteria while eukaryotic algae are impaired by higher sedimentation losses. Furthermore, cyanobacteria differ from eukaryotic algae in that they can fix dinitrogen, and new insights show that the nitrogen-fixation activity of heterocystous cyanobacteria can be strongly stimulated at elevated temperatures. Models and lake studies indicate that the response of cyanobacterial growth to rising CO₂ concentrations and elevated temperatures can be suppressed by nutrient limitation. The greatest response of cyanobacterial blooms to climate change is therefore expected to occur in eutrophic and hypertrophic lakes.

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* Corresponding author. Tel.: +31205257073; fax: +31205257832.

E-mail address: P.M.Visser@uva.nl (P.M. Visser).

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1. Introduction

It is well-established that in addition to anthropogenic nutrient enrichment, changes in the Earth's climate, specifically rising temperatures and altered hydrologic patterns, strongly influence the frequency, intensity, and duration of harmful cyanobacterial blooms ([Robarts and Zohary, 1987](#); [Peeters et al., 2007](#); [Suikkanen et al., 2007](#); [Wiedner et al., 2007](#); [Jöhnk et al., 2008](#); [Paerl and Huisman, 2008, 2009](#); [Wagner and Adrian, 2009](#); [O'Neil et al., 2012](#); [Paerl and Paul, 2012](#)). An expansion of cyanobacterial blooms is of great societal concern, because harmful cyanobacteria can impair safe drinking, irrigation, fishing and recreational waters that are critical for the growing global human population.

There is convincing evidence that a key driver of climate change is the concentration of atmospheric carbon dioxide (CO₂), which has been shown to modulate the Earth's surface and water temperatures via the 'greenhouse effect' ([IPCC, 2012](#)). Furthermore, long-term records of atmospheric CO₂ in ice cores and the atmosphere (e.g., at Mauna Loa, Hawaii) have shown that there is a well-defined parallel between increasing CO₂ concentrations and the rise of man-made fossil fuel combustion ([Tans et al., 1990](#)).

The relationships between rising atmospheric CO₂ levels, global warming and declining water quality are controlled through complex interactions with altered evaporation and rainfall patterns, changing hydrological flows and shifts in chemical and biological processes, all of which interact in non-linear ways ([Paerl and Paul, 2012](#)). This creates an enormous challenge in predicting the quantitative and qualitative ramifications for the many types of water bodies that are likely to be impacted. Furthermore, the transport and delivery of nutrients that are critical for development, proliferation and maintenance of cyanobacterial blooms are strongly influenced by climate-driven changes in precipitation patterns and biogeochemical processes ([Michalak et al., 2013](#)). All of these factors ultimately control planktonic communities, including cyanobacterial blooms ([Mitrovic et al., 2003](#); [Elliott, 2010](#); [Hall et al., 2013](#); [Michalak et al., 2013](#)).

In addition to its influence on global warming, rising atmospheric CO₂ levels may stimulate the proliferation of surface-dwelling cyanobacteria by providing them preferential access to a vast and rising pool of atmospheric CO₂ ([Paerl and Ustach, 1982](#); [Ibelings and Maberly, 1998](#); [Verspagen et al., 2014b](#)). An increase in atmospheric CO₂ increases its dissolution in water. Enhanced dissolution of CO₂ lowers pH, causing a slow acidification of the oceans ([Orr et al., 2005](#); [Doney et al., 2009](#)). In freshwaters, the impact of rising atmospheric CO₂ appears more complex than in most marine ecosystems. Freshwater systems range widely in pH and alkalinity ([Lazzarino et al., 2009](#); [Balmer and Downing, 2011](#)), which affects the speciation of inorganic carbon. Many freshwater ecosystems receive large amounts of organic carbon from terrestrial systems, which may result in CO₂ supersaturation, i.e., dissolved CO₂ concentrations that greatly exceed equilibrium with the atmosphere ([Cole et al., 1994](#); [Sobek](#)

[et al., 2005](#)). Conversely, in other lakes, CO₂ concentrations are strongly depleted as a consequence of the photosynthetic activity of dense phytoplankton blooms ([Talling, 1976](#); [Balmer and Downing, 2011](#); [Verspagen et al., 2014b](#)). Similar to the depletion of other resources, depletion of inorganic carbon (C_i) can limit growth ([Hein, 1997](#)), particularly in dense surface blooms of cyanobacteria ([Ibelings and Maberly, 1998](#)). Hence, the natural range of variation in CO₂ availability is much larger in lakes than in marine or terrestrial ecosystems, and bloom-forming cyanobacteria must cope with this variability.

This review focuses on the current state of knowledge on effects of climate change on harmful cyanobacteria. Although many reviews have already addressed this topic (e.g., [Paerl and Huisman, 2009](#); [Carey et al., 2012](#); [O'Neil et al., 2012](#)), most reviews focused on the direct or indirect effects of increased temperature, often in combination with accelerating eutrophication. In this review, effects of rising CO₂ concentrations on cyanobacteria are also addressed. The mechanistic underpinnings supporting cyanobacterial expansion in an atmospherically-CO₂ enriched, warmer, and nutrient-enriched world will be explored.

Physiological traits vary among species and strains and may direct the response of cyanobacterial species to a changing climate. First, an overview of these responses to elevated CO₂ concentrations will be provided, with special emphasis on CO₂-concentrating mechanisms (CCMs). Then the focus will be on direct and indirect temperature effects on cyanobacterial growth and competition, followed by a further exploration of interactive effects of climate change with nutrient availability. Key questions to be addressed are, for instance, whether global change is likely to lead to a proliferation of cyanobacteria at the expense of eukaryotic phytoplankton species, and whether the composition of cyanobacterial blooms may change.

2. Response to rising CO₂

2.1. Does rising CO₂ intensify bloom development?

Rising atmospheric CO₂ levels are often thought to have only minor impacts on bloom development in freshwater ecosystems. This assumption is based on two common misconceptions. It is often argued (1) that the CO₂ concentrations in freshwater lakes are sufficiently high to cover the carbon demands of phytoplankton populations, because many lakes are "supersaturated" with CO₂ ([Cole et al., 1994](#); [Sobek et al., 2005](#); [Jansson et al., 2012](#)) and (2) that changes in CO₂ availability have little effect on bloom development, because most cyanobacteria can also utilize bicarbonate as C source.

Concerning the first misconception, it is true that the pCO₂ in many lakes worldwide is well above atmospheric equilibrium (i.e., supersaturated; [Cole et al., 1994](#)). Most carbon input in lakes originates from terrestrial primary production in the surrounding watershed and not from atmospheric CO₂ ([Cole and Caraco, 2001](#);

Pacala et al., 2001; Richey et al., 2002; Maberly et al., 2013), which is subsequently mineralized, causing pCO₂ levels that commonly exceed 1500 ppm. Even in these “supersaturated waters”, however, the actual concentration of dissolved CO₂ (CO₂(aq)) is still quite low, and cyanobacterial blooms can easily turn a supersaturated lake into an undersaturated lake (Ibelings and Maberly, 1998; Verspagen et al., 2014b). For instance, consider a supersaturated lake with a pCO₂ of 1500 ppm. According to Henry's law, assuming a solubility constant of K_H = 0.034 mol L⁻¹ atm⁻¹, the CO₂(aq) concentration in this lake would be only ~50 μmol L⁻¹. This concentration is certainly not enough to cover the photosynthetic carbon demand of a dense cyanobacterial bloom. The photosynthetic activity of dense blooms can be as high as 12.5–50 μmol C L⁻¹ h⁻¹ (Hein, 1997), depleting the CO₂(aq) concentration in this lake within a few hours (Talling, 1976; Maberly, 1996). In some lakes, the CO₂(aq) concentration can even be drawn down to less than 0.1 μmol L⁻¹, corresponding to a pCO₂ of only a few ppm (Lazzarino et al., 2009; Balmer and Downing, 2011).

Data from Lake Volkerak, a large eutrophic lake in The Netherlands, are provided in Fig. 1 (Verspagen et al., 2006; 2014b). In this figure, the CO₂(aq) concentration that would be predicted from equilibrium with the atmosphere (i.e., [CO₂^{*}]) has also been indicated. This expected equilibrium CO₂(aq) concentration shows some seasonal variations, as the solubility of CO₂ in water is temperature dependent. Seasonal variation of the measured CO₂(aq) concentration in Lake Volkerak is much larger, however, because biological consumption and production of CO₂ act at a much faster rate than the equilibration of CO₂ between water and atmosphere. In winter and spring, the measured CO₂(aq) concentration in Lake Volkerak largely exceeds the CO₂(aq) concentration that would be predicted from equilibrium with the atmosphere, and hence in winter and spring the lake is supersaturated with CO₂. Conversely, dense blooms of the harmful cyanobacterium *Microcystis* occur in Lake Volkerak in summer and early fall. The photosynthetic activity of these blooms depletes the CO₂(aq) concentration to 1 μmol L⁻¹ (~30 ppm), such that the lake becomes severely undersaturated with CO₂ in summer while the pH rises above 9 for several months (Fig. 1). These data illustrate that the CO₂(aq) concentration in eutrophic lakes can vary from supersaturation in winter to undersaturation in summer.

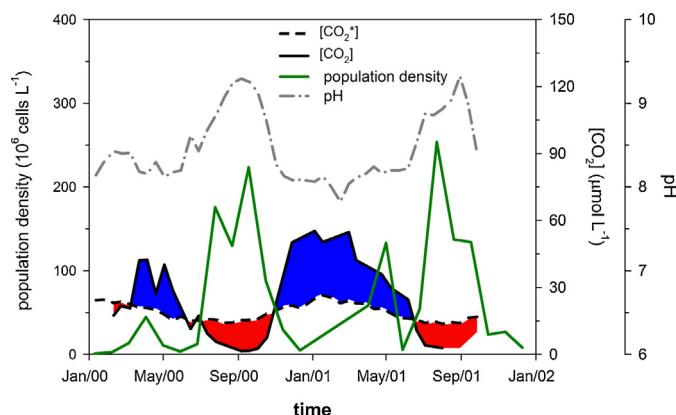


Fig. 1. Seasonal changes in phytoplankton population density (green line), dissolved CO₂ concentration ([CO₂], black solid line) and pH (grey dash-dotted line) in Lake Volkerak during two consecutive years. The black dashed line is the expected dissolved CO₂ concentration ([CO₂^{*}]) when assuming equilibrium with the atmospheric pCO₂ level. Blue shading indicates that the lake is supersaturated with CO₂, whereas red shading indicates undersaturation. In the months July–October, the cyanobacterium *Microcystis* comprised 75–98% of the phytoplankton population. Adjusted from Verspagen et al. (2014b).

The drawdown of the CO₂(aq) concentration by cyanobacterial blooms turns lakes into a sink for atmospheric CO₂ (Balmer and Downing, 2011). The CO₂ gas influx depends on the CO₂ deficit. More specifically, the CO₂ influx (g_{CO_2}) is proportional to the difference between the expected concentration of CO₂(aq) in equilibrium with the atmosphere (calculated from Henry's law) and the observed CO₂(aq) concentration (Siegenthaler and Sarmiento, 1993; Cole et al., 2010):

$$g_{\text{CO}_2} = v(K_{\text{H}}p_{\text{CO}_2} - \text{CO}_2(\text{aq})) \quad (1)$$

Here, v is the gas transfer velocity (also known as piston velocity) across the air–water interface, K_{H} is the solubility constant of CO₂ gas in water, and pCO₂ is the partial pressure of CO₂ in the atmosphere. If it is assumed that the dense cyanobacterial bloom has stripped the surface layer of CO₂(aq), this equation simplifies to $g_{\text{CO}_2} = vK_{\text{H}}p_{\text{CO}_2}$. The gas transfer velocity depends on several parameters, especially wind speed. A typical value for the gas transfer velocity of lakes is $v = 0.02 \text{ m h}^{-1}$ (Crusius and Wanninkhof, 2003; Cole et al., 2010). Assuming an atmosphere with pCO₂ = 400 ppm, the CO₂ influx during a dense cyanobacterial bloom would amount to ~7 mmol m⁻² d⁻¹. This influx is substantial and can be intercepted by the surface-dwelling cyanobacterial bloom for supporting photosynthesis (Paerl and Ustach, 1982; Ibelings and Maberly, 1998). A doubling of the atmospheric CO₂ concentration, to 800 ppm, would roughly double the CO₂ influx to ~14 mmol m⁻² d⁻¹. Moreover, this might still be an underestimate for dense cyanobacterial blooms. At pH > 9, which is typical for dense blooms, the chemical reaction of CO₂ with the abundant hydroxide ions further increases CO₂ transfer across the air–water surface by a process known as chemically enhanced diffusion (Emerson, 1975; Bade and Cole, 2006). This simple calculation shows that, in principle, an increase in atmospheric CO₂ levels may provide a sufficient influx of C to enable a substantial increase in the productivity of surface-dwelling cyanobacterial blooms.

Models and laboratory experiments have shown that rising CO₂ concentrations may indeed exacerbate cyanobacterial blooms (Schippers et al., 2004; Verspagen et al., 2014b). Verspagen et al. (2014b) performed chemostat experiments with *Microcystis* CYA140 under nutrient-saturating conditions. At a low atmospheric pCO₂ level of 200 ppm (half the current ambient pCO₂), the *Microcystis* population increased until it reached a steady state, at which it had depleted the dissolved CO₂(aq) concentration to 0.2 μmol L⁻¹ and raised the pH to 10 (Fig. 2A, C, and E). The same experiment was repeated at an elevated atmospheric pCO₂ level of 1200 ppm (three times ambient pCO₂), which resulted in a doubling of the *Microcystis* biomass, whereas the CO₂(aq) concentration was much less depleted and the pH was raised to only 8.5 (Fig. 2B, D, and F). The model predictions matched the experiments. These results demonstrate, both in theory and lab experiments, that bloom-forming cyanobacteria such as *Microcystis* can become carbon-limited, and that rising pCO₂ levels can increase cyanobacterial biomass (Verspagen et al., 2014b).

The second misconception is that changes in CO₂ availability have little effect on bloom development, because most cyanobacteria can also utilize bicarbonate. Indeed, it is true that many if not most cyanobacteria can use bicarbonate. However, whereas CO₂ passively diffuses through the cell membrane, utilization of bicarbonate requires investments in sodium-dependent and ATP-dependent bicarbonate uptake systems as well as in sodium antiporters that excrete the sodium accumulated in the cells (Price, 2011; Burnap et al., 2015; Sandrini et al., 2015b). These costs of bicarbonate utilization may have repercussions for the growth rates that can be achieved. For instance, *Synechococcus leopoliensis* grows at 80% of its maximum growth rate when bicarbonate is its

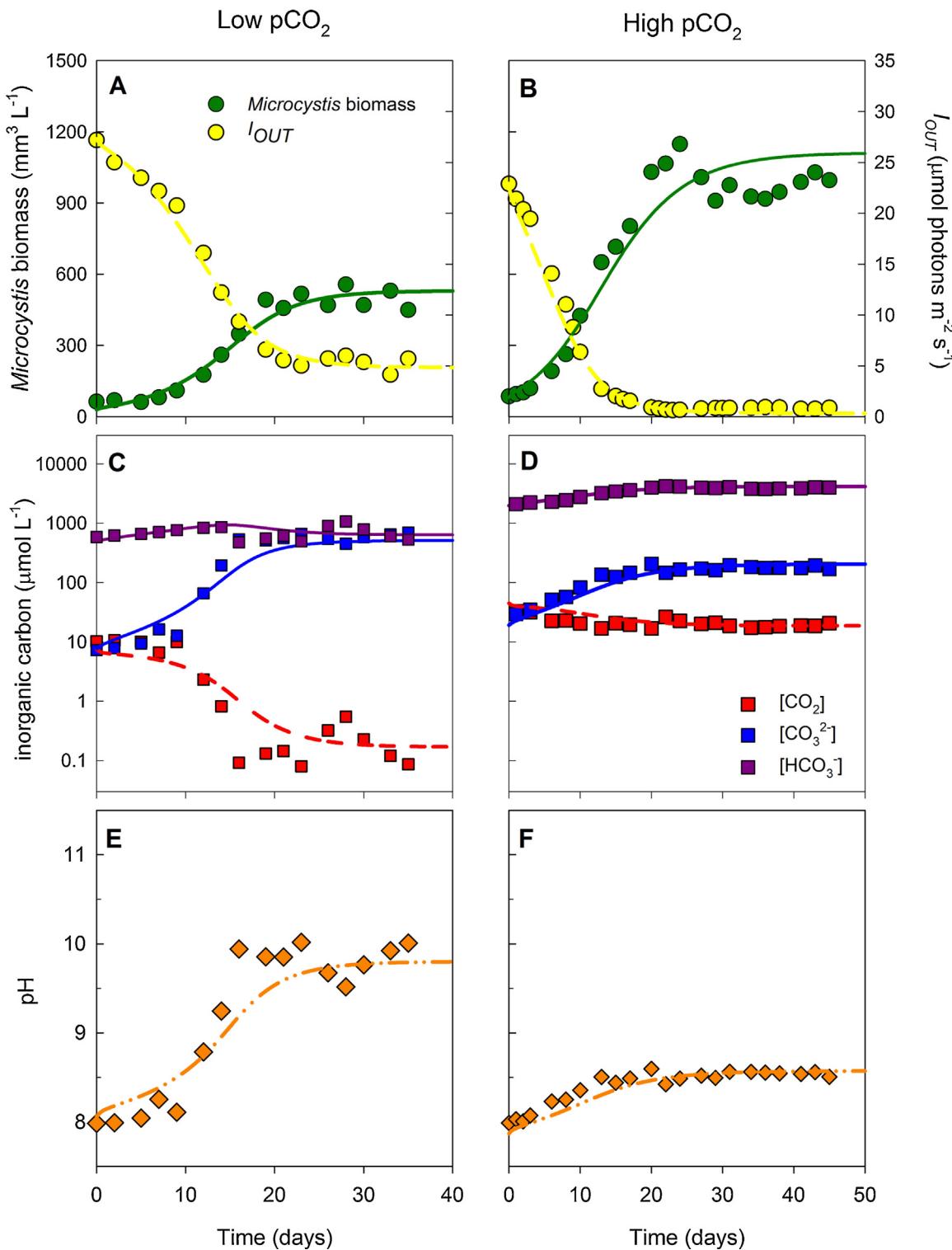


Fig. 2. Cyanobacterial growth and inorganic carbon chemistry at two different $p\text{CO}_2$ levels. Left panels: chemostat experiment with low $p\text{CO}_2$ of 200 ppm in the gas flow and 500 $\mu\text{mol L}^{-1}$ bicarbonate in the mineral medium. Right panels: chemostat experiment with high $p\text{CO}_2$ of 1200 ppm in the gas flow and 2000 $\mu\text{mol L}^{-1}$ bicarbonate in the mineral medium. Both chemostats were inoculated with *Microcystis* CYA140. (A, B) *Microcystis* biomass (expressed as biovolume) and light intensity penetrating through the chemostat (I_{OUT}). (C, D) Dissolved CO_2 , bicarbonate and carbonate concentrations. (E, F) pH. Symbols represent measurements, lines show model predictions. Adjusted from Verspagen et al. (2014b).

main carbon source (Miller et al., 1984). *Microcystis* HUB5-2-4, which lacks the high-flux bicarbonate transporter BicA but does contain the two high-affinity bicarbonate uptake systems SbtA and BCT1 (see Section 2.2), grows at only 35% of its maximum growth rate on bicarbonate alone (Verspagen et al., 2014b). In chemostat experiments, this *Microcystis* strain could barely sustain a small

population when CO_2 was largely removed from the gas flow, even though bicarbonate was provided at a saturating concentration of 2000 $\mu\text{mol L}^{-1}$ (Verspagen et al., 2014b). An increase from near-zero $p\text{CO}_2$ levels (0.5 ppm) to saturating $p\text{CO}_2$ levels (2800 ppm) led to an almost 20-fold increase of the *Microcystis* biomass. These laboratory experiments show that addition of CO_2 may strongly

promote cyanobacterial growth even in bicarbonate-rich waters. Yet, other cyanobacterial species such as *Cylindrospermopsis raciborskii* appear to be more effective bicarbonate users, and for these species rising CO₂ concentrations may have a smaller effect on growth rates when bicarbonate is available as an alternative C_i source (Holland et al., 2012). Hence, the effect of rising CO₂ on cyanobacterial growth is species specific. Moreover, the next sections will show that there is tremendous variation in CO₂ response within species.

2.2. The CO₂-concentrating mechanism of cyanobacteria

Phytoplankton use CO₂ and bicarbonate available in the environment for carbon fixation with the RuBisCO enzyme. To overcome the low affinity of RuBisCO for CO₂, most phytoplankton, including cyanobacteria, evolved a CO₂-concentrating mechanism (CCM) (Kaplan and Reinhold, 1999; Giordano et al., 2005; Badger et al., 2006; Price et al., 2008; Price, 2011). The typical cyanobacterial CCM is based on the uptake of CO₂ and bicarbonate from the environment, conversion of the acquired CO₂ into bicarbonate in the cytoplasm, and subsequent diffusion of the accumulated bicarbonate into specialized compartments called carboxysomes (Fig. 3). In the carboxysomes, carbonic anhydrases convert the accumulated bicarbonate back to CO₂, surrounding RuBisCO by a high CO₂ concentration. RuBisCO incorporates CO₂ into the Calvin–Benson–Bassham cycle, which assimilates the acquired carbon into organic molecules.

In cyanobacteria, five different C_i uptake systems have been identified, three for the uptake of bicarbonate and two for the conversion of CO₂, that diffuses into the cell, to bicarbonate (Fig. 3). These uptake systems have different physiological properties (Price et al., 2004; Price, 2011; Sandrini et al., 2015b). Two of the bicarbonate transporters, BicA and SbtA, are sodium-dependent symporters (Shibata et al., 2002; Price et al., 2004). BicA has a low affinity for bicarbonate ($K_{0.5} = 70\text{--}350 \mu\text{M}$ bicarbonate), but high flux rate. Conversely, SbtA has a high affinity for bicarbonate ($K_{0.5} < 5 \mu\text{M}$ bicarbonate), but low flux rate (Price et al., 2004). The third bicarbonate transporter, BCT1, is ATP-dependent, and similar to SbtA it has a high affinity for bicarbonate ($K_{0.5} = 10\text{--}15 \mu\text{M}$ bicarbonate) but a low flux rate (Omata et al., 1999, 2002). All three

bicarbonate uptake systems are located in the plasma membrane (Price, 2011).

The two CO₂ uptake systems, NDH-I₃ and NDH-I₄, convert CO₂ that passively diffuses into the cell to bicarbonate in a NADPH-dependent reaction (Price et al., 2002; Price, 2011). NDH-I₃ has a high affinity for CO₂ ($K_{0.5} = 1\text{--}2 \mu\text{M}$ CO₂) but a low flux rate (Maeda et al., 2002; Price et al., 2002). Conversely, NDH-I₄ has a lower affinity for CO₂ ($K_{0.5} = 10\text{--}15 \mu\text{M}$ CO₂) but a high flux rate (Maeda et al., 2002; Price et al., 2002). This diverse array of C_i uptake systems enables cyanobacteria to respond effectively to changes in C_i availability.

Eukaryotic algae can also employ a CCM, but it works differently from the CCM of cyanobacteria. In the green alga *Chlamydomonas reinhardtii*, the CCM is based on a light-driven pH gradient that is set up across the chloroplast thylakoid membrane, converting bicarbonate transported into the thylakoid lumen into CO₂ near the pyrenoids where CO₂-fixation takes place (Moroney and Ynalvez, 2007; Moroney et al., 2011). An interesting selection experiment where *C. reinhardtii* was exposed to elevated CO₂ for 1000 generations revealed that some cell lines lost the ability to induce high-affinity CO₂ uptake (Collins and Bell, 2004; Collins et al., 2006). This was attributed to mutations in CCM genes. Hence, similar to cyanobacteria, *C. reinhardtii* likely also possesses high-affinity and low-affinity C_i uptake genes. This experiment demonstrates that eukaryotic algae can evolve in response to elevated CO₂. Yet, much less is known about the CCM genes and proteins of algae than those of cyanobacteria.

2.3. Genetic diversity of C_i uptake systems in *Microcystis*

Microcystis is a potentially toxic cyanobacterium that forms dense blooms in eutrophic lakes all over the world (Verspagen et al., 2006; Qin et al., 2010; Michalak et al., 2013), and can produce the hepatotoxin microcystin (Codd et al., 2005; Dittmann et al., 2012). The genomes of 20 strains of *Microcystis aeruginosa* (Kützing) (sensu Otsuka et al., 2001) were screened recently, which revealed that these strains differ in the combination of C_i uptake systems (Sandrini et al., 2014). Genes encoding the ATP-dependent bicarbonate transporter BCT, and the two CO₂ uptake systems NDH-I₃ and NDH-I₄ were found in all 20 strains. Most

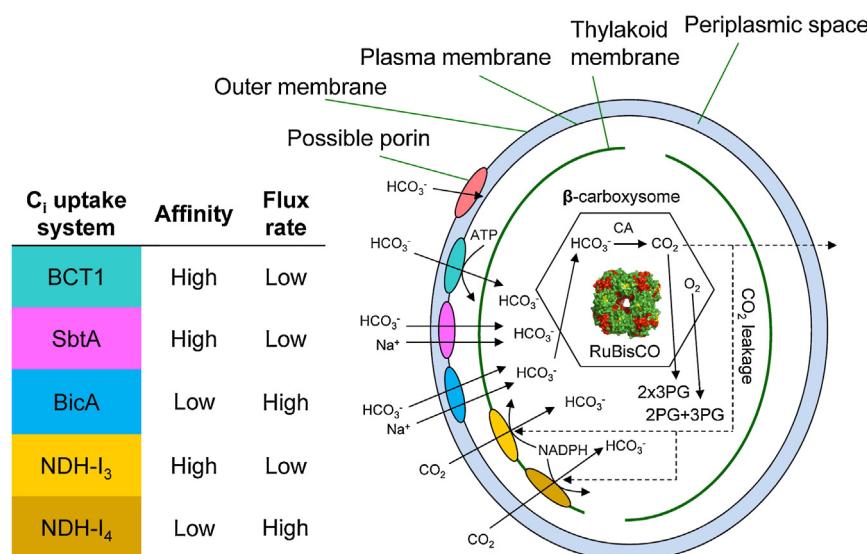


Fig. 3. Schematic overview of the CCM in cyanobacteria. Five different C_i uptake systems are known in cyanobacteria, including the ATP-dependent bicarbonate uptake system BCT1, two sodium-dependent bicarbonate uptake systems (BicA and SbtA) and two CO₂ uptake systems (NDH-I₃ and NDH-I₄). The C_i uptake systems differ in their affinities and flux rates. Accumulated bicarbonate is converted to CO₂ by carbonic anhydrases (CA) in the carboxysomes. CO₂-fixation by RuBisCO leads to the formation of 3-phosphoglycerate (3PG), whereas the reaction with O₂ (photorespiration) produces toxic 2-phosphoglycolate (2PG). The dashed lines indicate CO₂ leakage from the carboxysome, which can partly be intercepted by the CO₂ uptake systems.

Table 1The presence/absence of genes for C_i uptake systems and microcystin synthesis in cyanobacteria.

Strain	Origin	C _i uptake systems					Microcystin genes	
		Bicarbonate			CO ₂			
		BicA	SbtA	BCT1	NDH-I ₃	NDH-I ₄		
8 <i>Microcystis</i> strains	Sandrini et al. (2014)	+	+	+	+	+	+/-	
<i>Microcystis</i> PCC 7806	Sandrini et al. (2014)	+	-	+	+	+	+	
11 <i>Microcystis</i> strains	Sandrini et al. (2014)	-	+	+	+	+	+/-	
<i>Anabaena cylindrica</i> PCC 7122	GBR (Cambridge)	+	-	?	+	+	-	
<i>Anabaena</i> sp. 90	FIN (Lake Vesijarvi)	-	-	+	+	+	+	
<i>Anabaena</i> sp. PCC 7108	USA (Moss Beach, California)	-	+	+	+	+	-	
<i>Anabaena variabilis</i> ATCC 29413	USA (Mississippi)	+, ≠	+	+	+	+	-	
<i>Aphanizomenon flos-aquae</i> NIES-81	JPN (Lake Kasumigaura)	-	-	+	+	+	-	
<i>Planktothrix agardhii</i> NIVA-CYA 15	NOR (Lake Kolbotnvatnet)	-	+	+	+	+	+	
<i>Planktothrix agardhii</i> NIVA-CYA 34	NOR (Lake Kolbotnvatnet)	-	-	+	+	+	+	
<i>Planktothrix agardhii</i> NIVA-CYA 56/3	FIN (Lake Steinsfjorden)	-	+	+	+	+	+	
<i>Planktothrix prolifica</i> NIVA-CYA 98	FIN (Lake Steinsfjorden)	-	-	+	+	+	+	
<i>Planktothrix agardhii</i> NIVA-CYA 126/8	FIN (Lake Langsjon)	+	-	+	+	+	+	
<i>Planktothrix mougeotii</i> NIVA-CYA 405	FIN (Lake Steinsfjorden)	-	-	+	+	+	+	
<i>Planktothrix prolifica</i> NIVA-CYA 406	FIN (Lake Steinsfjorden)	-	-	+	+	+	+	
<i>Planktothrix rubescens</i> NIVA-CYA 407	FIN (Lake Steinsfjorden)	-	-	+	+	+	+	
<i>Planktothrix prolifica</i> NIVA-CYA 540	FIN (Lake Steinsfjorden)	-	+	+	+	+	+	
<i>Synechocystis</i> sp. PCC 6803	USA (California)	+	+	+	+	+	-	
<i>Synechococcus</i> sp. PCC 7002	PRI (Magueyes Island)	+	+	-	+	+	-	
<i>Synechococcus</i> sp. PCC 7942	USA (Texas)	-	+	+	+	+	-	

The C_i uptake systems SbtA, BCT1 and NDH-I₃ have a high substrate affinity and low flux rate, whereas BicA and NDH-I₄ have a low substrate affinity and high flux rate. The model cyanobacteria *Synechocystis* PCC 6803, *Synechococcus* PCC 7002 and *Synechococcus* PCC 7942 are shown for comparison with sequenced *Microcystis*, *Anabaena*, *Aphanizomenon*, and *Planktothrix* strains.

+ indicates that the gene is present. – indicates that the gene is absent. ≠ indicates that only a small fragment of the gene is present. ? indicates that a similar gene is present, but it is not clear if it encodes for the BCT1 bicarbonate transporter (*cmpABCD*), or possibly a different transporter.

The origins of the strains are indicated with three-letter codes of the different countries (ISO 3166-1 a-3).

The presence of C_i uptake genes is based on high similarity of the protein sequences with the reference protein sequences in *Microcystis* PCC 7806, *Microcystis* NIES-843, *Synechocystis* PCC 6803, *Synechococcus* PCC 7002 and *Synechococcus* PCC 7942. The presence of microcystin genes indicates potentially toxic strains.

other CCM genes are also widespread in *Microcystis*. However, *Microcystis* strains differ in the presence of the two sodium-dependent bicarbonate transporters BicA and SbtA. Three C_i uptake genotypes were found (Table 1). Some *Microcystis* strains possess all five C_i uptake systems and these are referred to as C_i uptake generalists. Moreover, in *Microcystis*, these C_i uptake generalists co-transcribe *bicA* and *sbtA* (Sandrini et al., 2014). Other strains contain the gene *sbtA* encoding for the high-affinity bicarbonate uptake system SbtA but lack the gene *bicA*, and will be referred to as high-affinity specialists. And again other strains contain the gene *bicA* encoding for the low-affinity but high-flux bicarbonate uptake system BicA, but lack the gene *sbtA*. These strains will be called high-flux specialists.

Eleven of the 20 investigated *Microcystis* strains produced the hepatotoxin microcystin. Microcystin-producing strains were found among the C_i uptake generalists, high-affinity specialists and high-flux specialists, and did not form distinct clusters in phylogenetic trees based on the *bicA* and *sbtAB* sequences (Sandrini et al., 2014). There is no relationship between the C_i uptake genotypes and the presence of microcystin production.

Within the C_i uptake genotypes, several genetic variants were discovered. For instance, one of the strains had a functional *sbtA* gene but a defective *bicA* gene caused by a transposon insert, and other strains combined *sbtA* with only a small remaining fragment of the *bicA* gene (Sandrini et al., 2014). These strains were classified among the high-affinity specialists, because their *bicA* gene is no longer functional. These results indicate that during the course of evolution some strains may have lost the ability to produce specific C_i uptake systems, in this case the loss of BicA. Presumably, in environments with low C_i availability the production of this low-affinity but high-flux bicarbonate transporter is an unnecessary burden, and its loss may therefore offer a selective advantage.

Consistent with these evolutionary considerations, laboratory experiments confirmed that the genetic variation in C_i uptake systems of *Microcystis* has phenotypic consequences (Sandrini

et al., 2014). High-affinity specialists with *sbtA* but without *bicA* grow better at a low partial pressure of CO₂ (pCO₂), but perform poorly at high pCO₂ conditions. Conversely, high-flux specialists with *bicA* but without *sbtA* grow poorly at low pCO₂, but perform well at high pCO₂ levels. Finally, C_i uptake generalists containing all five C_i uptake systems grow well across a wide range of pCO₂ levels (from 20 to 10,000 ppm) (Sandrini et al., 2014).

Competition experiments by Van de Waal et al. (2011) showed that rising pCO₂ levels can lead to a reversal in competitive dominance among *Microcystis* strains. These authors interpreted this result by differences in toxin production between the two strains, because one of the strains used in the experiments produced the hepatotoxin microcystin (strain CYA 140) whereas the other was non-toxic (strain CYA 43). The genetic analysis of Sandrini et al. (2014) revealed that these two strains also differed in their C_i uptake systems, which provides a much more parsimonious explanation for the observed reversal in competitive dominance. *Microcystis* strain CYA 140 was a high-affinity specialist (only *sbtA*), and won the competition at low pCO₂ levels. In contrast, *Microcystis* strain CYA 43 (=PCC 7005) was a C_i uptake generalist with both *bicA* and *sbtA*, and won the competition at high pCO₂ levels. While these experiments demonstrate that rising pCO₂ may shift strain dominance, this shift can thus be attributed to differences in the C_i uptake traits of the strains rather than to differences in their microcystin production. In particular, the results of these competition experiments support the hypothesis that natural selection favors the *sbtA* gene at low CO₂ conditions, whereas *bicA*-containing strains are favored at high CO₂ conditions.

2.4. C_i uptake systems of other harmful cyanobacteria

The CCMs of other harmful freshwater cyanobacteria have not been studied in detail, partly because genomic data are still largely lacking. But now the genomes of four *Anabaena* strains (Wang et al.,

2012; Shih et al., 2013; Thiel et al., 2014), one *Aphanizomenon* strain (Cao et al., 2014) and nine *Planktothrix* strains (Tooming-Klunderud et al., 2013; Christiansen et al., 2014) have been sequenced. We analyzed the CCM genes present in these genomes, based on high similarity of the protein sequences with the reference protein sequences from *Microcystis* PCC 7806, *Microcystis* NIES-843, *Synechocystis* PCC 6803, *Synechococcus* PCC 7002 and *Synechococcus* PCC 7942. This analysis revealed that *Anabaena*, *Aphanizomenon* and *Planktothrix* also display variation in the presence of the *bicA* and *sbtA* genes, whereas the two CO₂ uptake systems and the BCT1 bicarbonate transporter are widespread among all four cyanobacterial genera (Table 1). Interestingly, in addition to the three genotypes described in *Microcystis*, a fourth genotype that lacks both *bicA* and *sbtA* was detected in *Anabaena*, *Aphanizomenon* and *Planktothrix* (Table 1). Strains with this strategy might be called “C_i uptake minimalists”.

Hence, similar to *Microcystis*, other genera of harmful cyanobacteria also show genetic variation in their C_i uptake systems. Presumably, this genetic diversity produces a phenotypic variation similar to *Microcystis*, with a selective advantage for *sbtA*-containing strains at low CO₂ conditions but a selective advantage for *bicA*-containing strains in high-CO₂ environments. The phenotypic niche of *Anabaena*, *Aphanizomenon* and *Planktothrix* strains that lack both *bicA* and *sbtA* is intriguing, and has not yet been investigated. The absence of both sodium-bicarbonate symporters might imply that bicarbonate uptake has been taken over by the ATP-dependent bicarbonate transporter BCT1, as an adaptation to environments with low sodium concentrations. It is also possible that these C_i uptake minimalists are largely specialized in CO₂ uptake and have only a very limited capacity for bicarbonate uptake, and hence are mainly found in soft waters with pH < 6 where bicarbonate uptake is of little advantage.

It is often argued that cyanobacteria have a very effective CCM, and are therefore particularly strong competitors at low CO₂ levels in comparison to eukaryotic phytoplankton (Shapiro, 1990). If so, one might expect that low CO₂ concentrations will favor cyanobacteria, whereas eukaryotic phytoplankton tend to become more dominant at elevated CO₂ concentrations. A number of competition experiments between cyanobacteria and eukaryotic phytoplankton seems to support this hypothesis (Shapiro, 1997; Caraco and Miller, 1998; Low-Décarie et al., 2011, 2015). In other experiments, however, eukaryotic phytoplankton dominated over cyanobacteria at low CO₂ but not at elevated CO₂ concentrations (Verschoor et al., 2013). Indeed, the new insights reviewed above indicate that not all cyanobacteria are strong competitors at low CO₂. The genetic diversity of C_i uptake systems shows that there is major variation in the effectiveness of the cyanobacterial CCM, even among different strains within the same genus. Some cyanobacterial strains perform well at low CO₂, whereas other strains are much better competitors under high CO₂ conditions. This genetic and phenotypic variation in C_i uptake systems provides cyanobacterial communities with the potential for rapid evolutionary adaptation to changing CO₂ conditions, with a major selective advantage for cyanobacteria with high-flux C_i uptake systems in high-CO₂ environments.

3. Response to rising temperature

3.1. Enhanced growth rates

Rising temperatures promote cyanobacterial population dynamics in multiple ways. Temperatures of up to ~25 °C directly increase cyanobacterial photosynthesis and growth rate (Robarts and Zohary, 1987; Coles and Jones, 2000; Davis et al., 2009; Mehnert et al., 2010; Lürling et al., 2013). Most phytoplankton species reach their optimum temperature for growth in the range

of 20–25 °C, although there are exceptions, like the thermophilic cyanobacteria of hot springs (e.g., Allewalt et al., 2006) and stenotherm species (as suggested for *Planktothrix rubescens* (Findenegg, 1947)).

The general consensus is that cyanobacteria have a higher optimal growth temperature than eukaryotic algae. Paerl and Huisman (2008, 2009) based their ‘Blooms like it hot’ statement on experimental growth rate data of different species by Butterwick et al. (2005) and Reynolds (2006), and of seasonal phytoplankton data in a lake by Jöhnk et al. (2008). As a follow-up, Paerl et al. (2011) and Paerl (2014) showed literature data from several experimental studies, and their graphical compilation of these data also clearly indicate that the temperature optima of cyanobacteria are higher than those of most algae. The temperature optima of cyanobacteria were in the range of 27–37 °C overlapping those of green algae with optima in the range of 27–32 °C, while those of dinoflagellates (17–27 °C) and diatoms (17–22 °C) were distinctly different (Paerl, 2014). An experimental study with different species of cyanobacteria and green algae did not reveal a difference in optimum temperatures between these two taxonomic groups (Lürling et al., 2013). An extensive overview of temperature-dependent growth rates from several other studies yielded a slightly higher optimal temperature for cyanobacteria (27.2 °C) than for green algae (26.3 °C), but this difference was not significant (Lürling et al., 2013). Summarizing, the temperature optima for cyanobacteria and green algae can overlap, and this is likely dependent on species and culture conditions. Overall, the difference in temperature optima between cyanobacteria and green algae on one side and dinoflagellates and diatoms on the other appears to be considerable.

In addition to temperature optima, it is of interest to investigate how fast the growth rate increases with temperature. The temperature dependence of the growth rates of species has gained much interest in the context of the metabolic theory of ecology (Gillooly et al., 2001; Brown et al., 2004), and can be calculated from the Arrhenius equation:

$$\mu = c \exp\left(-\frac{E_A}{kT}\right) \quad (2)$$

Here, μ is the growth rate, c is a normalization constant, E_A is the activation energy, k is Boltzmann's constant (8.62×10^{-5} eV K⁻¹), and T is absolute temperature in Kelvin. The activation energy is a measure of the increase of the growth rate with temperature (below the temperature optimum). It can be estimated from an Arrhenius plot, where the natural logarithm of the growth rates ($\ln \mu$) is plotted against the inverse of temperature ($1/kT$). The value of the activation energy is then obtained from the (negative) slope of a linear regression of $\ln \mu$ versus $1/kT$. This approach was applied to the growth data of Lürling et al. (2013). For the cyanobacterium *Aphanizomenon gracile*, for example, this yields an activation energy of $E_A = 0.64$ eV while for the green alga *Scenedesmus acuminatus* it yields 0.11 eV (Fig. 4; Table 2).

For the cyanobacterial species investigated by Lürling et al. (2013), E_A ranged from 0.50 to 1.23 eV (Table 2). On average, the E_A (\pm s.d.) of the cyanobacteria was 0.70 (\pm 0.35) eV, whereas that of green algae was 0.43 (\pm 0.25) eV. This indicates that the growth rate of cyanobacteria increases faster with temperature than that of green algae, although the variation among species is considerable, and the difference between cyanobacteria and green algae was therefore at best marginally significant (Two-sample Student's *t*-test (for equal variances), $df = 14$, $p = 0.096$). To facilitate comparison with the literature, Q_{10} values (which measure the change in growth rate for a temperature increase of 10 °C) were also calculated from the data of Lürling et al. (2013). Over the temperature range from 20 to 27.5 °C, this gave Q_{10} values of 2.63 ± 0.94 for cyanobacteria and 2.03 ± 1.02

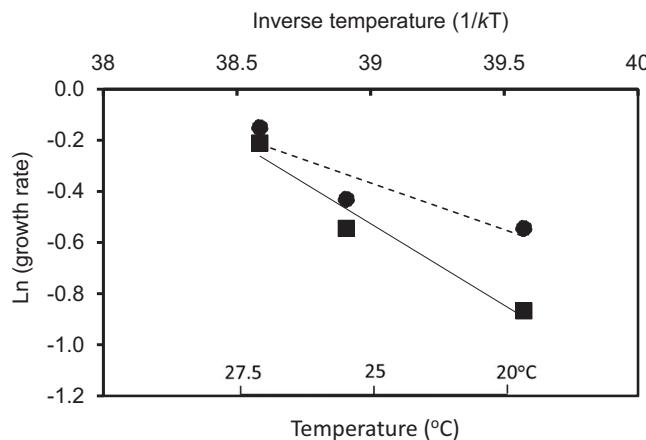


Fig. 4. Effect of temperature (expressed as $1/(kT)$, where k is Boltzmann's constant and T is absolute temperature in degrees Kelvin) on the growth rate of the cyanobacterium *Aphanizomenon gracile* (squares) and the green alga *Scenedesmus acuminatus* UTEX415 (circles) as determined in batch cultures (data from Lürling et al., 2013).

Table 2

Activation energies (E_A) and Q_{10} values of the growth rates of eight cyanobacteria and eight green algae. The values are calculated from the increase of growth rate with temperature, in the range of 20–27.5 °C, based on data from Lürling et al. (2013).

Species	E_A	Q_{10}
Cyanobacteria		
<i>Anabaena</i> sp. PCC7122	0.58	1.92
<i>Aphanizomenon gracile</i>	0.64	2.40
<i>Cylindrospermopsis raciborskii</i> CIRF-01	0.75	2.56
<i>Microcystis aeruginosa</i> PCC7941	0.54	2.21
<i>Microcystis aeruginosa</i> CYA140	1.23	4.63
<i>Planktothrix agardhii</i> CYA116	0.51	1.93
<i>Planktothrix agardhii</i> CYA126	0.50	
<i>Synechococcus elongatus</i> PCC6301	0.83	2.75
Green algae		
<i>Ankistrodesmus falcatus</i> CHL8	1.03	1.35
<i>Chlamydomonas reinhardtii</i> CHL13	0.85	1.26
<i>Desmodesmus bicellularis</i> CCAP276/14	0.37	1.92
<i>Desmodesmus quadricauda</i> UTEX614	0.46	2.99
<i>Monoraphidium minutum</i>	0.21	1.83
<i>Scenedesmus acuminatus</i> UTEX415	0.11	1.69
<i>Scenedesmus maximus</i> SAG39.81	0.37	4.09
<i>Scenedesmus obliquus</i> SAG276/3a	0.05	1.10

for green algae (Table 2). Other studies also reported high Q_{10} values for cyanobacteria, e.g., a study on seven cyanobacterial species by Mehnert et al. (2010) showed an average Q_{10} of 2.33 ± 0.87 . An exceptionally high Q_{10} (~ 9.6) for the growth-temperature dependence of *Microcystis* was reported by Reynolds (2006).

To unravel and understand these differences in temperature sensitivity between cyanobacteria and eukaryotic algae will require further study of the temperature dependencies of the different underlying physiological processes affecting phototrophic growth (e.g., carbon fixation, photorespiration and respiration; Bernacchi et al., 2001; Allen et al., 2005).

3.2. Increased stability of the water column

Temperature not only has a direct effect on the growth rates of organisms, but also has important indirect effects that can tip the competitive balance between species. In particular, high temperatures increase the stability of the water column, which suppresses turbulent mixing (Peeters et al., 2007; Jöhnk et al., 2008) and extends the duration of thermal stratification (De Stasio et al., 1996; Peeters et al., 2007). Thermal stratification is favorable for

buoyant cyanobacteria because weak mixing allows buoyant cyanobacteria to float to the upper water layers, where they have better access to light while shading the non-buoyant phytoplankton below (Reynolds et al., 1987; Huisman et al., 2004; Wagner and Adrian, 2009). Several cyanobacterial species contain gas vesicles, which are hollow protein structures in the cell that are filled with gas (Walsby, 1994). If the buoyancy provided by gas vesicles exceeds the ballast provided by other cell constituents, the cells float upwards.

Competition between buoyant cyanobacteria and non-buoyant eukaryotic phytoplankton can be described by a simple model. Consider a highly eutrophic or hypertrophic lake, in which nutrients are in ample supply and phytoplankton species such as diatoms, green algae and buoyant cyanobacteria mainly interact through mutual shading (i.e., competition for light). Let $N_i(z,t)$ denote the concentration of phytoplankton species i at depth z and time t . The population dynamics and vertical distribution of a number of n phytoplankton species can be described by the following set of partial differential equations (Klausmeier and Litchman, 2001; Huisman et al., 2004):

$$\frac{\partial N_i}{\partial t} = \mu_i(I(z))N_i + v_i \frac{\partial N_i}{\partial z} + D \frac{\partial^2 N_i}{\partial z^2} \quad i = 1, \dots, n \quad (3)$$

Here, $\mu_i(I(z))$ is the net specific growth rate of species i as function of light intensity I , where light intensity I decreases exponentially with depth z according to Lambert–Beer's law. Moreover, the light intensity at a given depth also depends on the phytoplankton concentrations above that depth, as denser phytoplankton blooms create more turbid conditions. For notational convenience, it was assumed that the net specific growth rate also includes phytoplankton losses due to, e.g., natural mortality, viral lysis and zooplankton grazing. Furthermore, v_i is the vertical sinking or flotation velocity of species i (with $v_i < 0$ for sinking species and $v_i > 0$ for buoyant species), and D is known as the turbulent diffusion coefficient or vertical eddy diffusivity. Zero-flux boundary conditions were assumed.

The model predicts that changes in the turbulent diffusion coefficient D , which is a measure of the mixing intensity, can cause a shift in the outcome of competition (Fig. 5; Huisman et al., 2004). In turbulent waters, vertical mixing dominates over the flotation

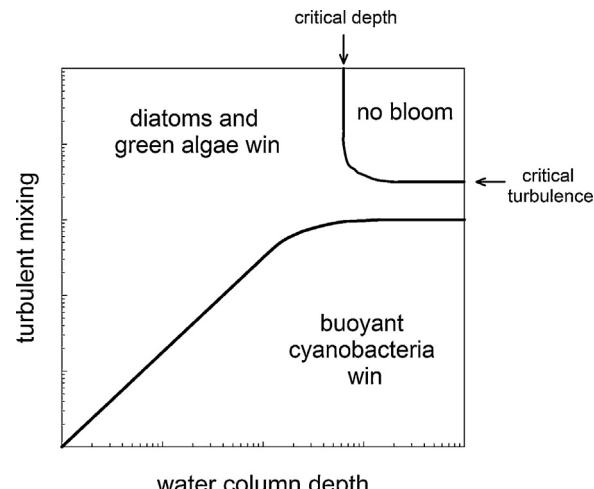


Fig. 5. Model prediction of competition between buoyant cyanobacteria and sinking diatoms and green algae, as a function of water-column depth and turbulent diffusion. The boundary line between the region of cyanobacterial dominance and the region where diatoms and green algae win depends on the ratio of the time scale of turbulent mixing and the vertical flotation velocity of the buoyant cyanobacteria. Hence, the exact position of this boundary line will vary among different species of cyanobacteria. For details, see Huisman et al. (2004).

and sinking rates of the different species. In this case, buoyant cyanobacteria cannot stay in the upper illuminated layers, because they are dispersed throughout the water column by intense vertical mixing. Under these conditions, diatoms and green algae are often superior competitors and displace buoyant cyanobacteria (Visser et al., 1996; Huisman et al., 2004). In contrast, in waters with only weak vertical mixing the flotation rate exceeds the mixing intensity, and buoyant cyanobacteria can float upwards and shade all non-buoyant species below. Moreover, in such stagnant waters diatoms and green algae are not entrained but may sink out of the surface layer. Hence, the model predicts that buoyant cyanobacteria displace the diatoms and green algae and dominate the phytoplankton community in waters with weak vertical mixing (Fig. 5).

All three terms in Eq. (3) are dependent on temperature. First, the specific growth rates ($\mu_i(I)$) of the different species vary with temperature. As seen in the preceding section, the growth rates of cyanobacteria tend to respond more strongly to rising temperatures than the growth rates of most eukaryotic phytoplankton. Second, according to Stokes' Law, the vertical flotation and sinking velocities (v_i) of the different species are inversely proportional to the viscosity of water. Since water becomes less viscous at higher temperature, this implies that buoyant cyanobacteria will float upwards faster whereas sinking diatoms will sink faster at higher temperature. Third, higher temperatures favor vertical stratification, which increases the stability of the water column and thereby suppresses vertical mixing by turbulent diffusion (D). Thus, higher temperatures lead to enhanced growth, a higher flotation velocity and a more stable water column, which all tend to favor the development of surface blooms of buoyant cyanobacteria at the expense of sinking diatoms and green algae.

The model predictions are supported by field studies that manipulated the turbulence structure of eutrophic lakes by artificial mixing. In these lake experiments, buoyant cyanobacteria were dominant when the water was stratified, but they were replaced by non-buoyant eukaryotic phytoplankton after the onset of artificial mixing (Reynolds et al., 1983; Visser et al., 1996; Huisman et al., 2004). An experiment in Lake Nieuwe Meer, The Netherlands, investigated an intermittent mixing regime, in which artificial mixing of the entire lake was alternately switched on and off at two-week intervals (Jöhnk et al., 2008). The experiment was performed during the summer of 2003, which turned out to be one of the hottest summers ever recorded in Europe. The increases and decreases in cell number of the buoyant cyanobacterium *Microcystis* during the summer heatwave coincided with the alternations in mixing intensity. When artificial mixing was switched off, the water column stratified and the *Microcystis* concentration in the surface layer rapidly increased both by vertical migration of the colonies and by population growth. And when artificial mixing was switched on again, the water column was well mixed and *Microcystis* declined. These results demonstrate that, in this lake, the direct effect of high temperatures on cyanobacterial growth rates was by itself not sufficient for cyanobacterial dominance. In addition, high temperatures also increased the stability of the water column, which enhanced the ability of buoyant cyanobacteria to float upwards and shift the competitive balance in their favor (Jöhnk et al., 2008).

3.3. Extension of growing season

Global warming may also affect the annual life cycle of cyanobacteria. Some cyanobacterial species like *Planktothrix agardhii* and *Planktothrix rubescens* form winter blooms or are persistent in the water column throughout the year in temperate lakes (Naselli-Flores et al., 2007; Akcaalan et al., 2014; Anneville et al., 2015). Most cyanobacterial species decrease in abundance

in winter, and several species overwinter as akinetes (specialized cells resistant to cold, desiccation and irradiation) or vegetative cells in lake sediments. N₂-fixing heterocystous cyanobacteria such as *Anabaena*, *Aphanizomenon* and *Gloeotrichia* may form akinetes (Sukenik et al., 2012, 2013; Cirés et al., 2013), whereas *Microcystis* colonies overwinter as vegetative cells (Brunberg and Boström, 1992; Misson et al., 2012). The overwintering populations of cyanobacteria in the sediment provide a potential inoculum for spring or summer blooms (Brunberg and Blomqvist, 2003; Kravchuk et al., 2011; Cirés et al., 2013). Model simulations indicate that the absence of recruitment from the sediment would decrease the subsequent summer bloom of *Microcystis* by about 50% (Verspagen et al., 2005). Increasing temperatures may initiate earlier germination of akinetes (Tsujimura and Okubo, 2003; Carey et al., 2014) and recruitment of *Microcystis* colonies (Trimbee and Prepas, 1988; Karlsson-Elfgren et al., 2004; Cao et al., 2008).

In temperate lakes, *Microcystis* has a bloom period in August–September, after which the population settles to the lake sediment (Reynolds et al., 1981; Takamura et al., 1984; Thomas and Walsby, 1986). This loss of buoyancy in autumn has been explained by accumulation of carbohydrates at decreasing temperatures (Thomas and Walsby, 1986; Visser et al., 1995). Akinete formation by filamentous cyanobacteria is often induced by the onset of physiological stress, e.g., by phosphate limitation, light limitation, or decreasing temperature (Sinclair and Whitton, 1977; Adams and Duggan, 1999; Meeks et al., 2002).

Since temperature is an important driver of recruitment from the sediment in spring and summer, of subsequent population growth, and of the initiation of the benthic life stages at the end of the season, global warming will likely cause an earlier onset and later cessation of cyanobacterial blooms. As a result, climate change may extend the growing season considerably. In Scandinavian lakes, for example, warmer winters and springs have increased spring and early summer biomass of cyanobacteria (Weyhenmeyer, 2001).

4. Interactive effects with nutrient availability

4.1. Effects of climate change depend on nutrient availability

In many aquatic systems the availability of nutrients determines primary production (Dzialowski et al., 2005; Xu et al., 2010; Lewandowska et al., 2014), and total nitrogen and total phosphorus concentrations are often good predictors of cyanobacterial biomass (Downing et al., 2001; Håkanson et al., 2007). At the physiological level, there are still many gaps in our understanding of how nutrient limitation may interact with changes in temperature or CO₂ availability (e.g., Spijkerman et al., 2011).

Verspagen et al. (2014a) developed a conceptual framework to predict how different nutrient loads may modify effects of rising CO₂ on phytoplankton biomass production. They investigated a stoichiometrically explicit model that describes phytoplankton growth as function of nutrient, CO₂ and light availability. Hence, there are three potentially limiting resources in this model (Fig. 6A). Inorganic carbon becomes limiting at very low pCO₂ levels, nutrients become limiting at very low nutrient loads, and light becomes limiting in dense phytoplankton blooms at high pCO₂ levels and high nutrient loads. Light limitation can of course also be induced by other mechanisms, such as a high background turbidity (due to high concentrations of dissolved organic matter or resuspended sediment particles), deep mixing, or low incident light intensities in winter. The resource limitation pattern in Fig. 6A can be used to sketch to what extent rising pCO₂ levels will increase phytoplankton biomass (Fig. 6B). In oligotrophic waters with low nutrient loads, rising pCO₂ levels will shift phytoplankton

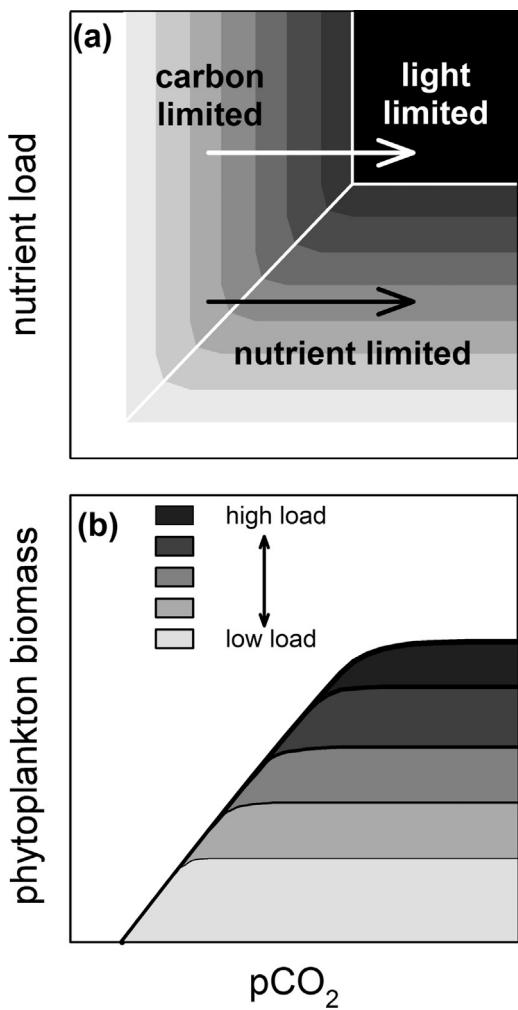


Fig. 6. (A) Hypothesized patterns of resource limitation, at different atmospheric CO₂ levels and nutrient loads. The arrows indicate that rising atmospheric CO₂ levels will cause a shift from carbon to nutrient limitation in systems with a low nutrient load (black arrow), but from carbon to light limitation in systems with a high nutrient load (white arrow). (B) The extent to which phytoplankton biomass will increase with rising CO₂ levels will depend on the nutrient load. Adjusted from Verspagen et al. (2014a).

growth from carbon-limited to nutrient-limited conditions (black arrow in Fig. 6A). In this case, the higher CO₂ availability stimulates some growth, but the phytoplankton biomass remains constrained by the low nutrient levels in the system (Fig. 6B). Conversely, in hypertrophic waters with high nutrient loads, rising pCO₂ levels will shift phytoplankton growth from carbon-to light-limited conditions (white arrow in Fig. 6A), which will allow a much larger increase in phytoplankton biomass (Fig. 6B). Chemostat experiments with the harmful cyanobacteria *Microcystis* CYA140 and HUB5-2-4 confirmed these model predictions (Verspagen et al., 2014a). The general message emerging from these results is that phytoplankton biomass will respond more strongly to rising pCO₂ levels in eutrophic and hypertrophic than in oligotrophic ecosystems.

Similarly, the response of cyanobacteria to temperature depends strongly on nutrient availability (Davis et al., 2009; Wagner and Adrian, 2009; Kosten et al., 2012; Tararu et al., 2012; Beaulieu et al., 2013; Rigosi et al., 2014). For instance, two recent studies analyzed data on cyanobacterial biomass in more than 1000 lakes in the USA (Beaulieu et al., 2013; Rigosi et al., 2014). They found that the cyanobacterial biomass in these lakes was

affected by both nutrients and temperature, where nutrients explained a larger proportion of the variation than temperature. Moreover, the relative importance of nutrients and temperature varied with the nutrient status of the lakes. Nutrient availability had a large impact on cyanobacterial biomass in oligotrophic lakes, whereas temperature was more important in mesotrophic lakes. In eutrophic and hyper-eutrophic lakes, nutrients and temperature had a synergistic effect on cyanobacterial biomass (Rigosi et al., 2014). Synergy of nutrients and temperature implies that in a warmer climate, nutrient concentrations will have to be decreased further in order to control cyanobacterial biomass (Kosten et al., 2012). Furthermore, analysis of this large data set revealed that the response to nutrients and temperature varied among the different cyanobacterial taxa. Some taxa, such as *Anabaena*, are more sensitive to changes in nutrient availability, whereas other taxa, such as *Microcystis*, are more sensitive to changes in temperature (Rigosi et al., 2014).

4.2. Effects of climate change on nitrogen fixation

Several genera of harmful cyanobacteria are capable of fixing atmospheric dinitrogen (N₂), including *Anabaena* (nowadays referred to as *Dolichospermum*; Wacklin et al., 2009), *Aphanizomenon*, *Cylindrospermopsis*, *Nodularia*, *Lyngbya* and *Nostoc*. In contrast, other harmful cyanobacterial genera such as *Microcystis* and *Planktothrix* cannot fix N₂. Nitrogen fixation is carried out by nitrogenase (Zehr et al., 2000). This enzyme complex is inhibited by oxygen (Gallion, 1992). Since photosynthesis produces oxygen, cyanobacteria need special adaptations to protect nitrogenase from oxygen inactivation. In freshwater ecosystems, diazotrophic cyanobacteria have developed a spatial separation of photosynthesis and nitrogen fixation by differentiating special cells, known as heterocysts (Muro-Pastor and Hess, 2012). Heterocysts have a thick glycolipid cell wall. This cell wall serves as a gas diffusion barrier to decrease the diffusion of O₂ into the heterocyst, but should of course allow a sufficient influx of N₂ to enable nitrogen fixation (Walsby, 1985). Hence, the glycolipid layer is not impermeable for gases. The O₂ that unavoidably diffuses into the heterocyst should be respiration fast enough to obtain near anoxic conditions inside the heterocyst.

Recently, theory has been developed that uses the different temperature dependencies of O₂ diffusion and respiration to predict how rising temperature will affect nitrogen fixation (Stal, 2009; Brauer et al., 2013). Diffusion of O₂ into the cell depends on the O₂ concentration in the surrounding medium, temperature, and the diffusion properties of the cell wall. All else being equal, the O₂ diffusion rate slightly increases with temperature at a Q₁₀ of ~1.1; i.e., with every 10 °C increase of temperature the influx of O₂ into the heterocyst increases only with a factor 1.1 (Stal, 2009). Respiration is an enzymatic process that increases (within its physiological limits) with temperature at a Q₁₀ of ~2; i.e. the respiration rate doubles with each 10 °C increase of temperature. The consequence is that at high temperature, the respiration rate may easily keep up with the diffusive influx of O₂. At low temperature, when respiration is slow, the influx of O₂ by diffusion may exceed the respiration rate and the heterocyst has difficulty to maintain anaerobic conditions. At low temperature, the glycolipid cell wall must therefore be more efficient as a gas diffusion barrier. For instance, in the cold waters of the northern Baltic Sea, heterocysts of *Anabaena* sp. possess a much more efficient gas diffusion barrier than those of *Nodularia spumigena* in the warmer waters of the Bornholm Sea in the southern Baltic (Stal, 2009). The downside of a more efficient gas diffusion barrier is a lower influx of N₂ into the heterocysts, thus suppressing their nitrogen fixation activity. The *Nodularia* blooms in the southern Baltic are indeed more productive in

terms of biomass and N₂-fixation activity than the *Anabaena* blooms in the northern Baltic.

These theoretical considerations suggest that increasing temperature will favor heterocysts with glycolipid cell walls that are less rigorous N₂ diffusion barriers and presumably less 'expensive' to synthesize. Therefore, it is likely that increasing temperature will result in a substantially higher N₂-fixation activity and greater competitive advantage for diazotrophic cyanobacteria in nitrogen-limited waters.

Rising CO₂ concentrations may also enhance nitrogen fixation rates, as has been reported for non-heterocystous marine cyanobacteria such as *Trichodesmium* and *Crocospaera* ([Hutchins et al., 2007](#); [Levitin et al., 2007](#); [Fu et al., 2008](#)). There are large strain-specific differences in CO₂ response, however, suggesting that individual strains of these diazotrophs are adapted to grow and fix nitrogen at different CO₂ concentrations ([Hutchins et al., 2013](#)). These strain-specific differences might again be related to variation in the presence and expression of different C_i uptake systems, similar to the genetic and phenotypic variation in CO₂ responses of *Microcystis* ([Sandrin et al., 2014, 2015b](#)). For example, *Trichodesmium erythraeum* IMS101 possesses only the high-flux CO₂ uptake system NDH-I₄ and the high-flux bicarbonate uptake system BicA ([Kranz et al., 2011](#)), which may explain why the nitrogen fixation and growth rate of this high-flux specialist increase strongly with a rise in ambient pCO₂ levels ([Hutchins et al., 2007](#); [Levitin et al., 2007](#); [Kranz et al., 2009](#)).

Nitrogen fixation rates in *Nodularia spumigena*, a heterocystous diazotroph from the Baltic Sea, showed contrasting CO₂ responses in different laboratory experiments ([Czerny et al., 2009](#); [Wannicke et al., 2012](#); [Eichner et al., 2014](#)). Part of this variation might be attributed to different growth conditions, as some experiments were performed under phosphate-limited conditions ([Wannicke et al., 2012](#)) whereas others used phosphate-replete conditions ([Czerny et al., 2009](#); [Eichner et al., 2014](#)). Mesocosms with natural phytoplankton assemblages from the Baltic Sea, including *Nodularia* and *Aphanizomenon* species, did not reveal any significant change in N₂-fixation activity in response to elevated CO₂ ([Paul et al., 2015](#)). To what extent the N₂-fixation rates of harmful cyanobacteria in lakes and reservoirs will respond to rising CO₂ is still largely an open question. Comparative studies of the genetic and phenotypic variation in CO₂ responses among diazotrophs may shed more light on this important gap in our knowledge.

5. Effects of climate change on cyanobacterial toxins

Cyanobacteria produce a range of bioactive compounds ([Welker and Von Döhren, 2006](#); [Leão et al., 2012](#)). Microcystins are the most well-known and most abundant ones in lakes and are toxic to animals ([Metcalf and Codd, 2012](#)). In predicting the effects of climate change on microcystin concentrations in a lake, one should focus on the effect of environmental conditions on: (1) cyanobacterial biomass, (2) the ratio of toxic (microcystin-producing) to non-toxic cyanobacteria, and (3) the microcystin production per cell.

Variation in cyanobacterial biomass causes the highest variation in cyanotoxin concentration in aquatic ecosystems: the more cyanobacterial biomass, the higher the toxin concentration. Cyanobacterial biomass is affected by CO₂, temperature, nutrients, and light, as has been described in the preceding sections when impacts on cyanobacterial growth were discussed.

The ratio of toxic to non-toxic strains is also a major determinant of the microcystin concentration in lakes ([Kardinaal and Visser, 2005](#)). [Davis et al. \(2009\)](#) found that during field experiments in four lakes in the northeast USA, toxic strains of *Microcystis* grew faster than their non-toxic counterparts when

water temperatures were increased 4 °C above ambient (average of the four lakes was 24 °C). Furthermore, they found that the interaction of increasing temperature and nutrients produced the highest growth rates in toxic strains, potentially leading to larger blooms with higher toxin contents.

Changes in microcystin production can be responsible for up to a fourfold variation of the microcystin content per cell ([Wiedner et al., 2003](#); [Kardinaal and Visser, 2005](#); [Van de Waal et al., 2009](#)). Many studies have investigated the impact of environmental variables on the microcystin production of toxic cells. The review of [Gehringer and Wannicke \(2014\)](#) indicates that microcystin production is stimulated by an ample supply of nutrients in combination with suitable temperature and light conditions for optimal growth. Under nutrient-rich conditions, elevated CO₂ levels stimulate a further increase of the microcystin content in *Microcystis* cells ([Van de Waal et al., 2009](#); [Sandrin et al., 2015a](#)). Furthermore, in a strain producing several different microcystin variants, elevated CO₂ levels in combination with high nitrogen concentrations shifted the microcystin composition towards the more N-rich but less toxic variant microcystin-RR ([Van de Waal et al., 2009](#)).

Almost all previous research on the effects of environmental conditions on microcystin production (reviewed by, e.g., [Sivonen and Jones, 1999](#); [Gehringer and Wannicke, 2014](#)) has been performed on free microcystins in the cells, while it is now known that a large fraction is covalently bound to proteins ([Zilliges et al., 2011](#); [Meissner et al., 2013, 2015](#)). These bound microcystins cannot be extracted using methanol. The fraction of bound microcystins is variable and dependent on the environmental conditions, e.g., the binding to proteins is associated with oxidative stress caused by high light ([Meissner et al., 2013](#)). This raises questions regarding the validity of previous studies as well as the potential toxicity of bound microcystins. Further research on the binding of microcystins to proteins is therefore recommended.

Cyanobacteria can also produce a variety of other toxins, including the hepatotoxins nodularin and cylindrospermopsin and the neurotoxins anatoxin and saxitoxin. These cyanotoxins are less widespread than microcystin, and only a few studies have investigated how their production is affected by environmental conditions (reviewed by [Neilan et al., 2013](#); [Boopathi and Ki, 2014](#)). The available studies indicate that nodularin production by *Nodularia spumigena* was stimulated at elevated temperature ([Lehtimäki et al., 1997](#); [Hobson and Fallowfield, 2003](#)). Saxitoxin production by *Aphanizomenon* sp. LMECYA was higher at 28 °C than at 22 °C ([Dias et al., 2002](#)), but saxitoxin production by *Cylindrospermopsis raciborskii* strain C10 was lower at 25 °C than at 19 °C ([Castro et al., 2004](#)). Anatoxin production by *Anabaena* and *Aphanizomenon* decreased at high temperature ([Rapala et al., 1993](#)). Hence, each of these cyanotoxins shows a different temperature response, which indicates that rising temperatures may alter the toxin composition of cyanobacterial blooms.

6. Future research needs and conclusions

One of the key points emphasized in this review is that dissolved inorganic carbon concentrations in eutrophic lakes can change dramatically on seasonal time scales, from supersaturation in winter to undersaturation in summer. Yet, the possible impacts of rising atmospheric CO₂ levels on freshwater ecosystems have received surprisingly little attention thus far. Models and laboratory experiments provide arguments that rising CO₂ levels are likely to stimulate cyanobacterial blooms. Field evidence is still limited, however, and the extent to which cyanobacterial blooms can sequester atmospheric CO₂ is still largely unexplored. Hence, there is a need for lake studies on the coupling of cyanobacterial blooms with seasonal and diurnal dynamics of the dissolved

inorganic carbon, and how these dynamics interact with exchanges of CO₂ with the atmosphere.

Furthermore, during recent years much more has become known about the molecular functioning and genetic diversity of cyanobacterial CCMs, both in model cyanobacteria such as *Synechocystis* PCC 6803 (Price, 2011; Burnap et al., 2015) and in environmentally relevant cyanobacteria such as *Microcystis* (Sandrin et al., 2014, 2015b). Little is known about the abundance, succession and geographical distribution of different C_i uptake genotypes in natural waters, or about evolutionary adaptation of cyanobacterial CCMs following prolonged exposure to elevated CO₂ concentrations. Hence, there is a need for biogeographical and eco-evolutionary studies investigating adaptive responses of cyanobacteria to changes in CO₂ availability.

Although effects of environmental conditions on microcystin production in *Microcystis* have been extensively investigated, there are many other toxins produced by many other species that have yet to be examined. Furthermore, the toxin concentrations in cyanobacteria-dominated lakes are largely determined by the relative abundances of toxic versus non-toxic strains. Yet, only a few studies have investigated how the competition between toxic and non-toxic strains is altered at elevated temperature (Davis et al., 2009) and elevated CO₂ (Van de Waal et al., 2011). Hence, there is a need for studies assessing how climate change will affect the toxicity of cyanobacterial blooms, and in particular under which circumstances toxic strains are able to outperform non-toxic strains and vice versa.

Cyanobacteria and eukaryotic algae may respond differently to climate change, which can lead to large changes in phytoplankton community composition. Yet, only a few studies have compared growth responses to temperature or CO₂ across a wide range of species (e.g., Butterwick et al., 2005; Lürling et al., 2013). The available studies provide little information on the impact of limiting resources (N, P, carbon, light) on the temperature-growth responses, and possible synergistic effects of rising CO₂ and elevated temperature have rarely been investigated (Fu et al., 2007; Karlberg and Wulff, 2013). To understand changes in community composition, there is a great need for controlled studies that compare growth responses to rising CO₂ and global warming across different species.

Comparative lake data have been analyzed to study the impact of temperature on cyanobacterial dominance across large geographical gradients (Kosten et al., 2012; Tararu et al., 2012; Beaulieu et al., 2013; Rigosi et al., 2014). To predict the impact of rising CO₂ concentrations, similar comparative lake studies should be carried out that focus on CO₂ dynamics and pH in relation to phytoplankton community composition. Furthermore, there is a particular need for long-term lake studies, so that changes over time can be quantified.

In conclusion, the effects of climate change on cyanobacteria are multifaceted and can be quite complex. There is broad consensus in the scientific literature that rising atmospheric CO₂ concentrations and global warming are likely to increase the occurrence, intensity and duration of harmful cyanobacterial blooms in eutrophic lakes. Additionally, the microcystin production of cyanobacteria will probably increase at elevated temperature and high CO₂ levels. There are still many intriguing open questions and uncertainties. Hence, there is a clear need for more laboratory and field research across a range of spatiotemporal scales. The risk that changes in climate and land use will cause a further deterioration of the water quality in many areas of the world generates a societal responsibility for scientists, water managers and policy makers to take further steps in our ability to understand, predict and mitigate the occurrence of toxic blooms in the surface waters across the changing landscapes of our planet.

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